

Pebble accretion near the snowline

Chris Ormel

University of Amsterdam
VIDI research group leader

w/ input from
Sebastiaan Krijt, Djoeke
Schoonenberg, Beibei Liu



Key questions/contents

- **Which are planet's building blocks**
kilometer size **planetesimals** or small **pebbles**?
Planetesimals form at H₂O iceline
- **Conditions for Pebble Accretion**
P.A. relies on planetesimals as “seeds”
- **Pebble accretion efficiency**
Dependence on position, stellar mass; effects of H₂O iceline
- **Implications to planets**
Planet formation solar system vs. low mass stars
Effects of P.A. on composition planets (rocky vs icy)

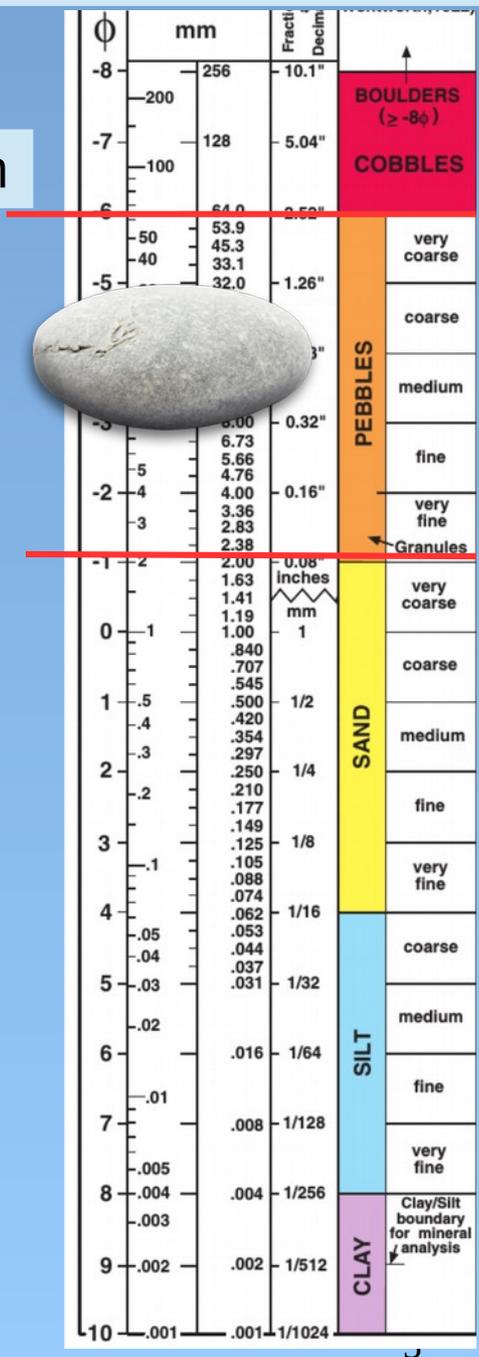
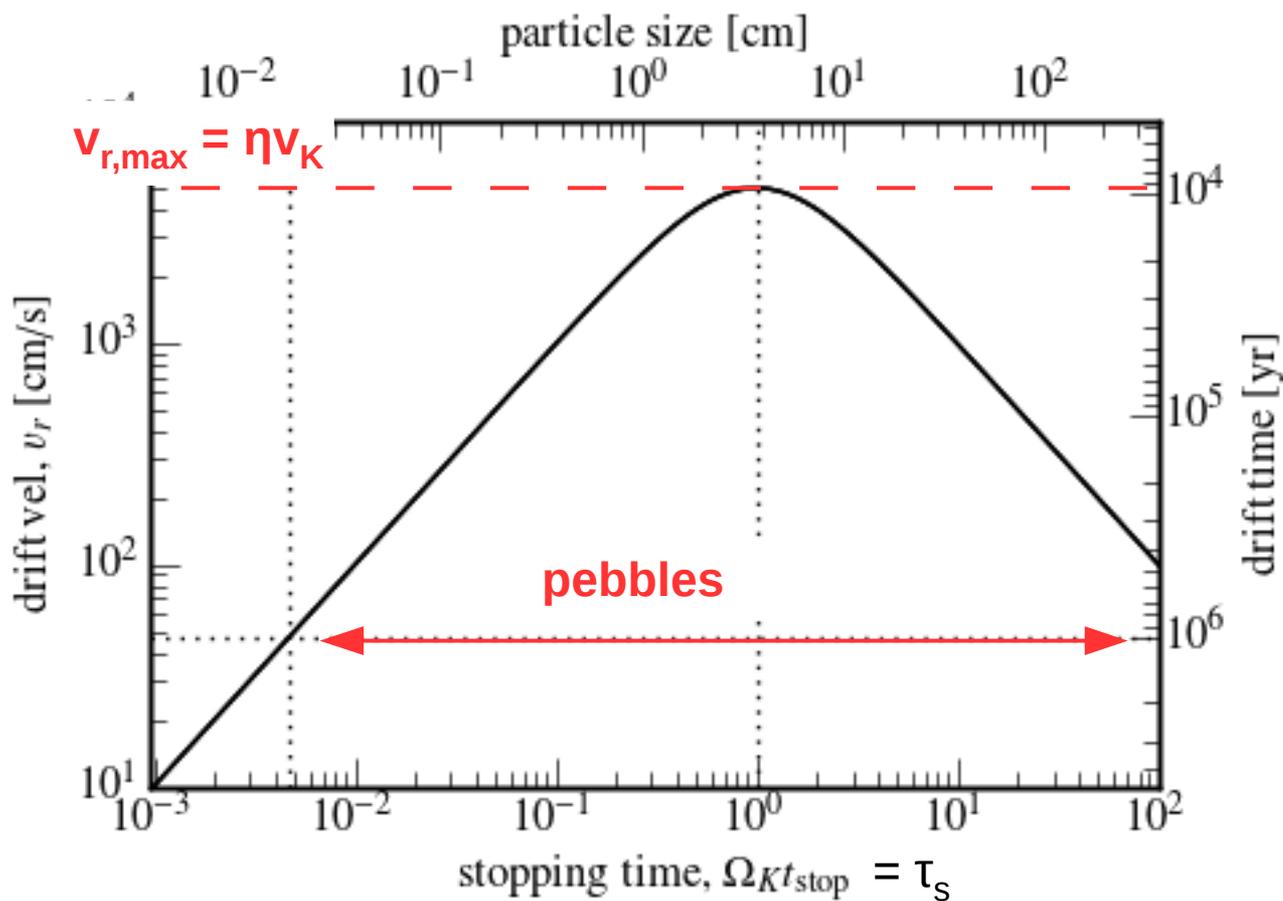
Pebble definition

