



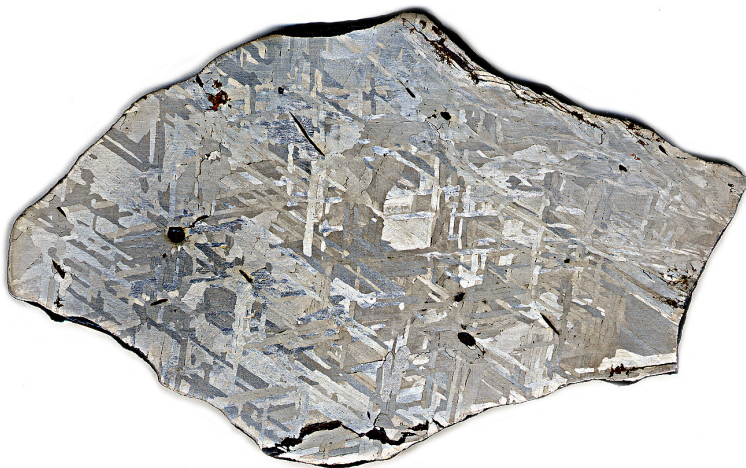
How Chondrules and CAIs Constrain Disk Processes and Planet Formation

Steve Desch

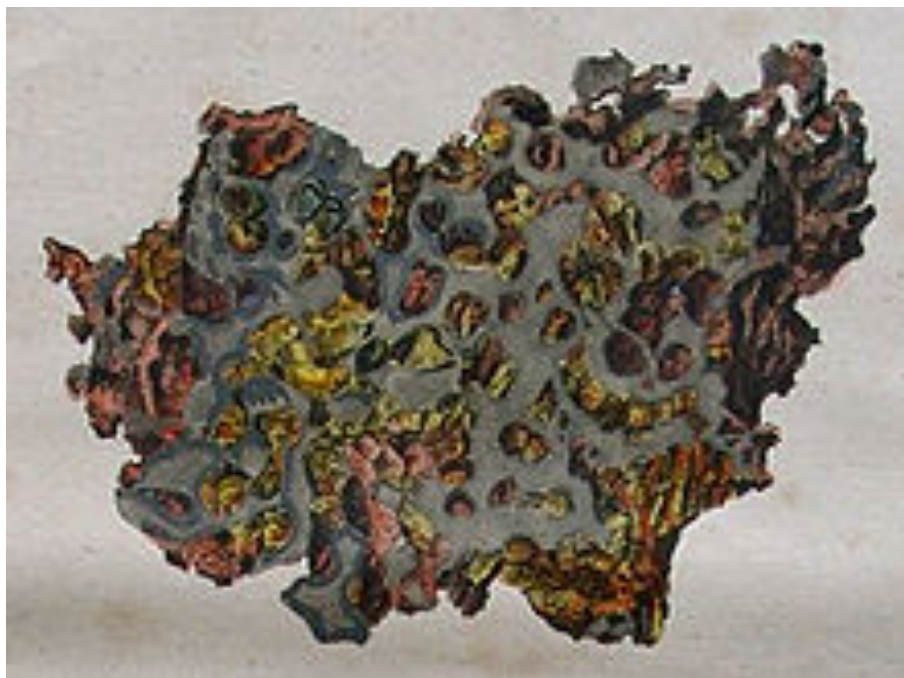
School of Earth and Space Exploration
Arizona State University

RESCEU/Planet Symposium
University of Tokyo
November 28, 2017

Iron Meteorites



Stony Achondrite
(angrite)



Pallasites

Carbonaceous Chondrites (CI Orgueil)