Geologist: *particles* $2 \text{ mm} < d < 6.4 \text{ cm}$

Astrophysical: *particles that drift*

6.4 cm

at 100 au for power-law disk



1 dm

1 cm

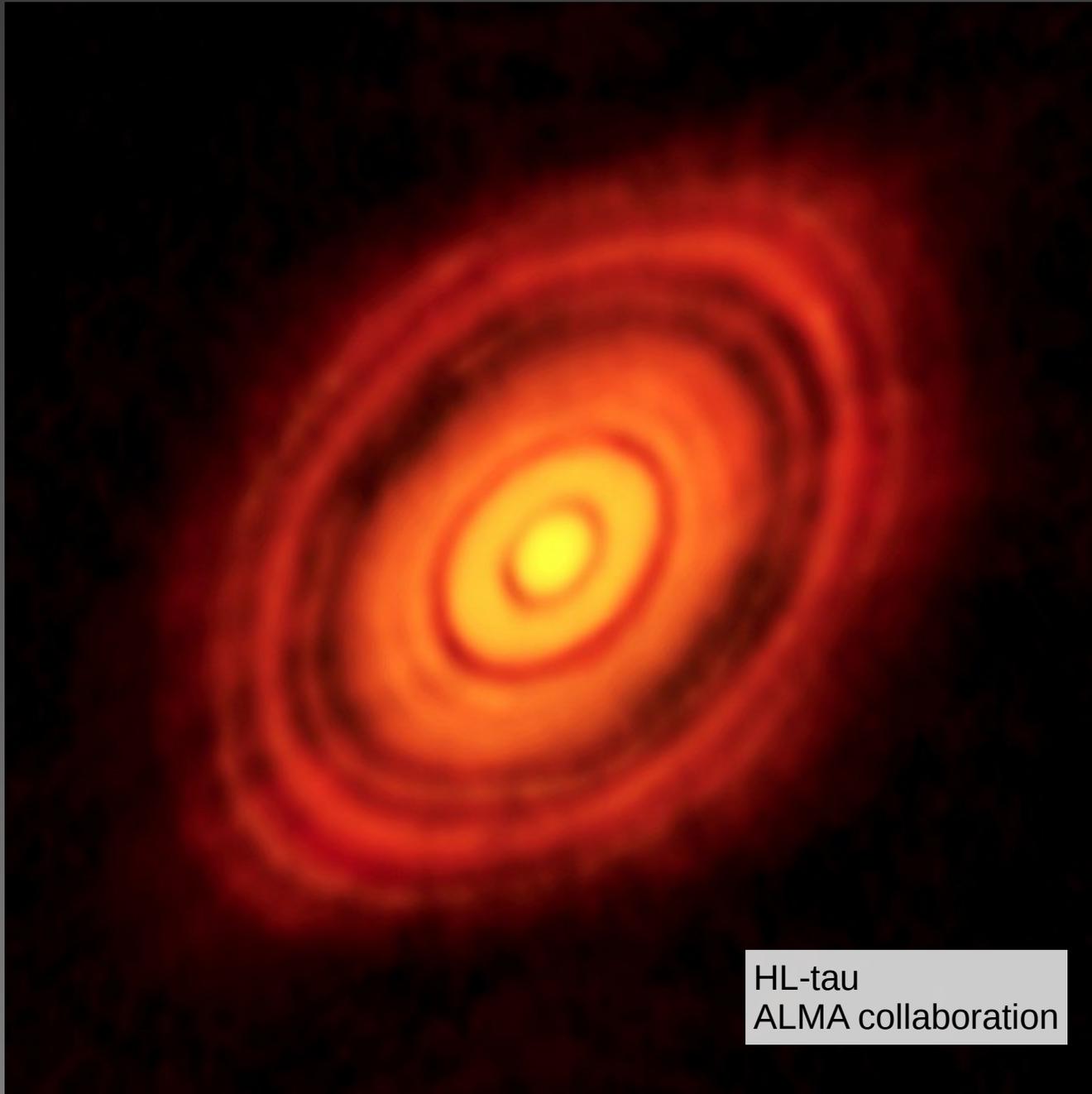
1 mm

100 μm

10 μm

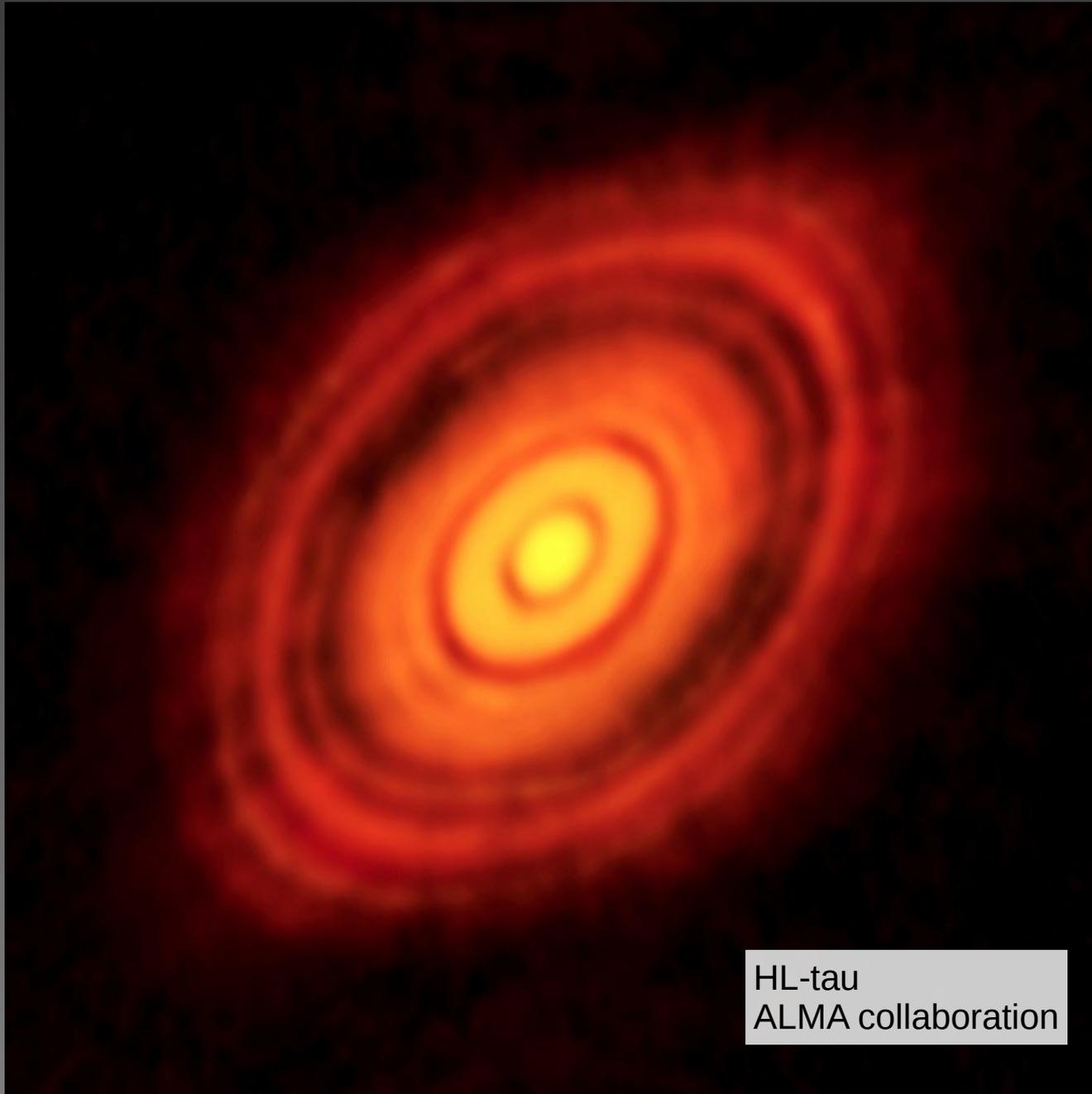
1 μm

A pebble disk



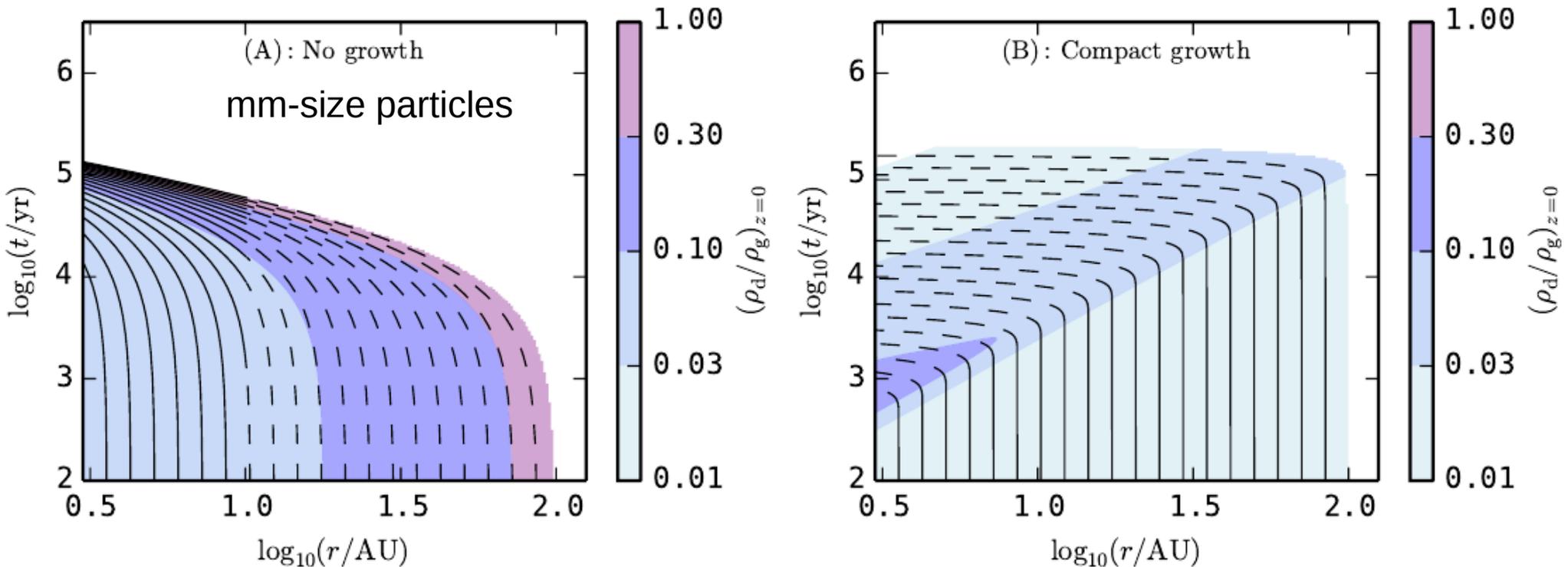
HL-tau
ALMA collaboration

A pebble disk



Disks evolve inside out

Krijt et al. (2016)



No growth: pileups

radial velocity decreases inward

cf. Youdin & Shu (2002); Youdin & Chiang (2004)

Growth + drift

No strong pileups!

cf. Birnstiel et al. (2012); Okuzumi et al. (2012); Lambrechts & Johansen 2014; Sato et al. (2016)

Need special conditions

Planetesimal formation near iceline

Planetesimal formation at iceline

By streaming instability once $\rho_{\text{solid}}/\rho_{\text{gas}} > \sim 1$

traffic jam effect

Ida & Guillot (2016)

Diffusion & recondensation of H₂O

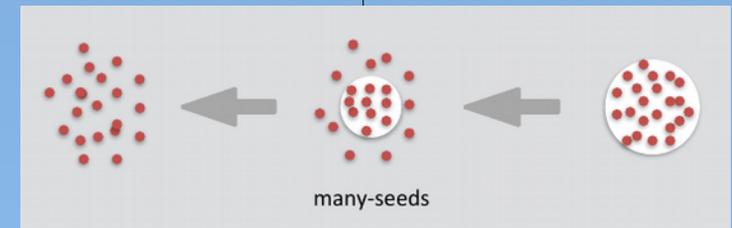
- Stevenson & Lunine (1988)
- Ros & Johansen (2013)
- Schoonenberg & Ormel (2017)

To summarize:

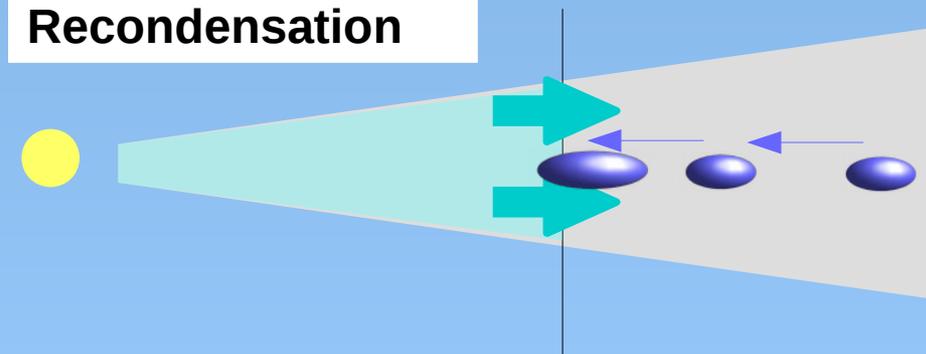
“Planetesimals form *first* at H₂O iceline”

Drazkowska & Alibert (2017)

Traffic jam

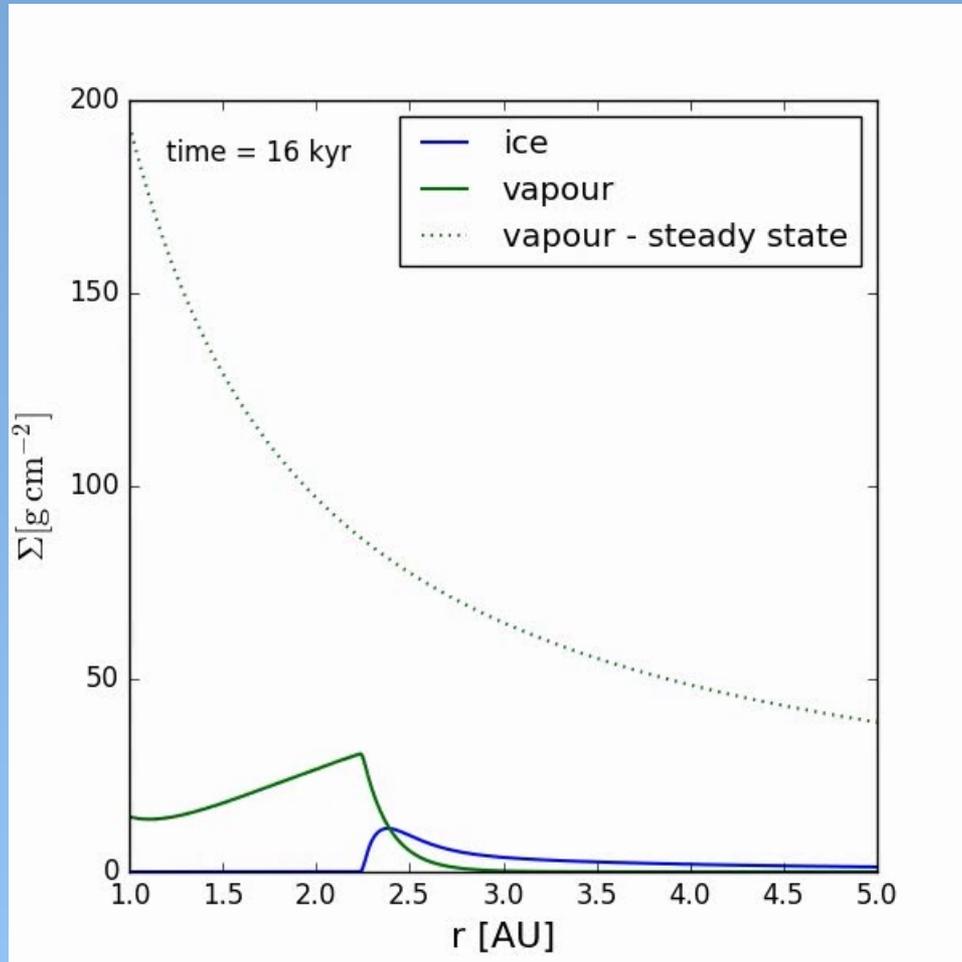


Recondensation



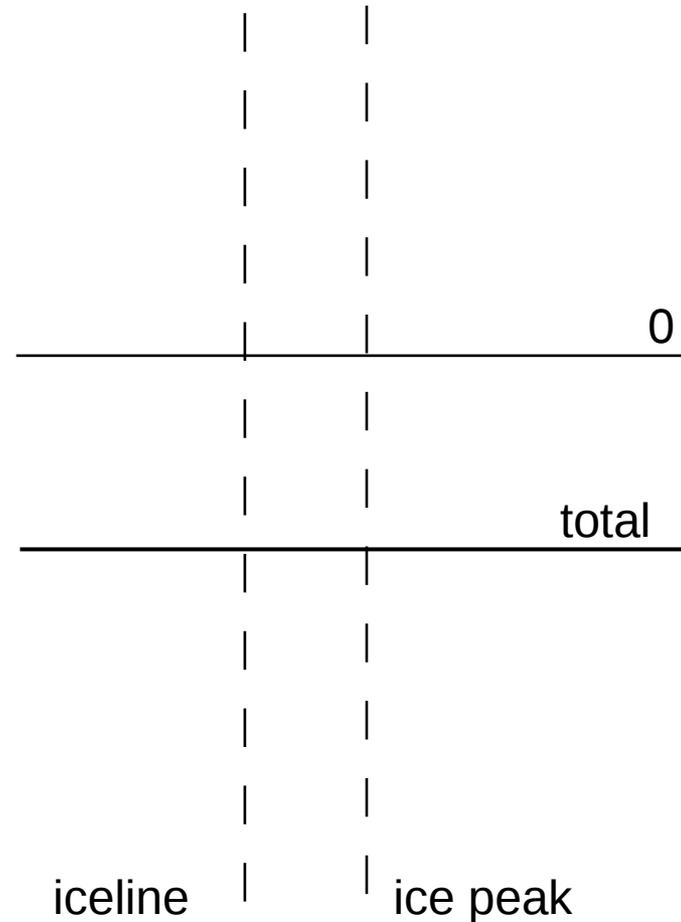
H₂O Iceline

Schoonenberg & Ormel (2017)



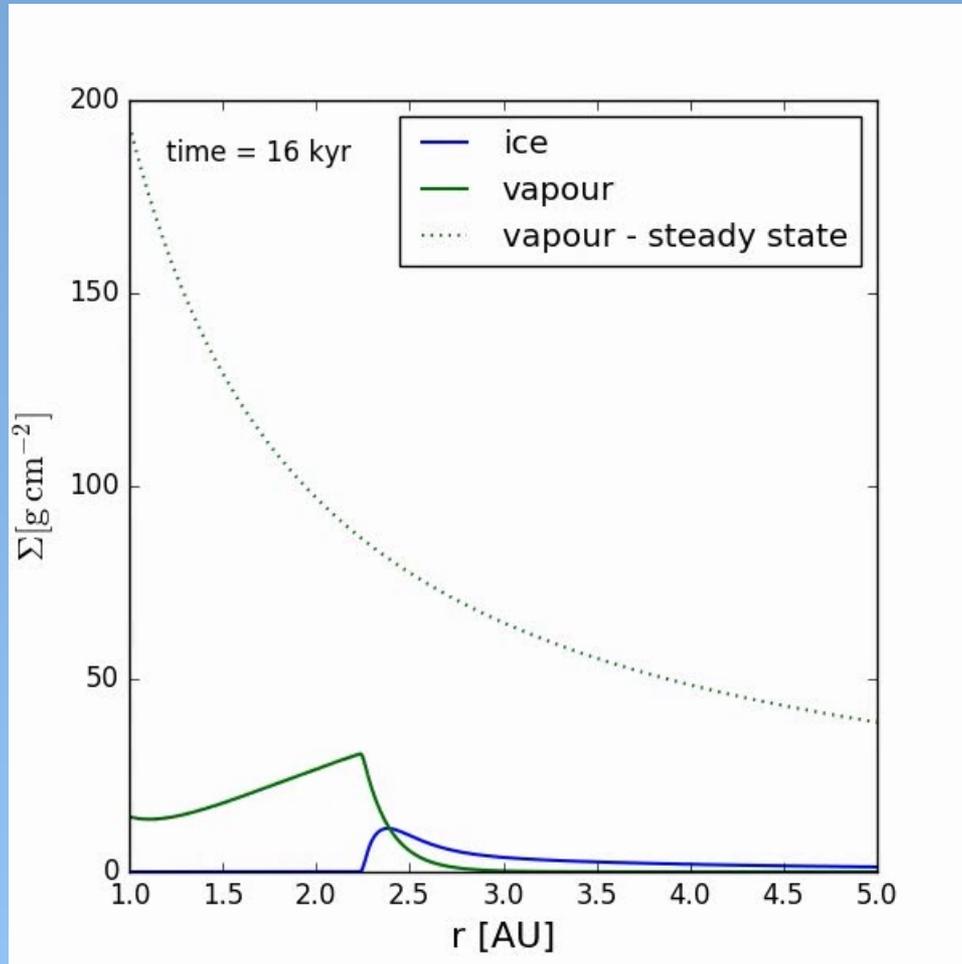
Schoonenberg & Ormel (2017)

$$\text{Mass flux} = \Sigma v + \Sigma_{\text{gas}} D_i \nabla \left(\frac{\Sigma_i}{\Sigma_{\text{gas}}} \right)$$

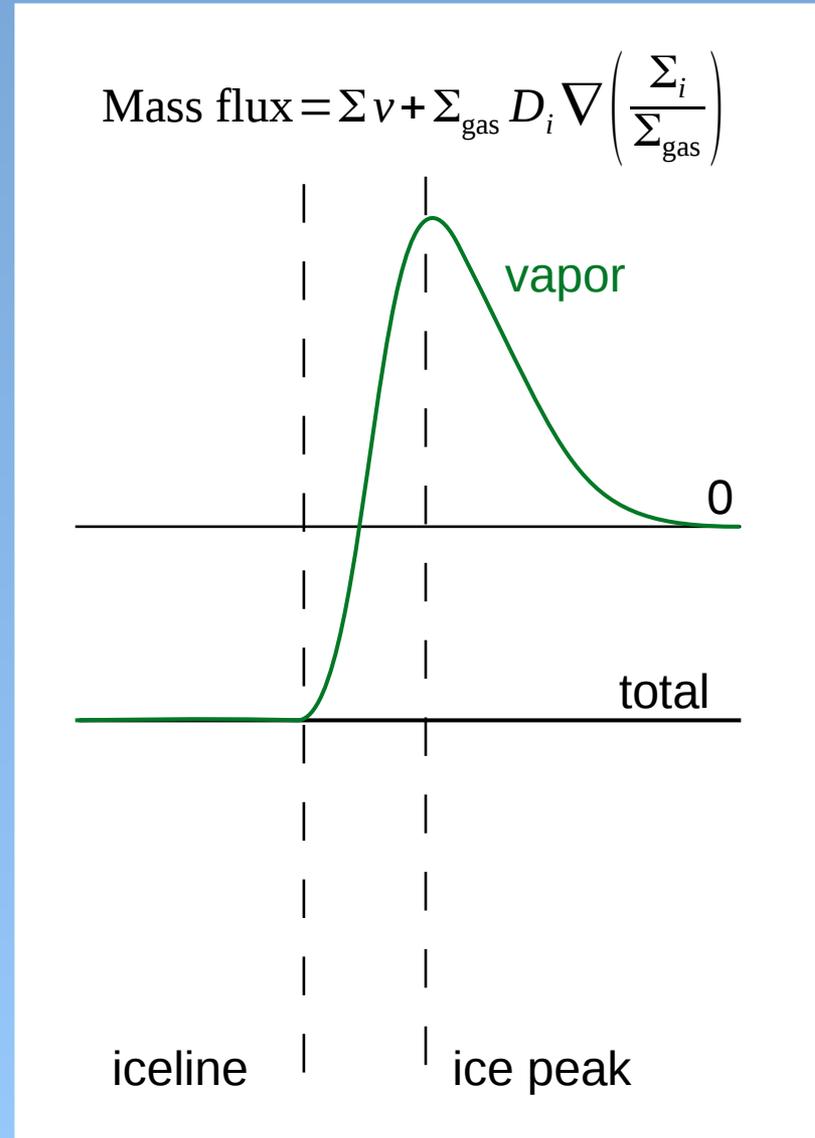


H₂O Iceline

Schoonenberg & Ormel (2017)

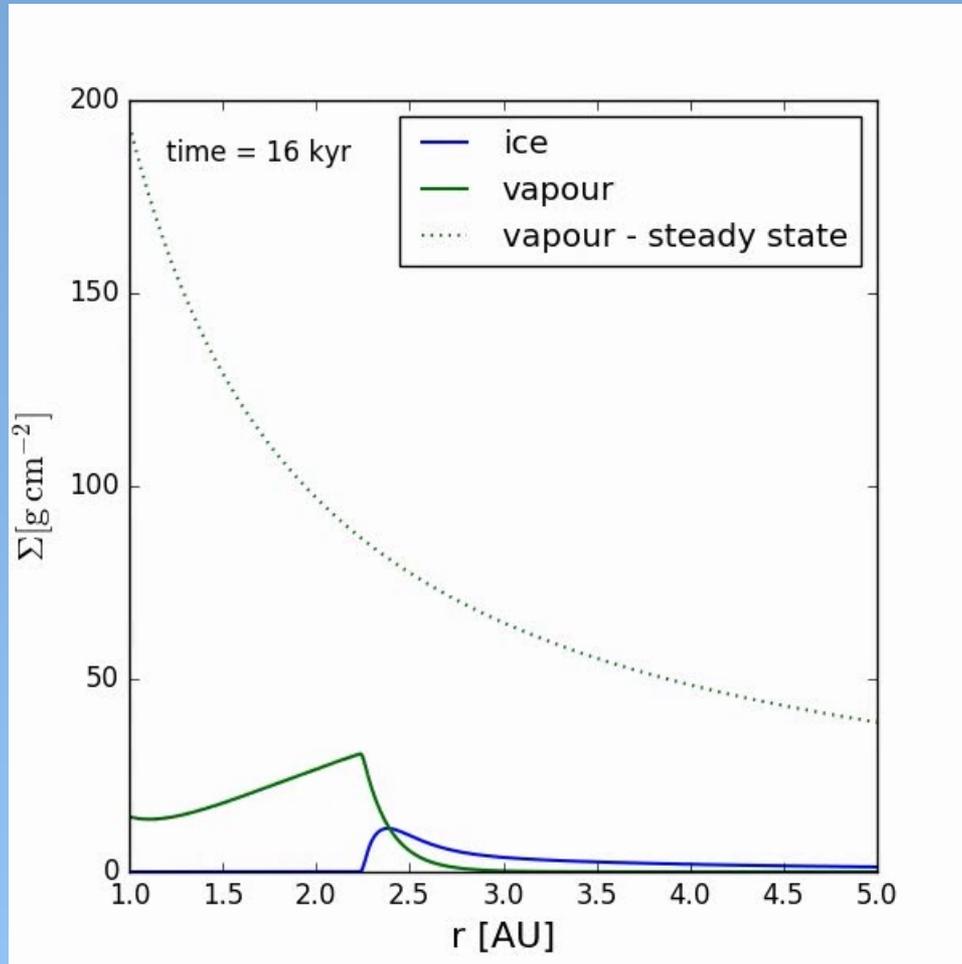


Schoonenberg & Ormel (2017)

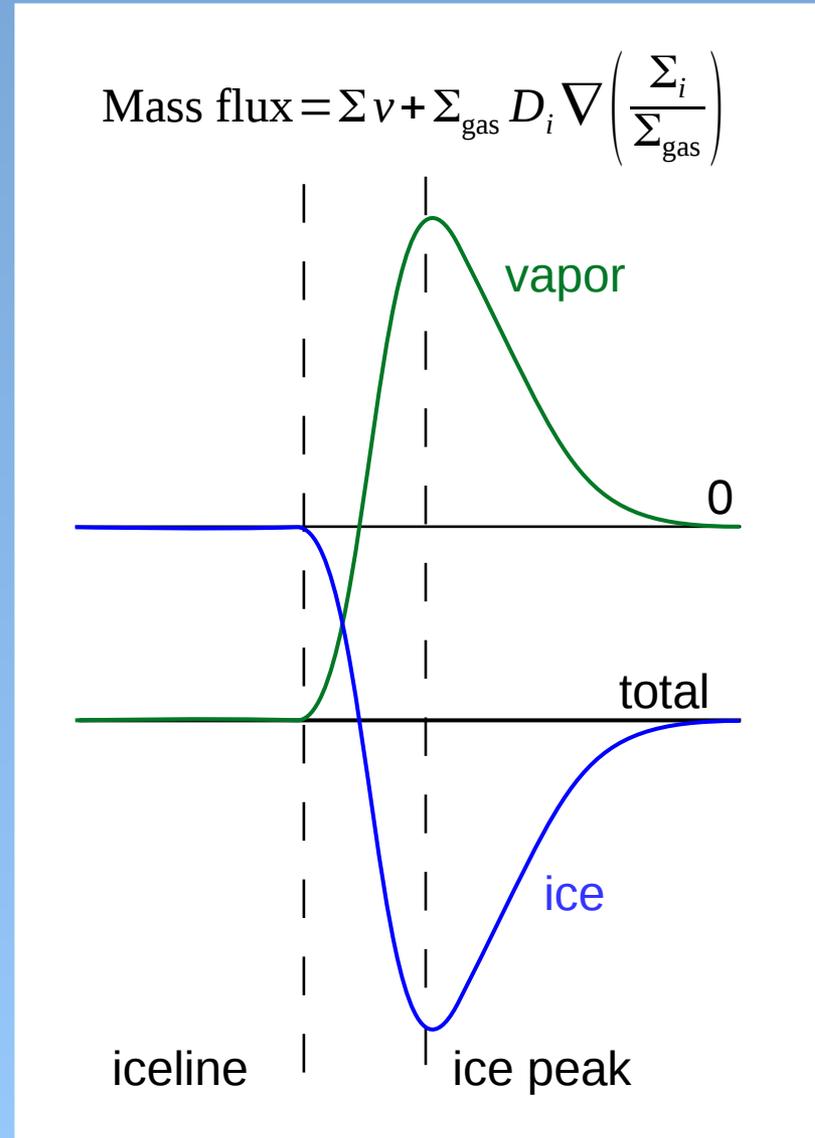


H₂O Iceline

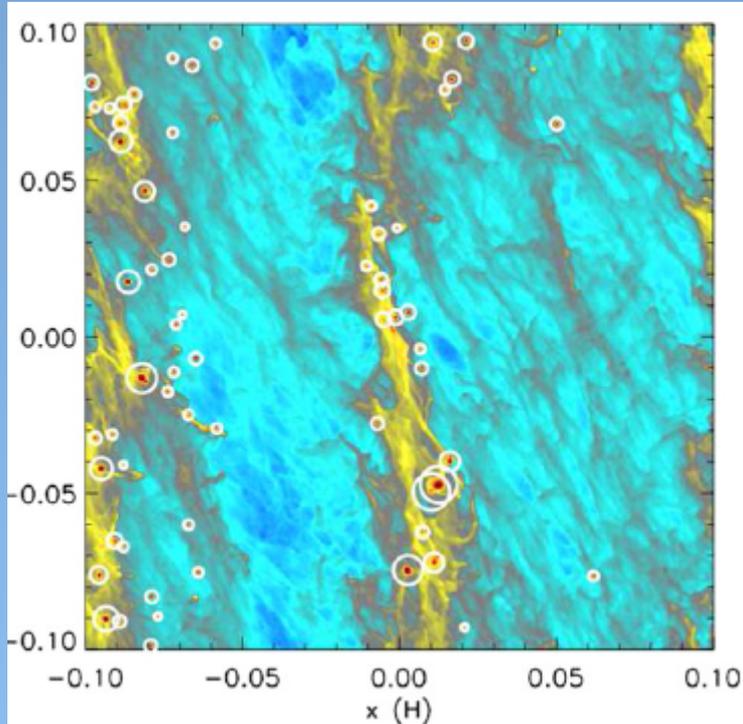
Schoonenberg & Ormel (2017)



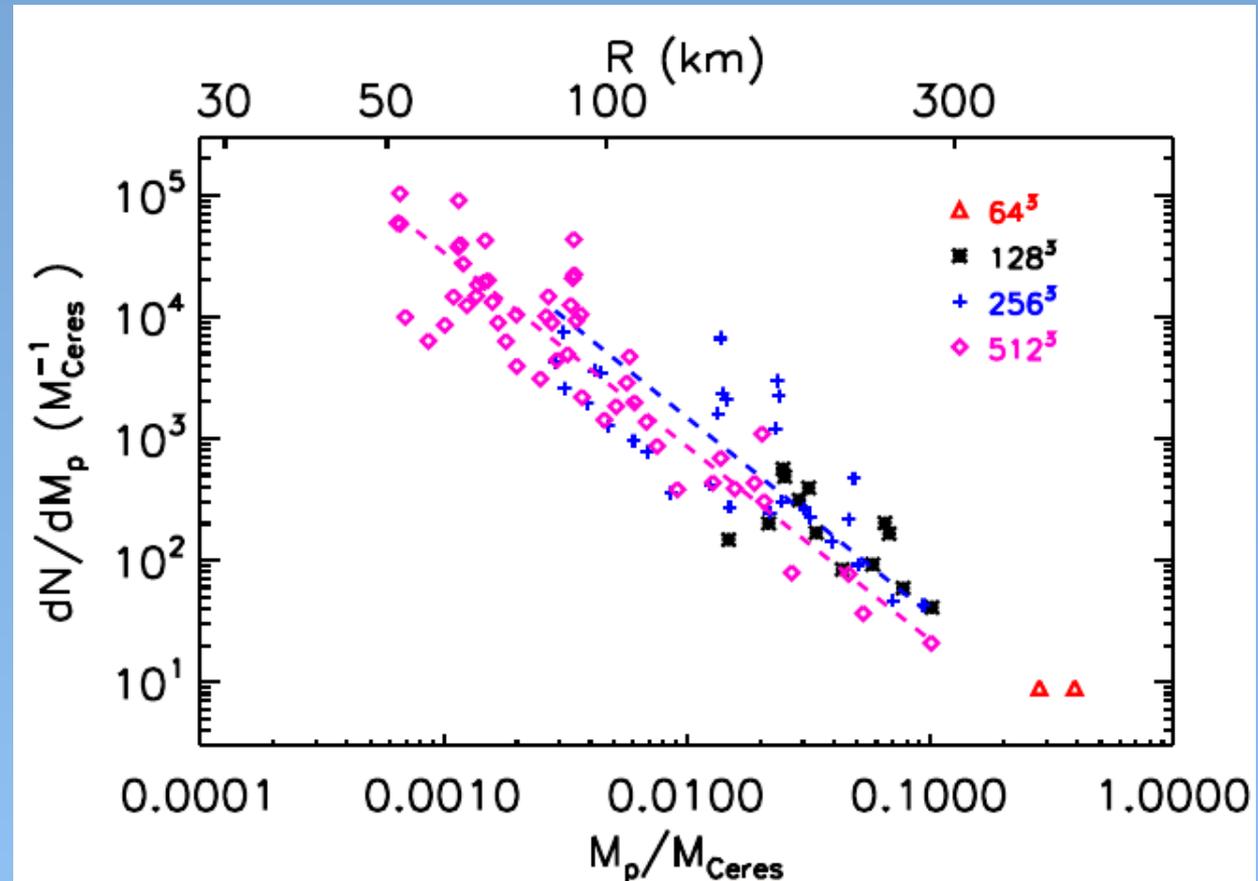
Schoonenberg & Ormel (2017)