CM, CR, CI = water-rich,
CCs formed farther from Sun

OCs, ECs = water-poor,
formed closer to Sun

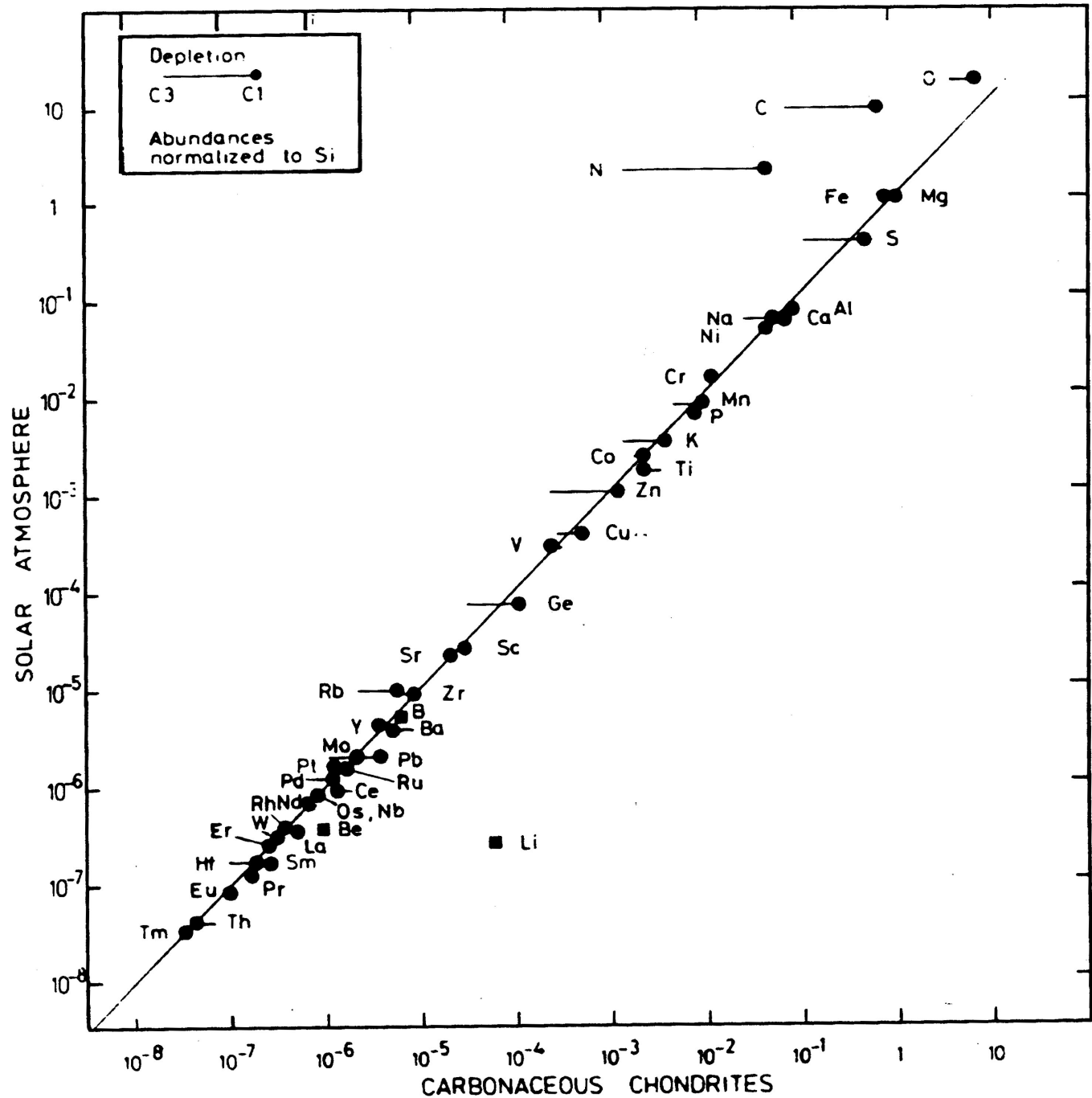
Enstatite Chondrites (EH3 SAH97079)

Ordinary Chondrites (L3-6 NWA 869)



CI Chondrites are chemically identical to the Sun (< 10-25%).

They contain 'no' refractory inclusions.



All other chondrites are full of chondrules---melted droplets of Fe,Mg silicates (olivine, pyroxene)

ORDINARY CHONDRITE



1 mm

ENSTATITE CHONDRITE



1 mm

Carbonaceous chondrites have many
Ca-rich, Al-rich inclusions (CAIs)

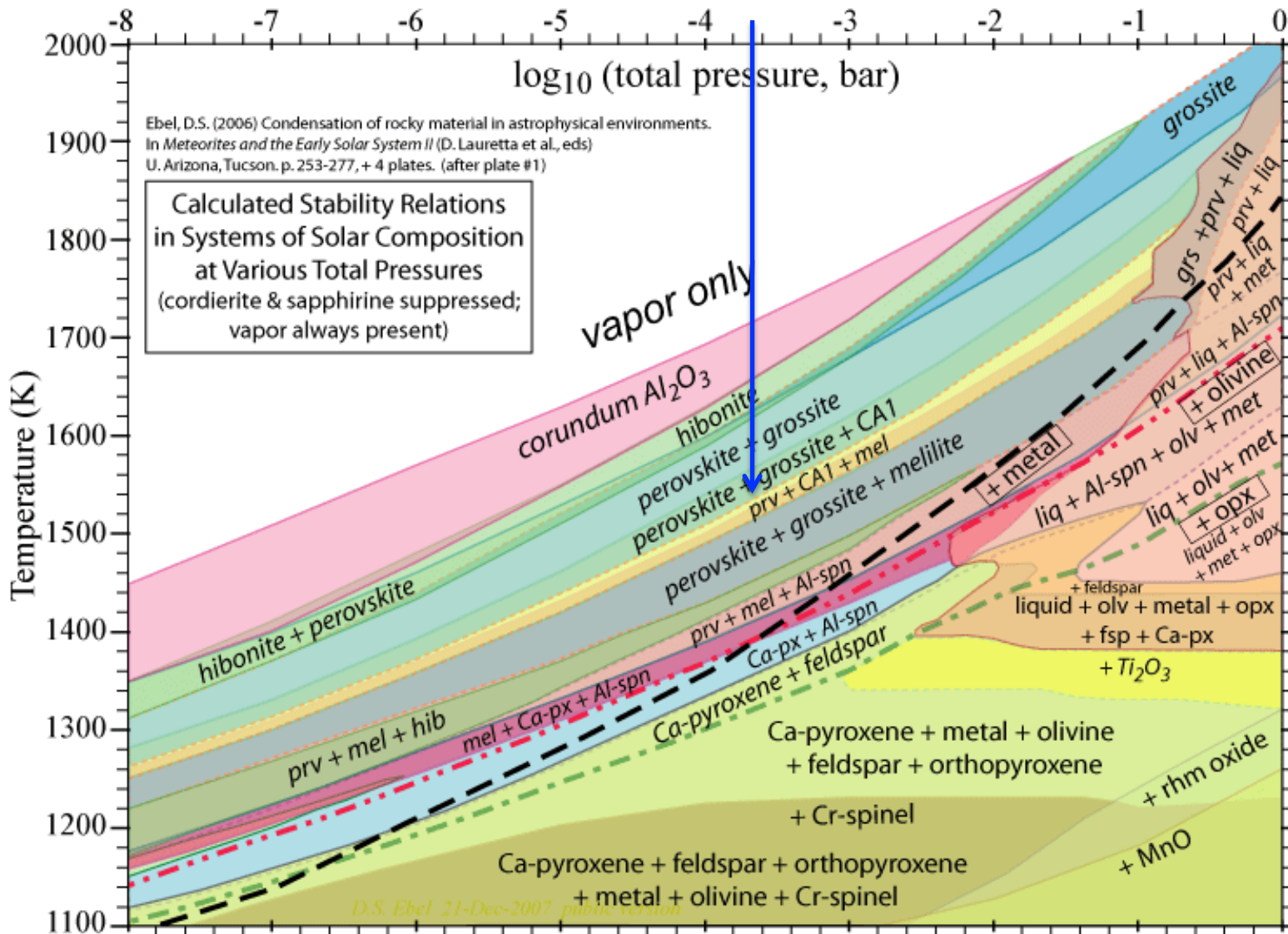
CARBONACEOUS CHONDRITE



1 mm

CAIs made of minerals that condense first from a cooling solar-composition gas.

Rich in refractory lithophiles (Ca, Al, Ti, Sc, Y, Hf, Zr, REEs)



CAIs condensed from gas at $T > 1400$ K (close to Sun)

Ebel (2006)

Table 1 of Desch, Kalyaan & Alexander, in revision (ApJ).

Why are Ordinary and Enstatite chondrites depleted in refractories and CAIs?

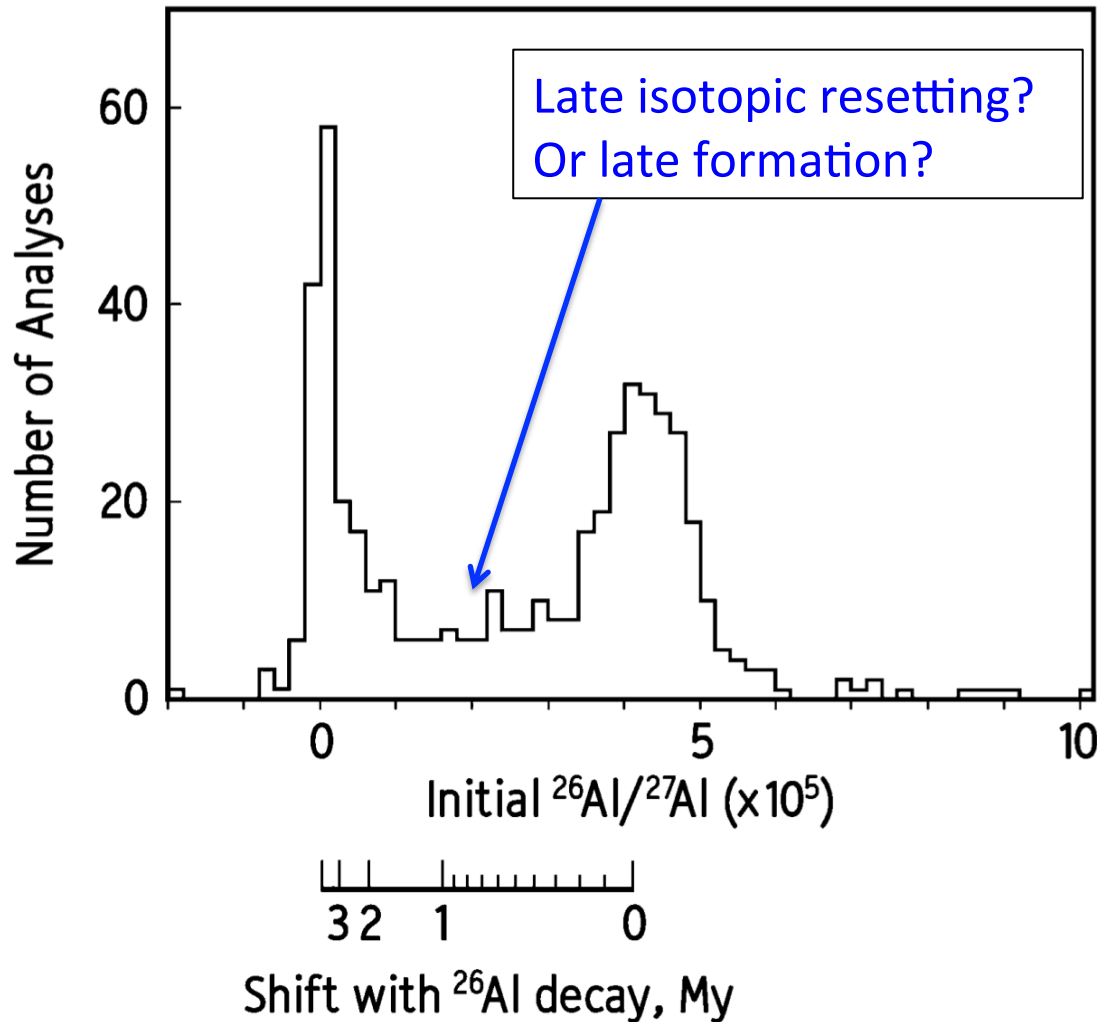
Why are Carbonaceous chondrites enriched?