Planetesimal IMF

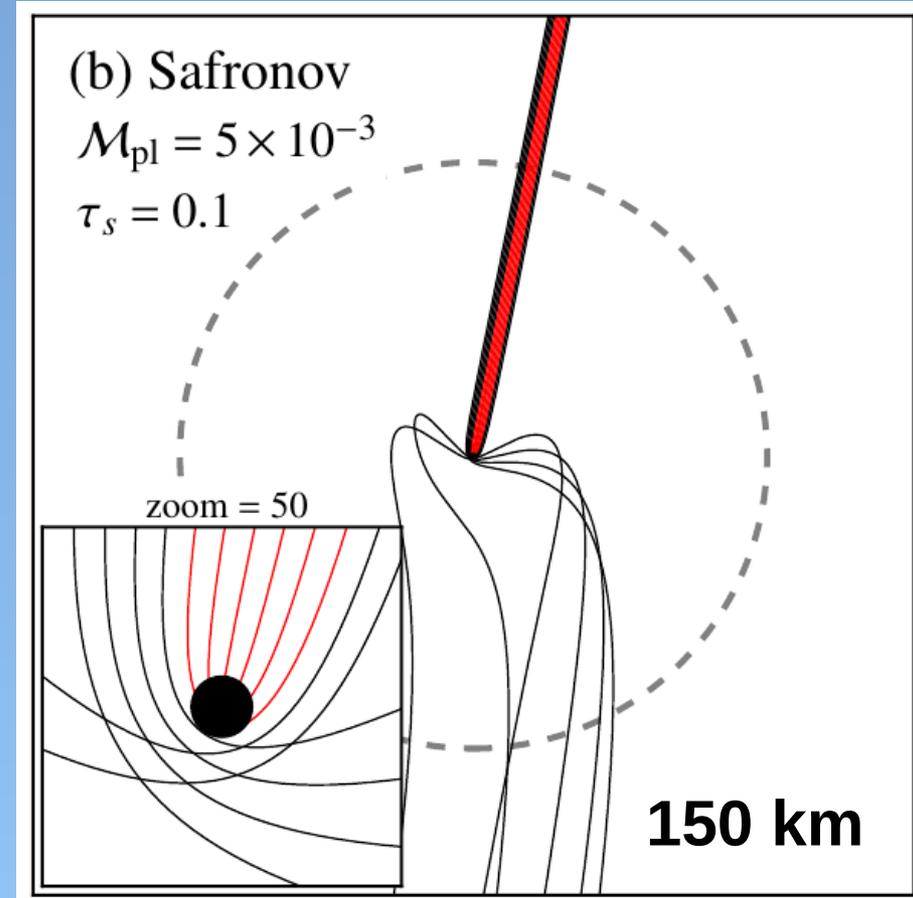
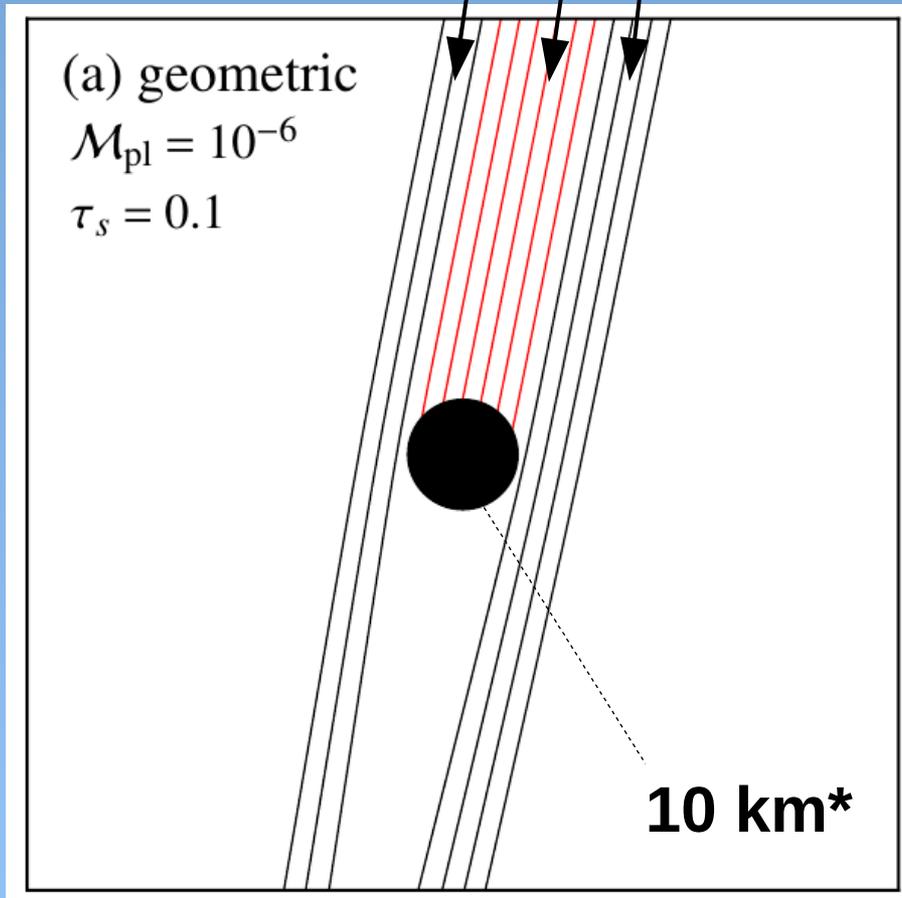


Johansen et al. (2015)
Simon et al. (2016)
Schaefer et al. (2017)



Is ~ 100 km sufficient for pebble accretion?

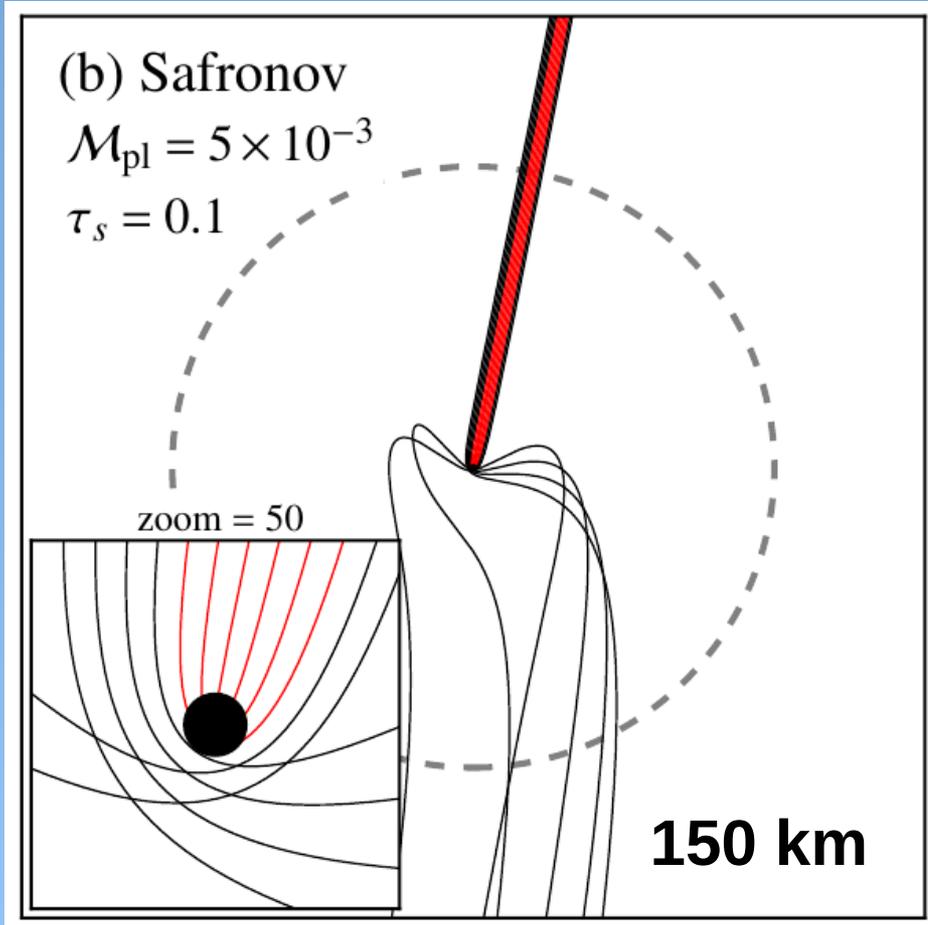
No pebble accretion



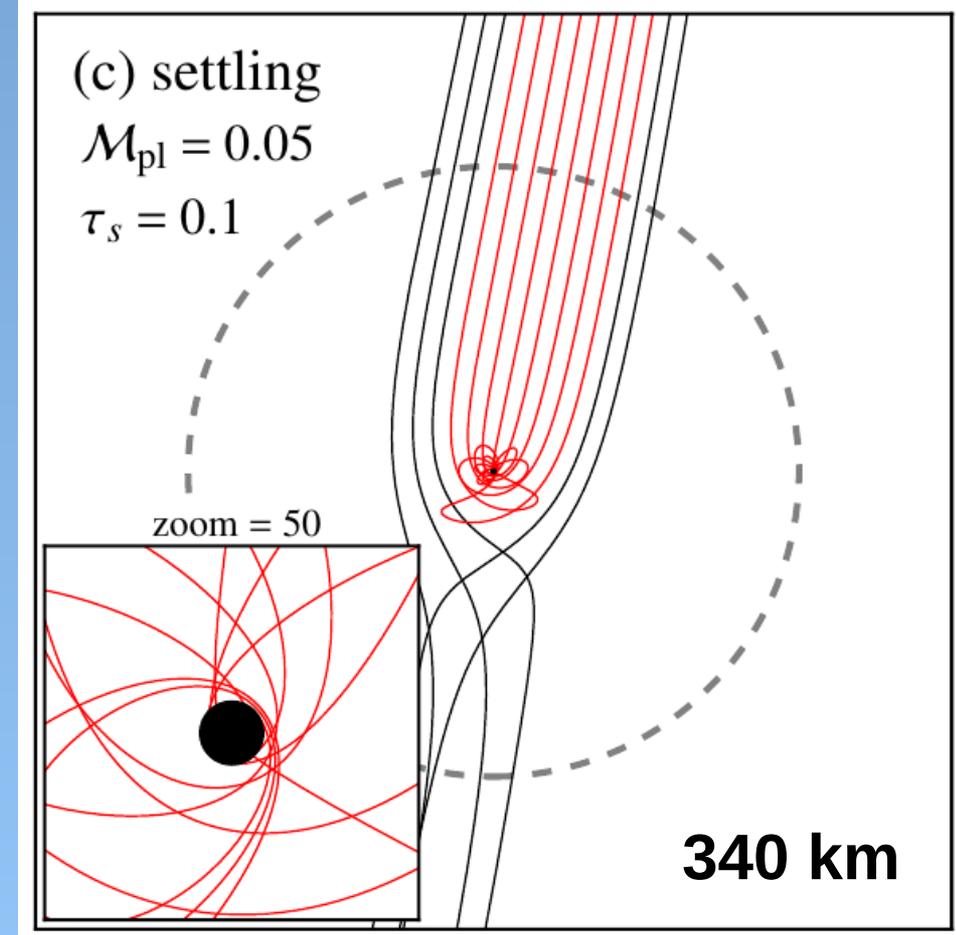
* 1 au, solar-mass star, disk headwind 50 m sec⁻¹

Ormel (2017)