Type	(X/Mg) / CI ^a	(X/Mg) / CI ^b	CAIs (vol%) ^a	CAIs (vol%) ^b	CAI (wt%) ^c	Adopted CAIs (wt%)
EH	0.87	0.884	< 0.1	0.01	0	< 0.1
EL	0.83	0.871	< 0.1	0.01	0	< 0.1
R	0.95	0.974	< 0.1	0.04	0	< 0.1
H	0.93	0.899	0.01 - 0.2	0.02	0	< 0.2
L	0.94	0.904	< 0.1	0.02	0	< 0.1
LL	0.90	0.890	< 0.1	0.02	0	< 0.1
CK	1.21	1.24	0.2	4	4	< 4
CV	1.35	1.35	3.0	3.0	6	> 3
CO	1.13	1.11	1.0	1.0	5	> 1
CM	1.15	1.13	1.2	1.2	2	1-2
CR	1.03	1.02	0.12	0.6	1	0.5-1
CI	≡ 1.00	1.00	0.00	0.0	0	0.0

a) Scott & Krot (2014); b) Rubin (2011); c) estimate based on Trinquier et al. (2009)

CAIs define “time zero” of solar nebula

Most CAIs had live ^{26}Al . Consistent with uniform $^{26}\text{Al}/^{27}\text{Al} = 5.23 \times 10^{-5}$ in solar nebula at $t=0$, most CAIs forming in < 0.4 Myr or less.



MacPherson et al. (2012, 2017); Kita et al. (2013); Ushikubo et al. (2017).

Or < 0.02 Myr? (Thrane et al. 2006)

MacPherson et al. (1995)

Type	chondrule ages (Myr)	SF14 accretion time (Myr) ^a	Other	Adopted Time of Accretion (Myr)
angrites		0.5 ± 0.4	$\leq 1.5^b$	~ 0.5
HEDs		0.8 ± 0.3	$\leq 2.2^{+1.1c},$ $\leq 0.6^{+0.5d}, < 1^e$	~ 0.8
acapulcoites-lodranites		1.3 ± 0.3	$1.5 - 2.0^f$	~ 1.5
ureilites		1.0 ± 0.3	$< 1.9^{+2.2g}, \sim 1.4^h,$ $\sim 1.6^i, \sim 0.6^j$	~ 0.6
winonaites			$\sim 1.8^k$	~ 1.8
EH		1.8 ± 0.1	$\sim 2^l$	1.7 - 1.9
EL		1.8 ± 0.1		1.7 - 1.9
R		2.1 ± 0.1		2.0 - 2.2
H	1.7 ± 0.7^m	2.1 ± 0.1	$1.8 - 2.7^n, \sim 2^{m,o}$	2.0 - 2.2
L	$1.6 - 2.2^p, 2.5 - 3.0^q$	2.1 ± 0.1	$1.8 - 2.5^r$	2.0 - 2.5
LL	$\sim 1.0 - 2.5^s$ $\sim 0.7 - 2.4^t$ $\sim 1.5 - 3.0^u$ $\sim 2.0^{+0.5v}_{-0.3}$	2.1 ± 0.1		2.0 - 2.5
CK		2.6 ± 0.2		2.4 - 2.8
CV	$\sim 1.5 - 4^w,$ $2.0 - 2.5^x$	3.0 ± 0.2	$1.8 - 2.5^r$	~ 3.0
CO	$1.7 - 3.0^y$	2.7 ± 0.2	$1.8 - 2.5^r$	$\sim 2.5 - 3.0$
CM		3.5 ± 0.5	$3.0 - 4.0^z$	3.0 - 4.0
CR	2.5 ± 1.2^{aa} $3.5 - 3.8^{ab}$	3.5 ± 0.5	$> 4.0^{+0.5ab}_{-0.3}$	3.8 - 4.5
CI		3.5 ± 0.5	$3 - 4^z$	3.0 - 4.0

Table 2 of Desch, Kalyaan & Alexander, in revision (ApJ).

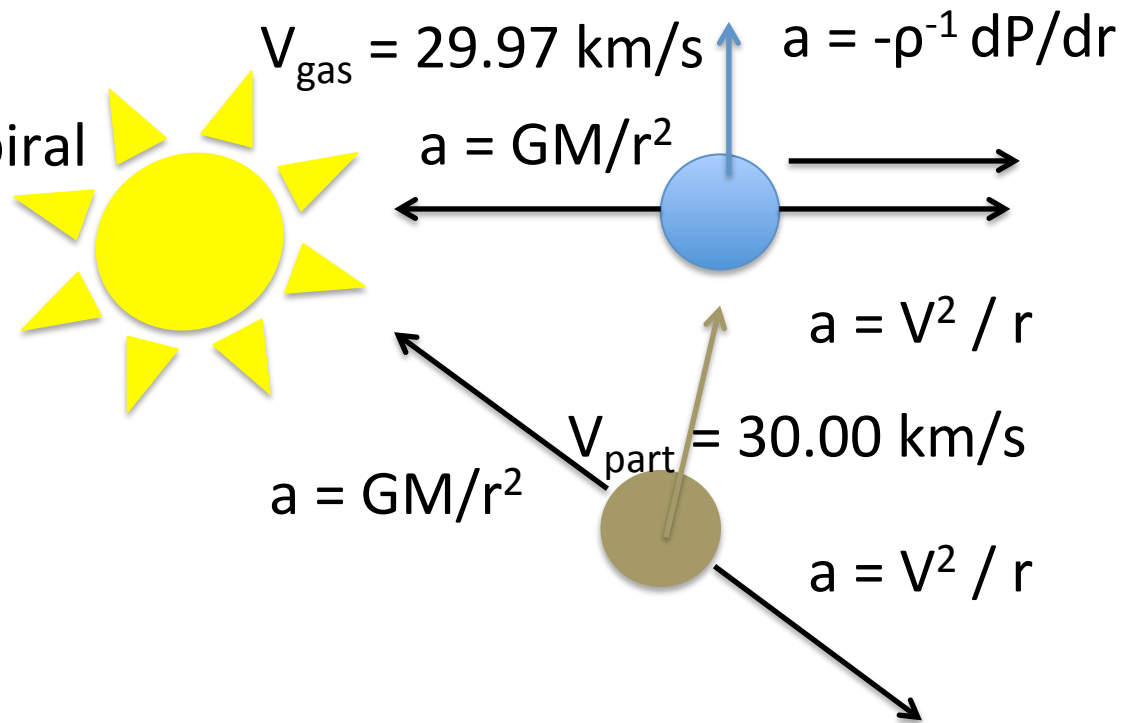
Chondrites formed 2-4 Myr after CAIs: Chondrule ages and thermal modeling, esp. Sugiura & Fujiya (2014)

a. Sugiura & Fujiya (2014); b. Hf-W: Kleine et al. (2012); c. Mn-Cr: Trinquier et al. (2008); d. Al-Mg: Schiller et al. (2011); e. Al-Mg: Neumann et al. (2014); f. Hf-W: Touboul et al. (2009) g. Al-Mg: Baker et al. (2012); h. Hf-W: von Kooten et al. (2014); i. Hf-W: Budde et al. (2015); j. Wilson & Goodrich (2016); k. Hunt et al. (2017); l. Shukolyukov & Lugmair (2004); m. Kleine et al. (2008) n. Gail et al. (2014); o. Henke et al. (2013); p. Al-Mg: Rudraswami & Goswami (2007); q. U-Pb: Bollard et al. (2017); r. Mn-Cr: Doyle et al. (2015); s. Al-Mg: Rudraswami et al. (2008); t. Al-Mg: Mostefaoui et al. (2002); u. Al-Mg: Villeneuve et al. (2009); v. Kita et al. (2000); w. Al-Mg: Hutcheon et al. (2009); x. Al-Mg: Nagashima et al. (2015); y. Al-Mg: Kurahashi et al. (2008); z. Mn-Cr: Fujiya et al. (2013); aa. Pb-Pb: Amelin et al. (2002); ab. Al-Mg: Schrader et al. (2017).

The “CAI Storage Problem”

CAIs are big enough to spiral in to the Sun in < 1 Myr (Weidenschilling 1977)

How did they remain in nebula for > 3 Myr to accrete into chondrites?



If they have to end up in any chondrite, why carbonaceous chondrites, which form farthest from the Sun? Why are they almost absent from ordinary and enstatite chondrites?