Onset pebble accretion

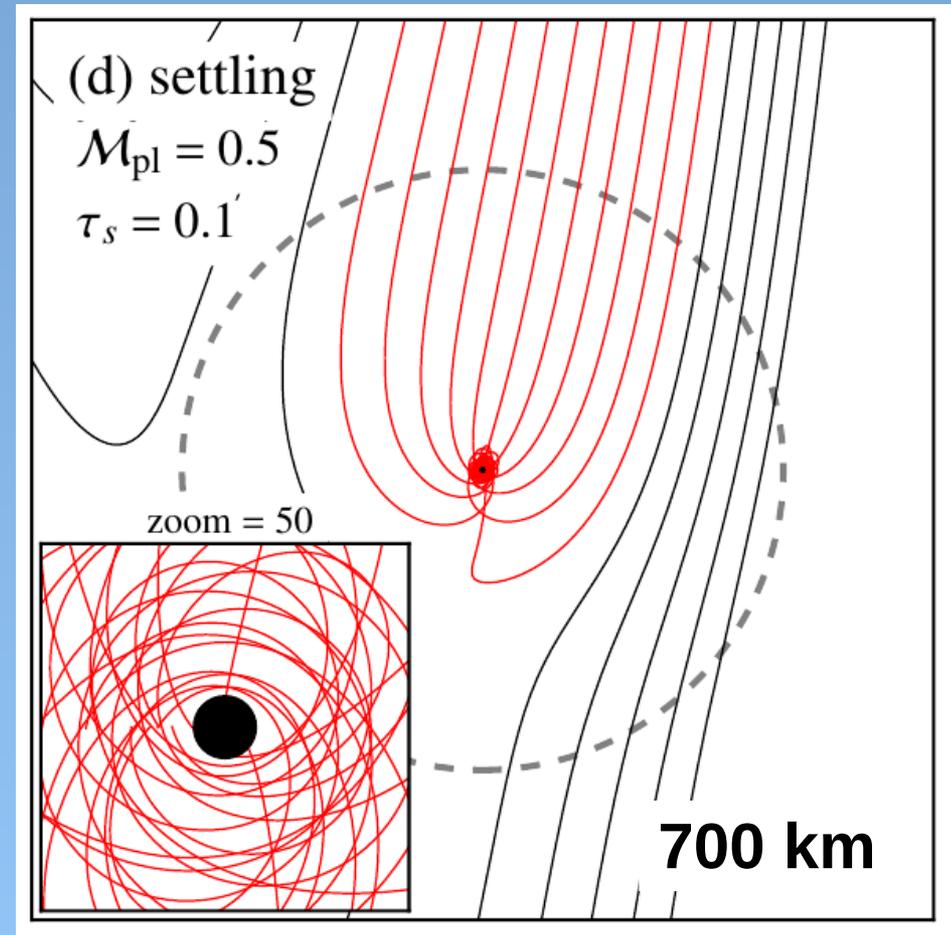
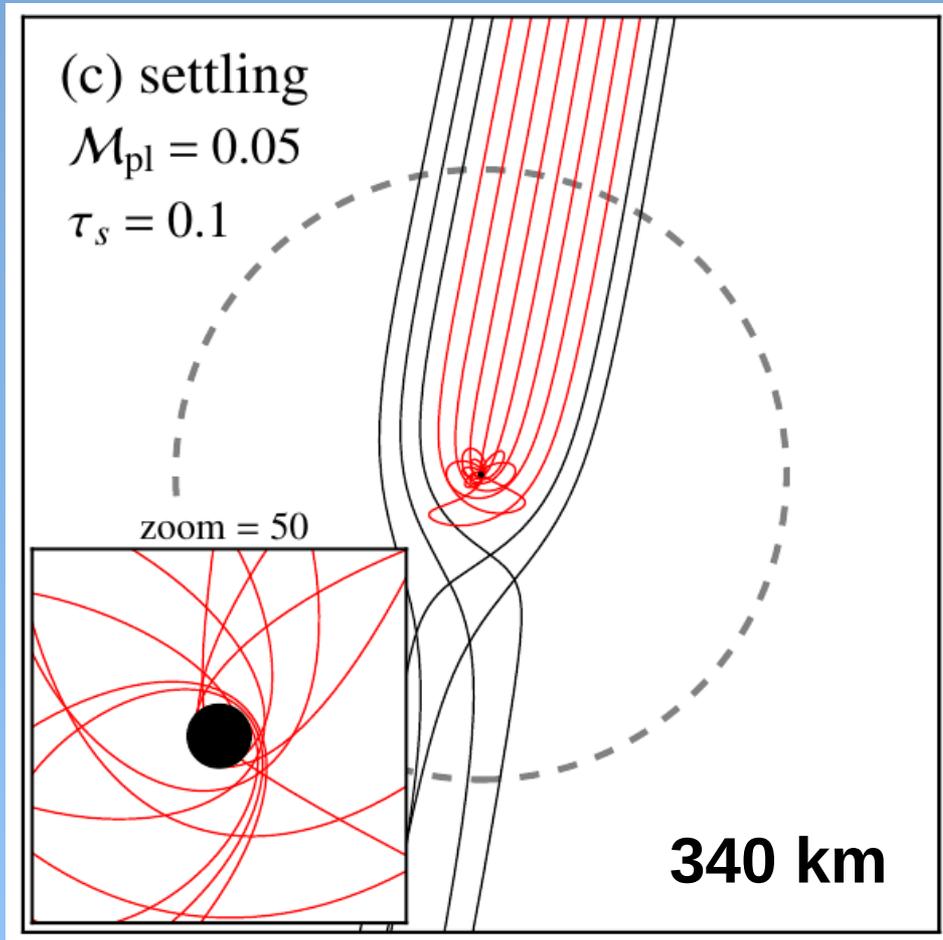


No Pebble accretion



Pebble accretion

Onset pebble accretion



Ormel (2017)

Onset of pebble accretion

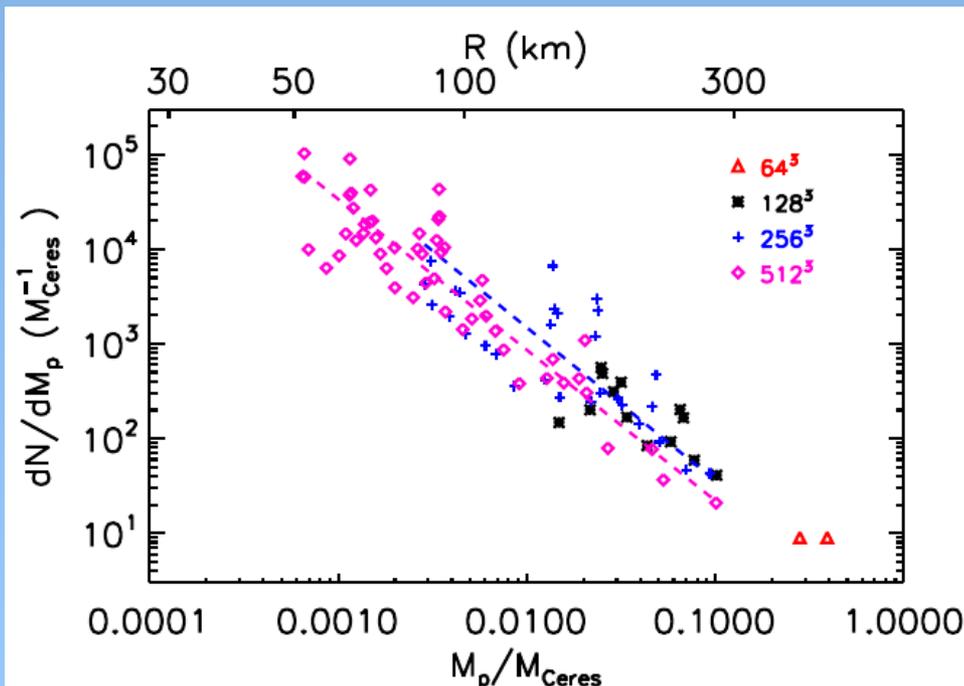
$$R_{\text{P.A.}} \approx 220 \text{ km} \left(\frac{v_{\text{hw}}}{50 \text{ m/sec}} \right) \left(\frac{r}{\text{au}} \right)^{0.42} \left(\frac{\tau_s}{0.1} \right)^{0.28}$$

Visser & Ormel (2016)

Smaller in colder disks or in pressure bumps

Increases with disk orbital radius

Larger for larger pebbles



Planetesimal → pebble-accretion bodies
 May need “classical” phase
 e.g., planetesimal runaway growth
 Levison et al. (2015)

Not all pebbles are accreted!

Efficiency = probability of accretion
In 2D limit

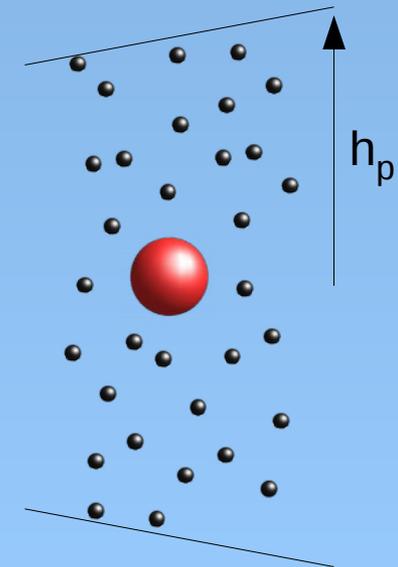
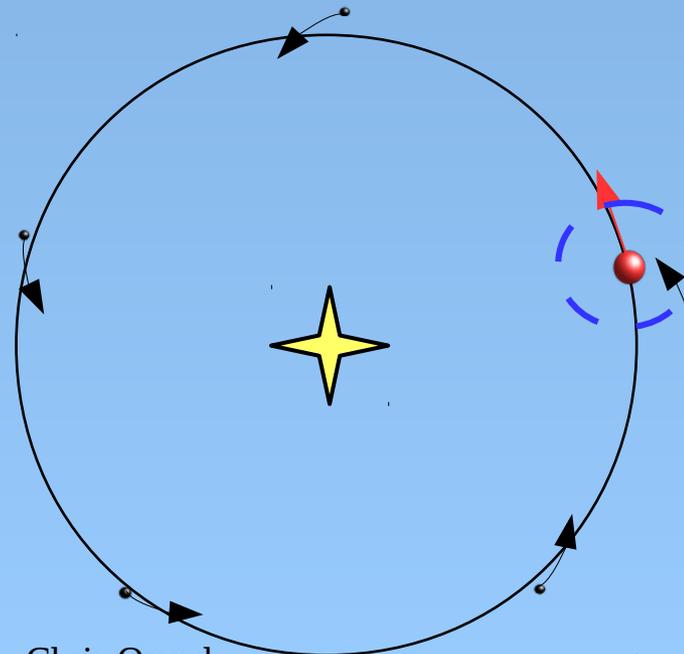
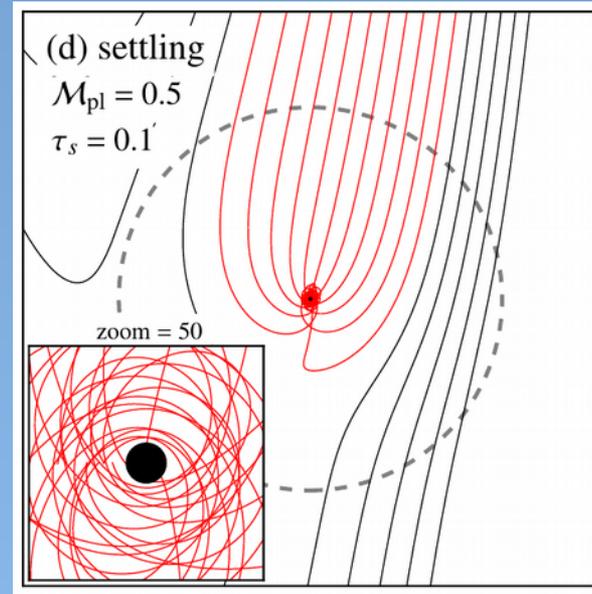
$$\epsilon_{2D} = \frac{0.31}{\eta} \left(\frac{M_p}{M_\star} \frac{\Delta v}{\tau_s v_K} \right)^{1/2}$$

In 3D limit

$$\epsilon_{P.A.} = 0.39 \frac{M_p}{M_\star} \frac{1}{\eta H_p / r}$$

Liu & Ormel, *subm.*

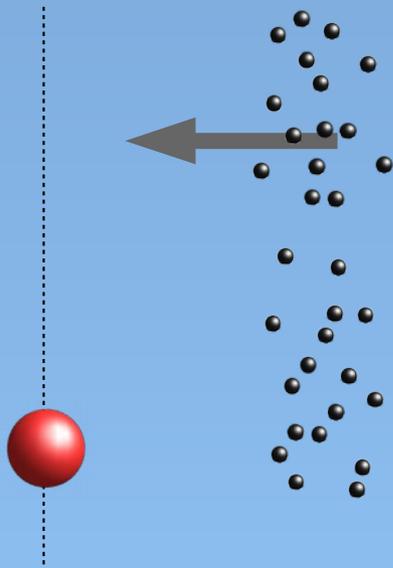
Expression consistent w/ literature
but prefactors different