Hints from Isotopes

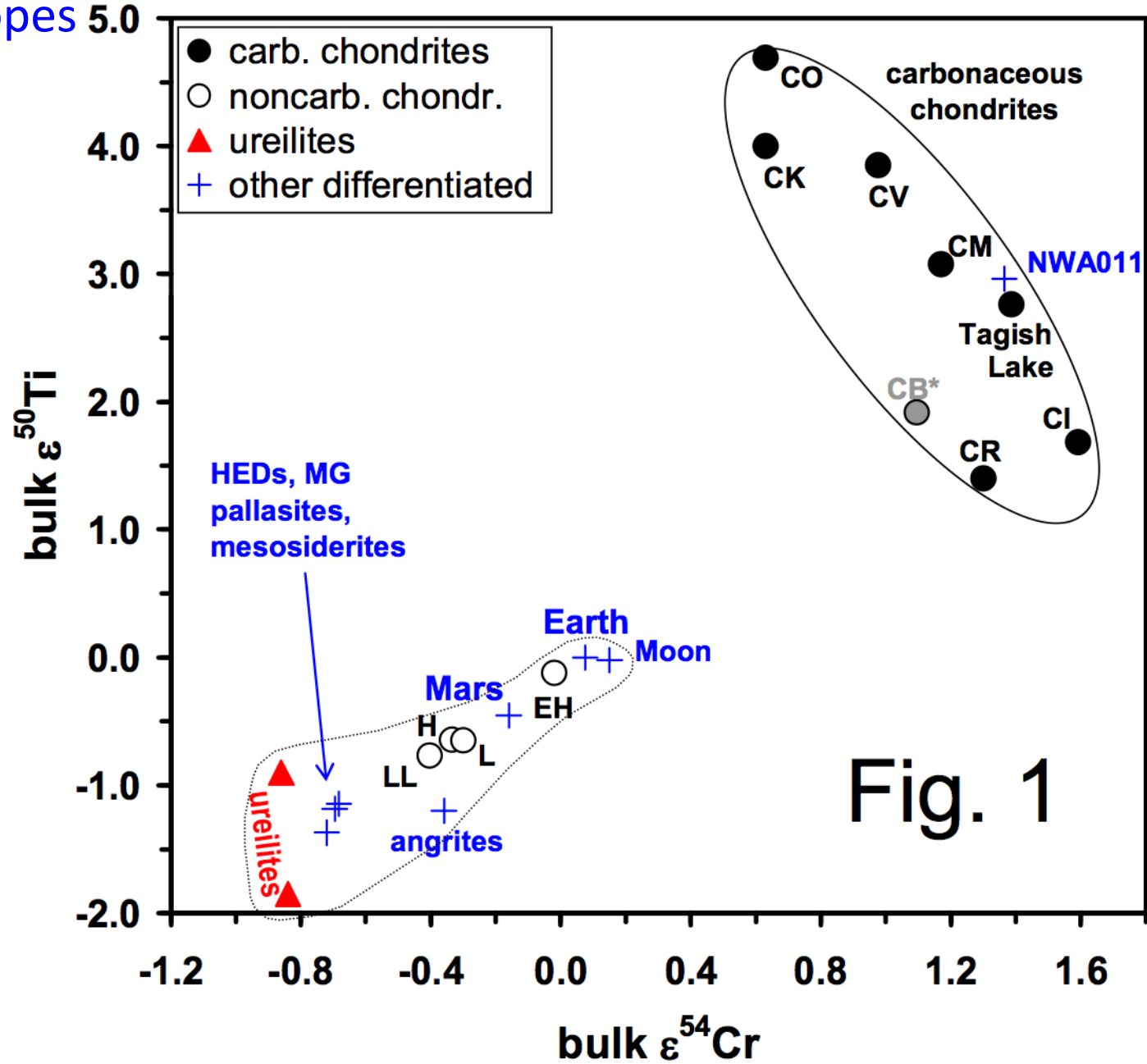
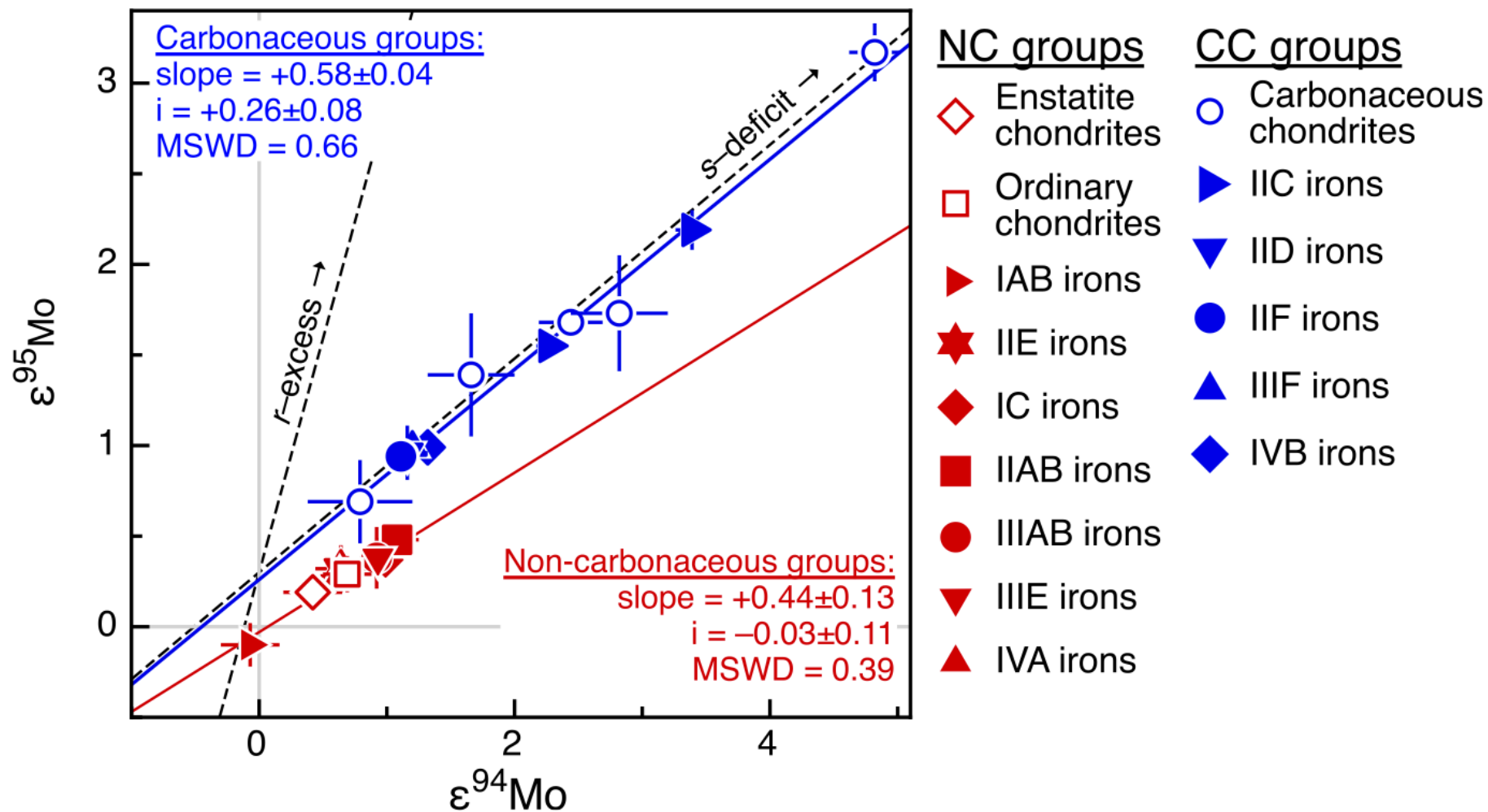


Fig. 1

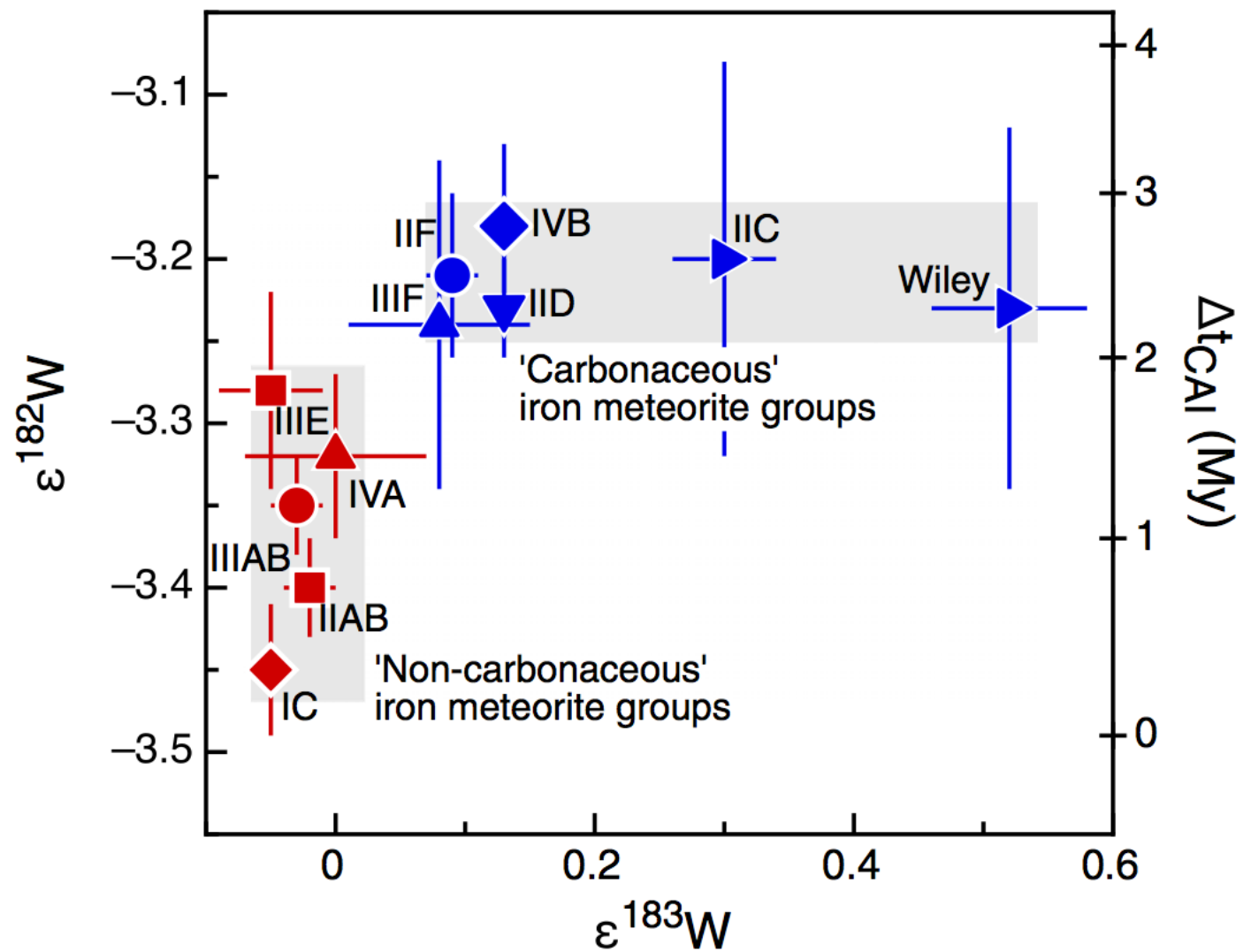
Hints from Isotopes



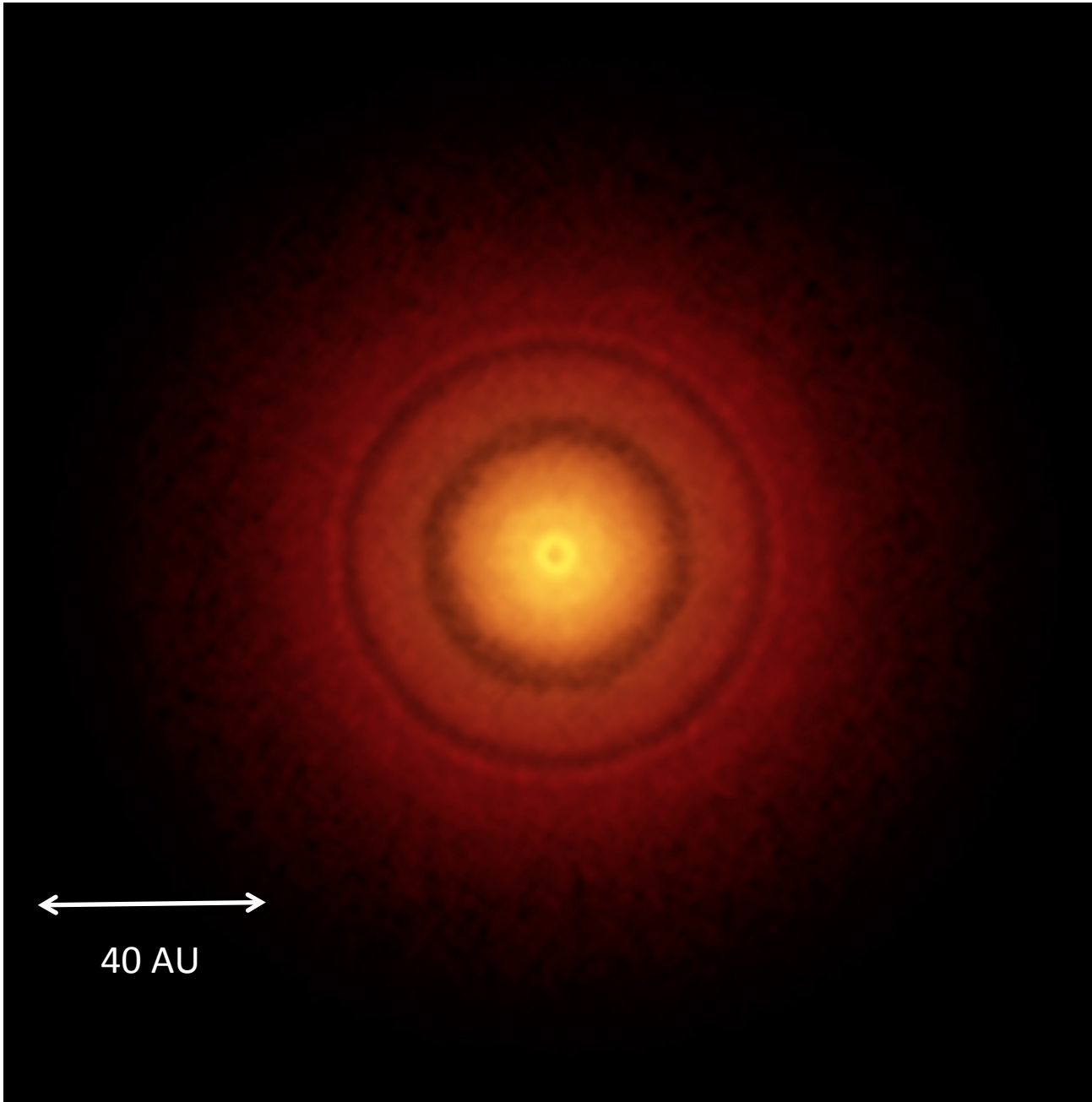
Kruijer et al. (2017)