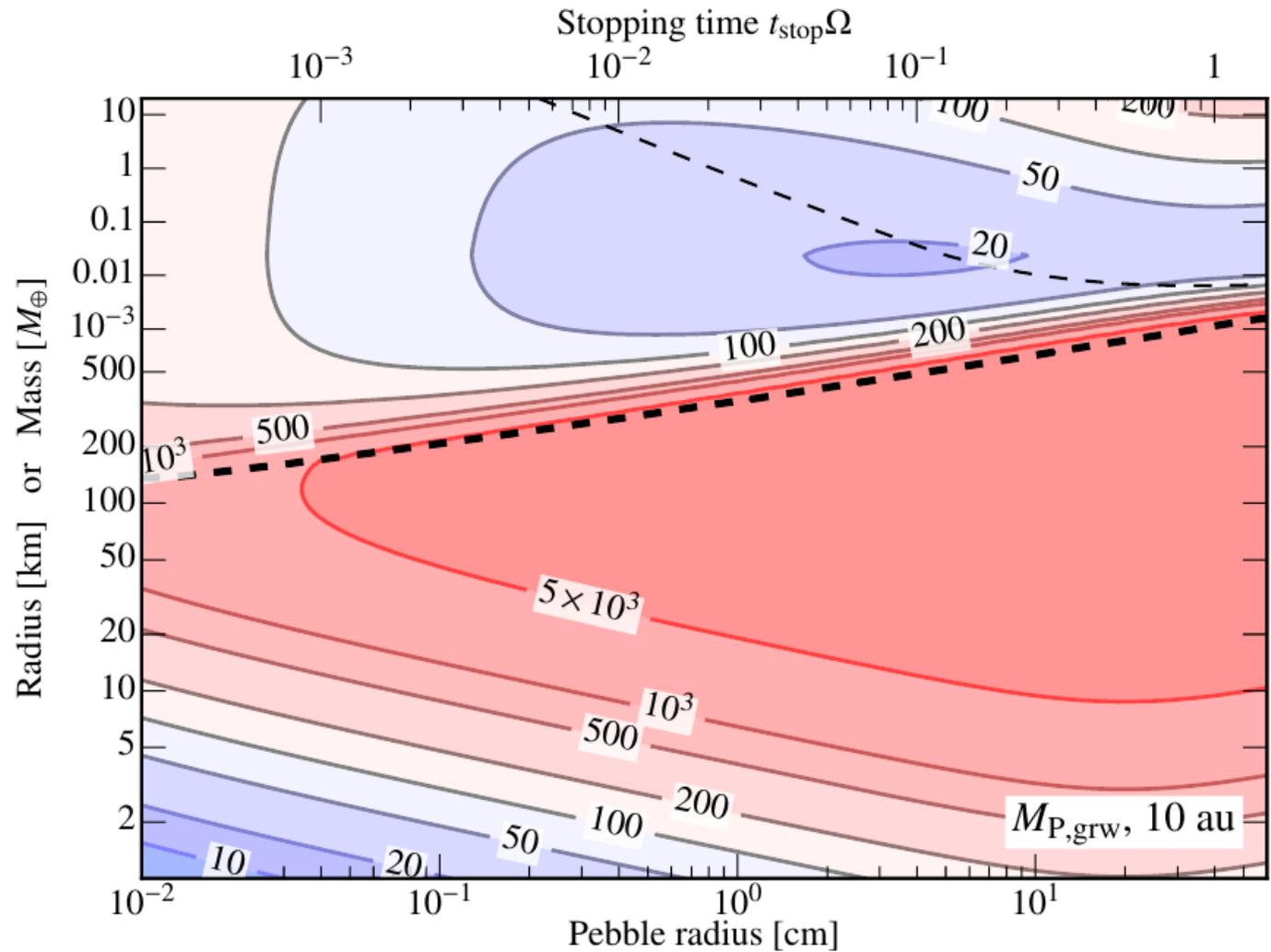
Pebble-doubling mass

(in Earth masses)

$$m_{P,grw} = M_p / \epsilon_{P.A.}$$



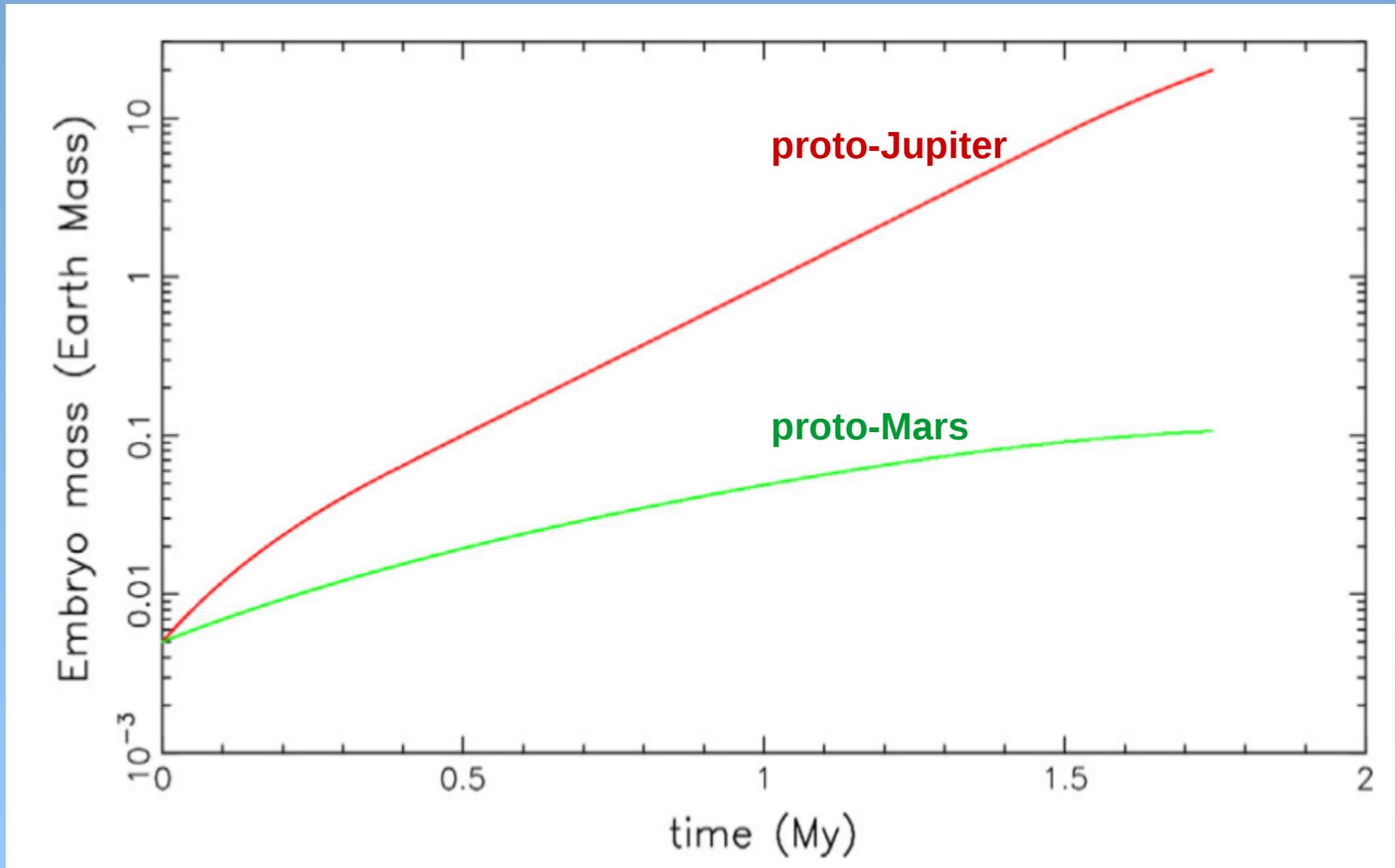
How many pebbles
[Earth masses]
needed to double the
mass of the planet →



Ormel (2017), download [here](#).

Application I: solar system

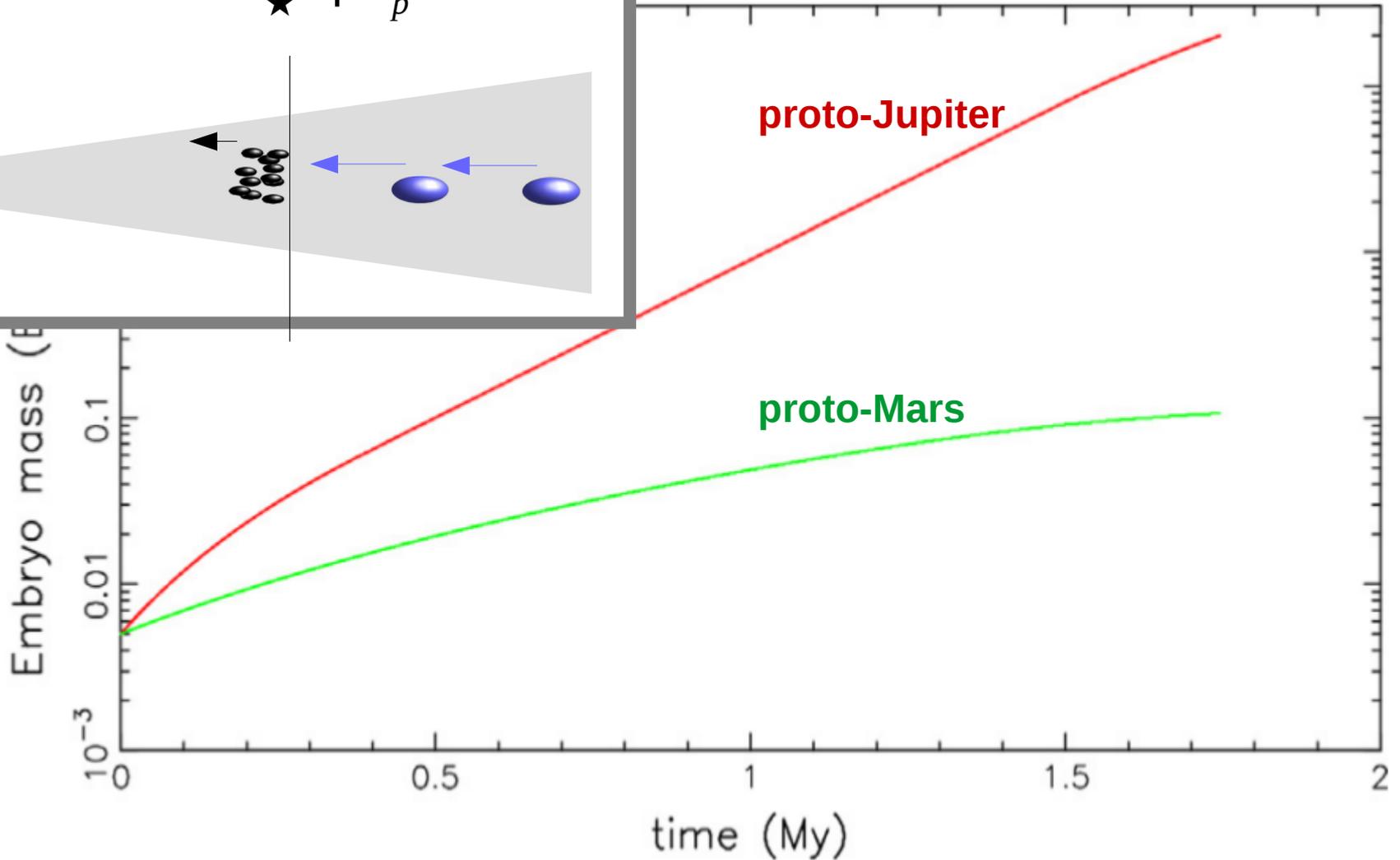
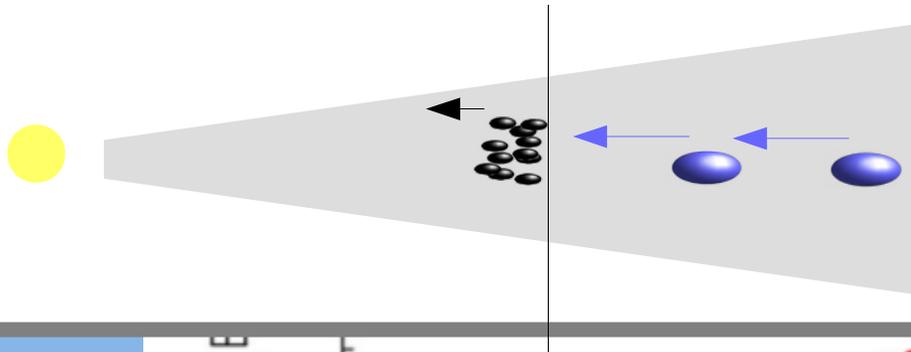
Morbidelli et al. 2015



Application I: solar system

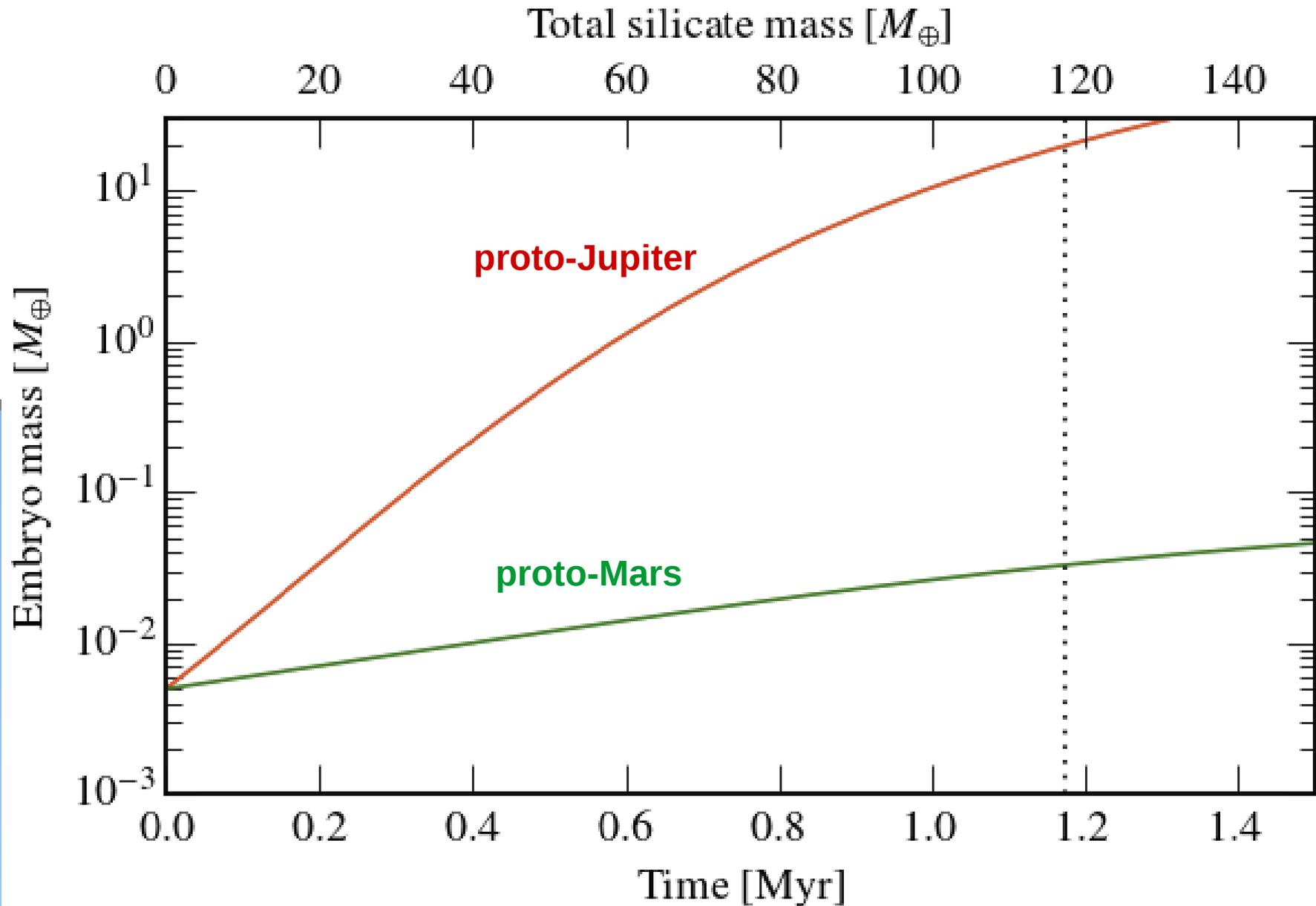
Morbidelli et al. 2015

$$\epsilon_{\text{P.A.}} = 0.39 \frac{M_p}{M_\star} \frac{1}{\eta H_p / r}$$



Application I: solar system

Morbidelli et al. 2015



Application II:

Icelines around low-mass stars

Pebble accretion is *far more efficient* around low-mass stars b/c:

- further in
- Hill radii larger

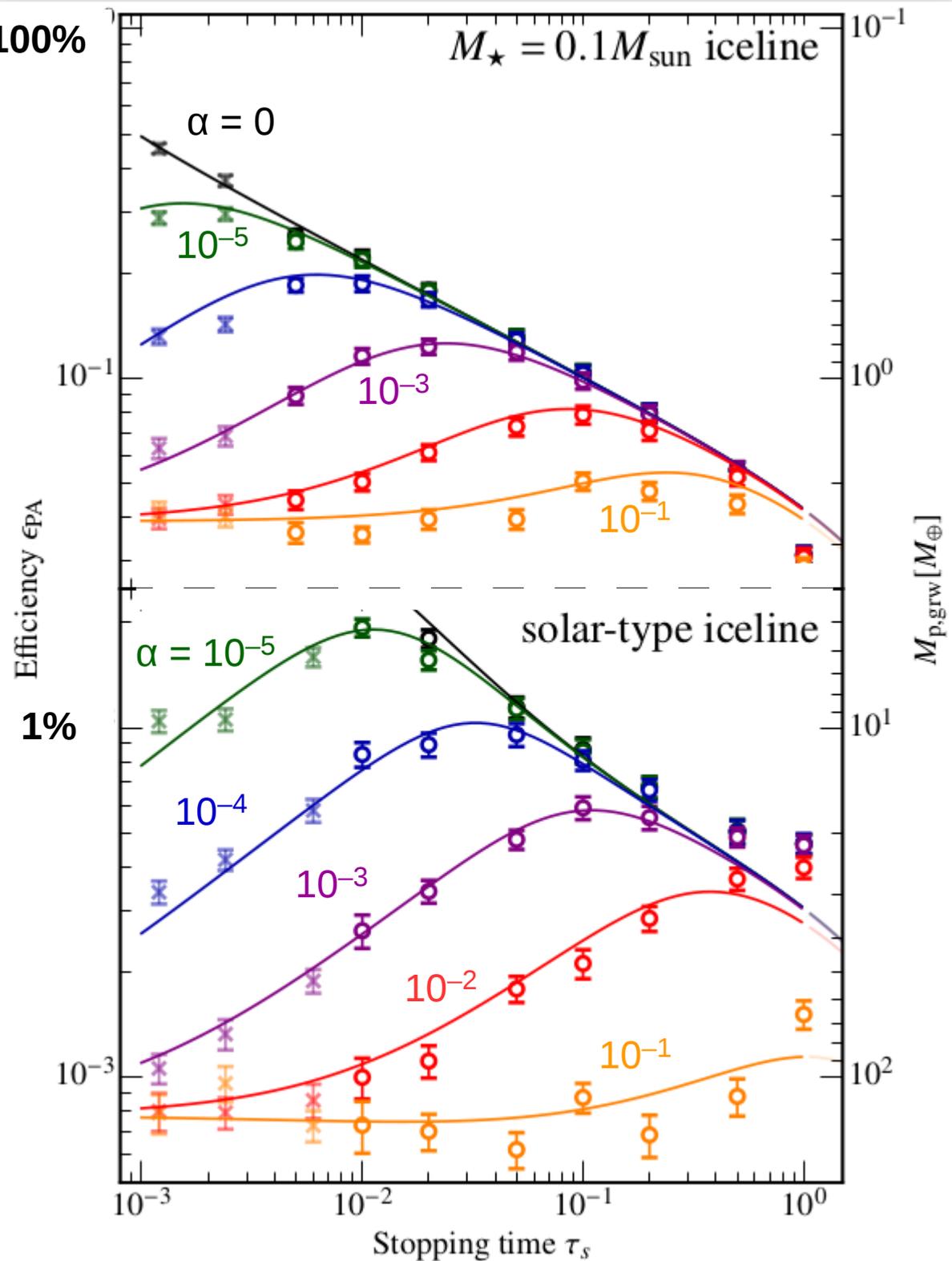
$$\epsilon_{\text{P.A.}} = \frac{0.39}{\eta} \frac{M_p}{M_\star} \frac{1}{H_p/r}$$

Efficiency and pebble doubling mass for a $0.1 M_{\text{Earth}}$ mass \rightarrow
 [Ormel & Liu, in prep.]

* solar-type star: $h=0.05$; $\eta=0.003$

* M-star: $h=0.03$; $\eta = 0.001$

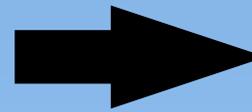
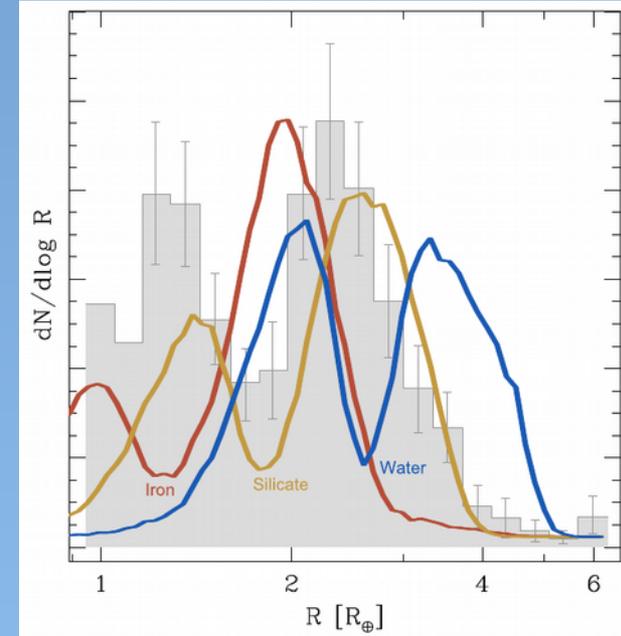
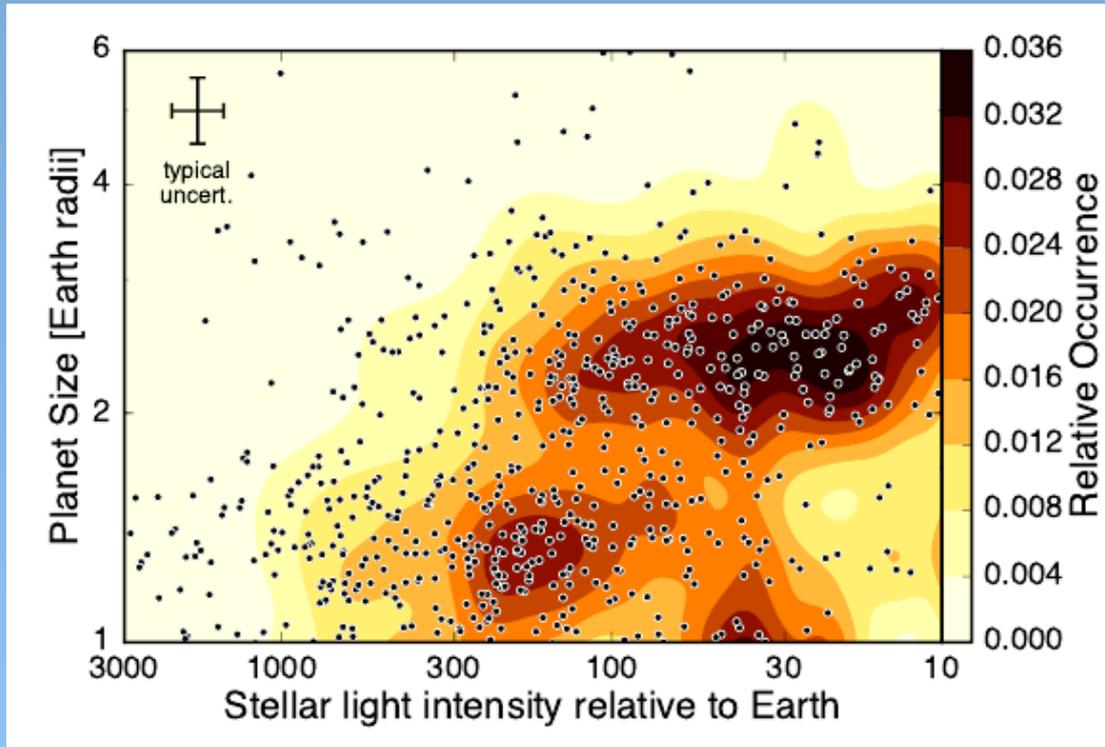
100%



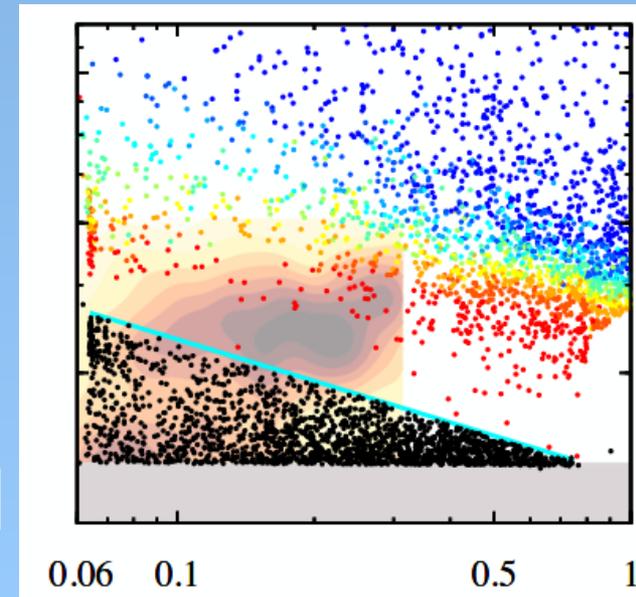
Composition of planets

How to keep them rocky?

Owen & Wu (2017)



Fulton et al. (2017)



Jin & Mordasini

1. Formation, then migration

Rock after ice

- start icy at snowline
- migrate interior to snowline
- accrete silicate pebbles afterwards
- end up w/ rocky composition

Scenario works esp. for low-mass stars

Make similar planets

similar size planets in system
indicative of formation at single location?

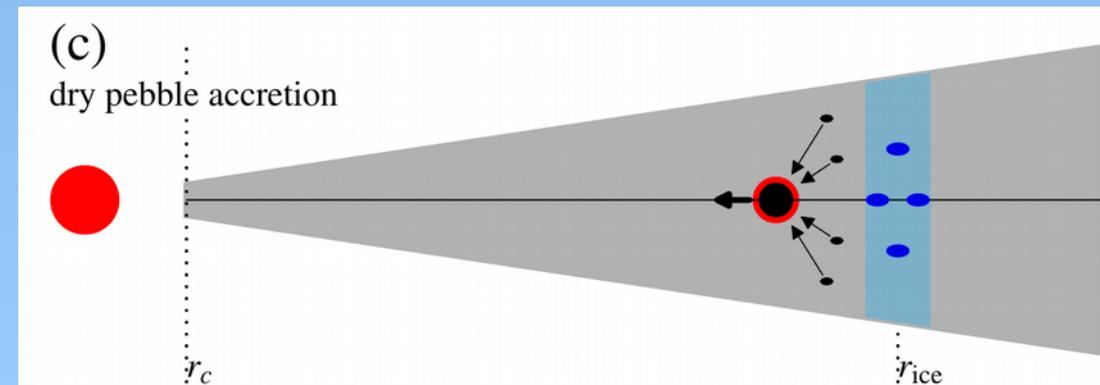
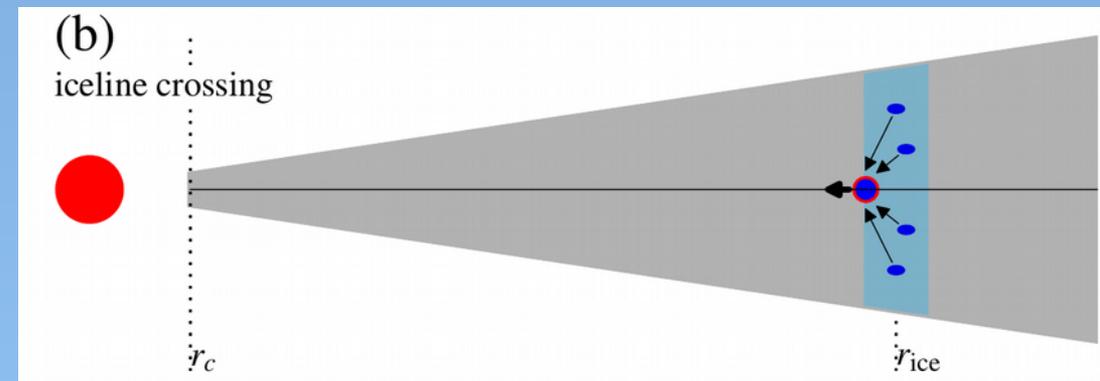
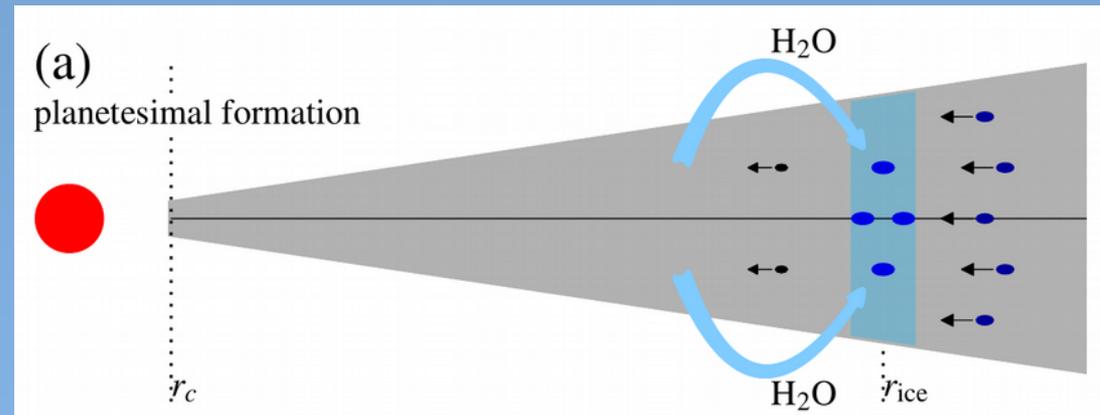
Weiss et al. (2017)

cf. “inside-out” planet formation models

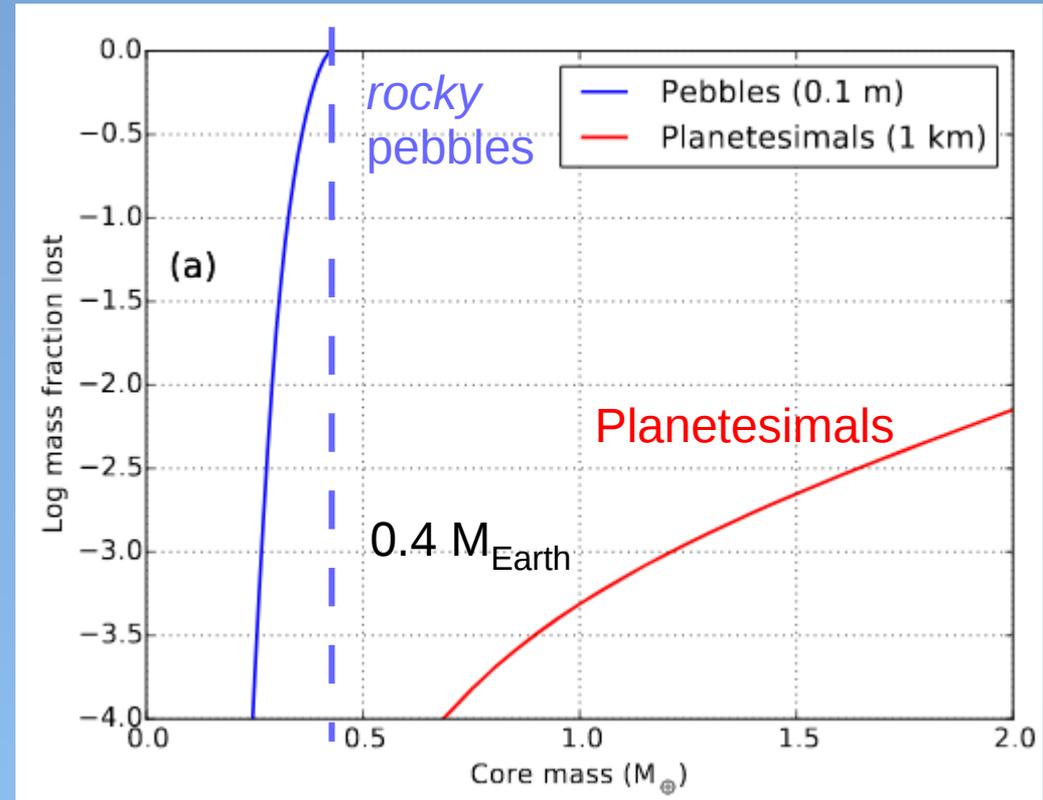
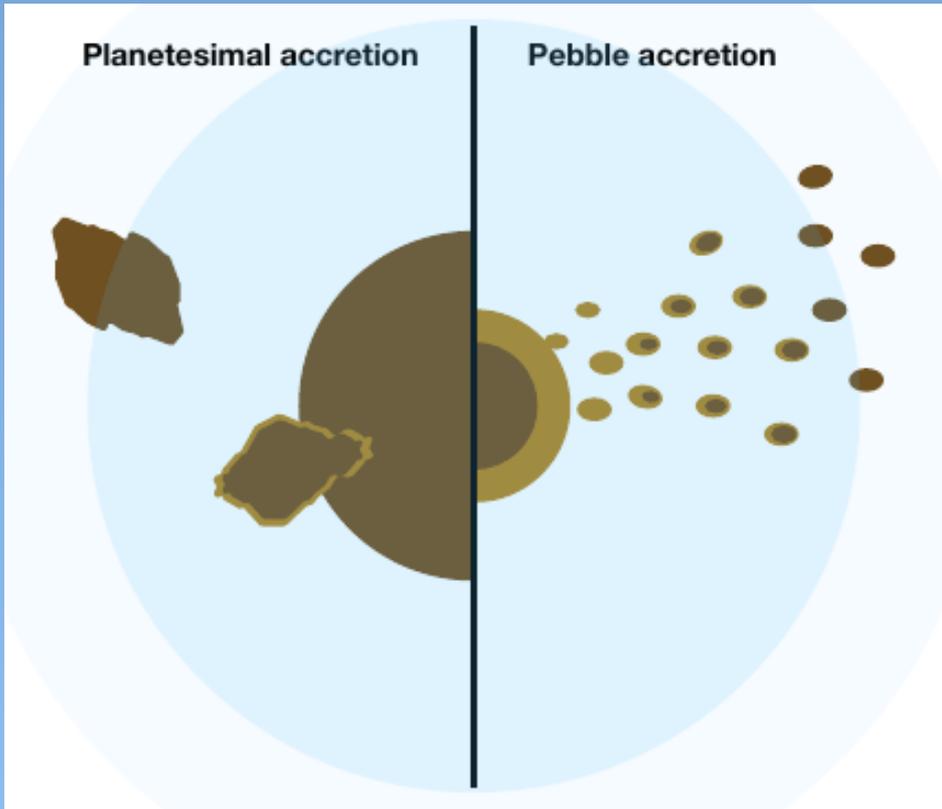
Chatterjee & Tan (2014, 2015); Hu et al. (2017)

Scenario for formation of TRAPPIST-1 system →

Ormel et al. (2017)



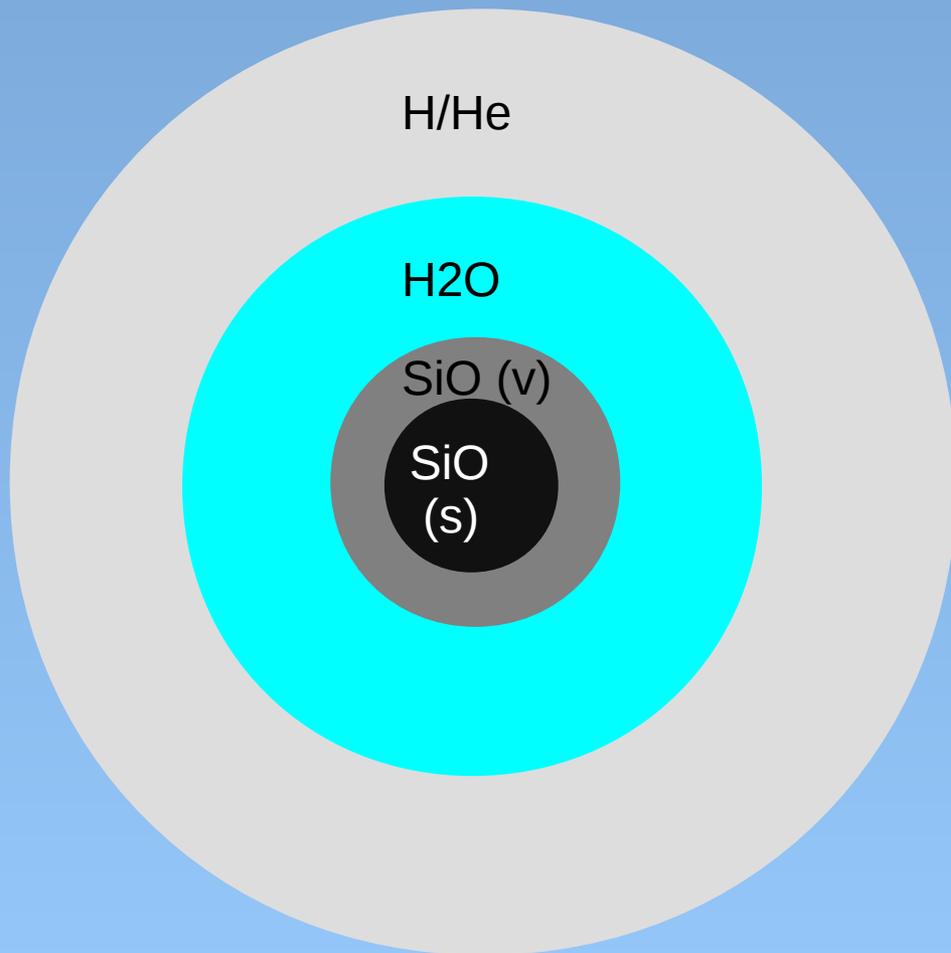
2. Pebble evaporation in envelope



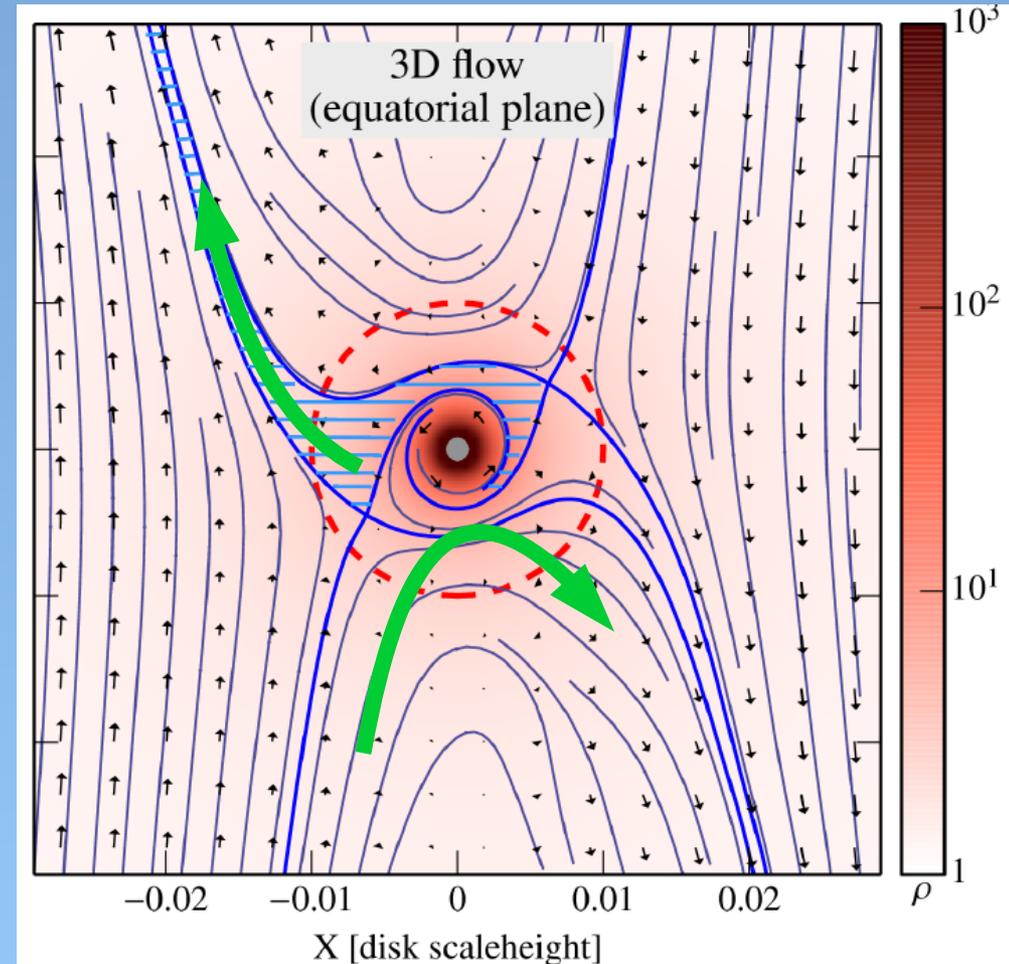
Brouwers, Vazan, Ormel (2017)

Evaporation + Recycling remove H₂O

1D structure



Recycling envelope $\leftarrow \rightarrow$ disk



Ormel et al. (2015); Fung et al. (2015); Cimerman et al. (2017); Lambrechts & Lega (2017)

Summary