Hints from Isotopes

Solar nebula split into two reservoirs
between 0.4 and 0.9 Myr



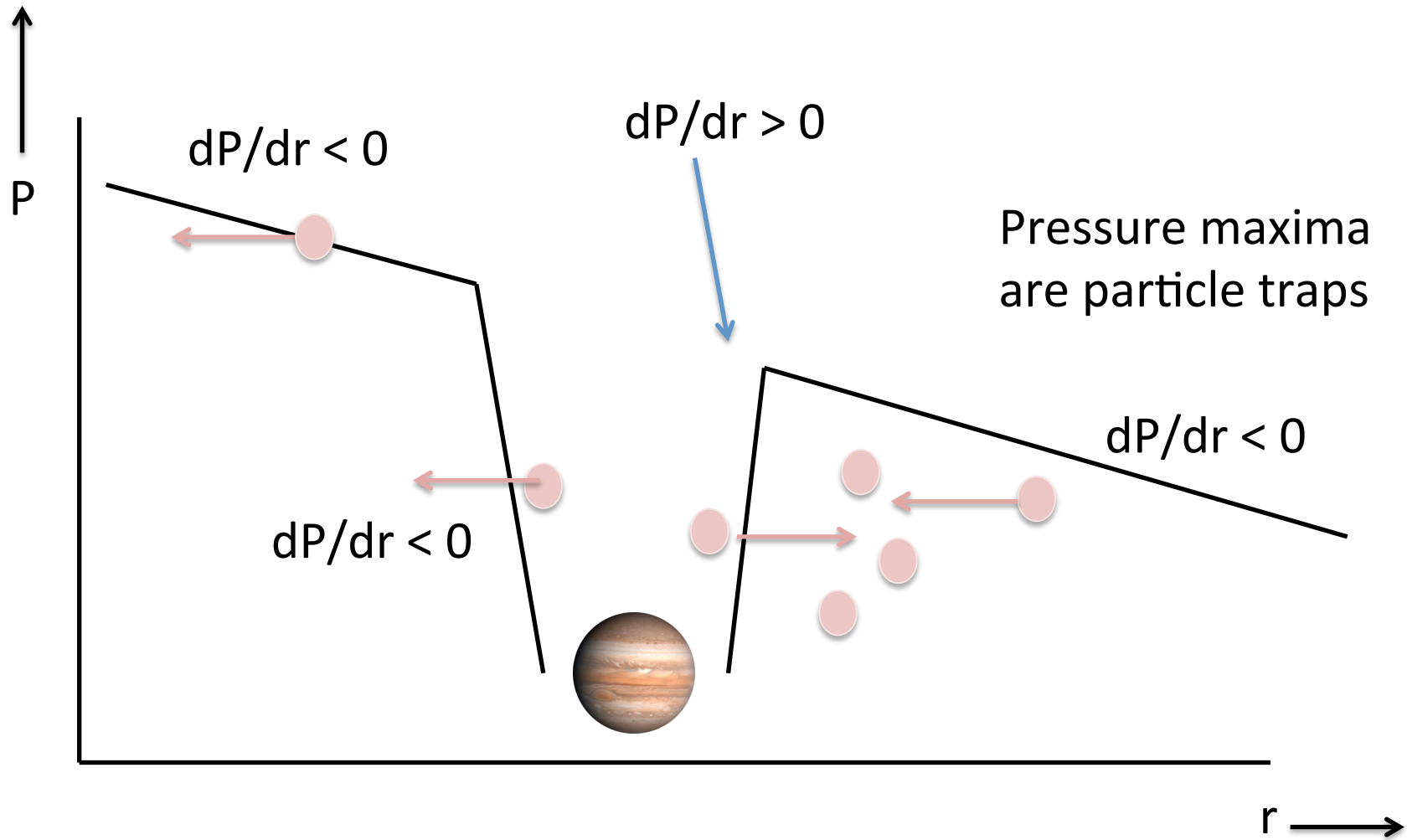
Hints from Astronomical Observations



ALMA millimeter observations are revealing many protoplanetary disks have gaps and rings

Andrews et al. (2016)

Hints from Astronomical Observations





Morbidelli et al. (2016) suggested Jupiter prevented flow of ice across gap.

Kruijer et al. (2017) suggested Jupiter divided solar system into two isotopic reservoirs.

Maybe Jupiter trapped CAIs...

A Comprehensive Model

Similar to Yang & Ciesla (2012)

Standard 1-D alpha disk with accretional heating (Min et al. 2011) and opacities (Semenov et al. 2003), initialized with compact ($R_1=1$ AU) self-similar profile (Hartmann et al. 1998). T capped at 1400 K.

Jupiter forms at 3 AU at 0.6 Myr. Starts at $30 M_E$, grows by accreting gas in Hill sphere.

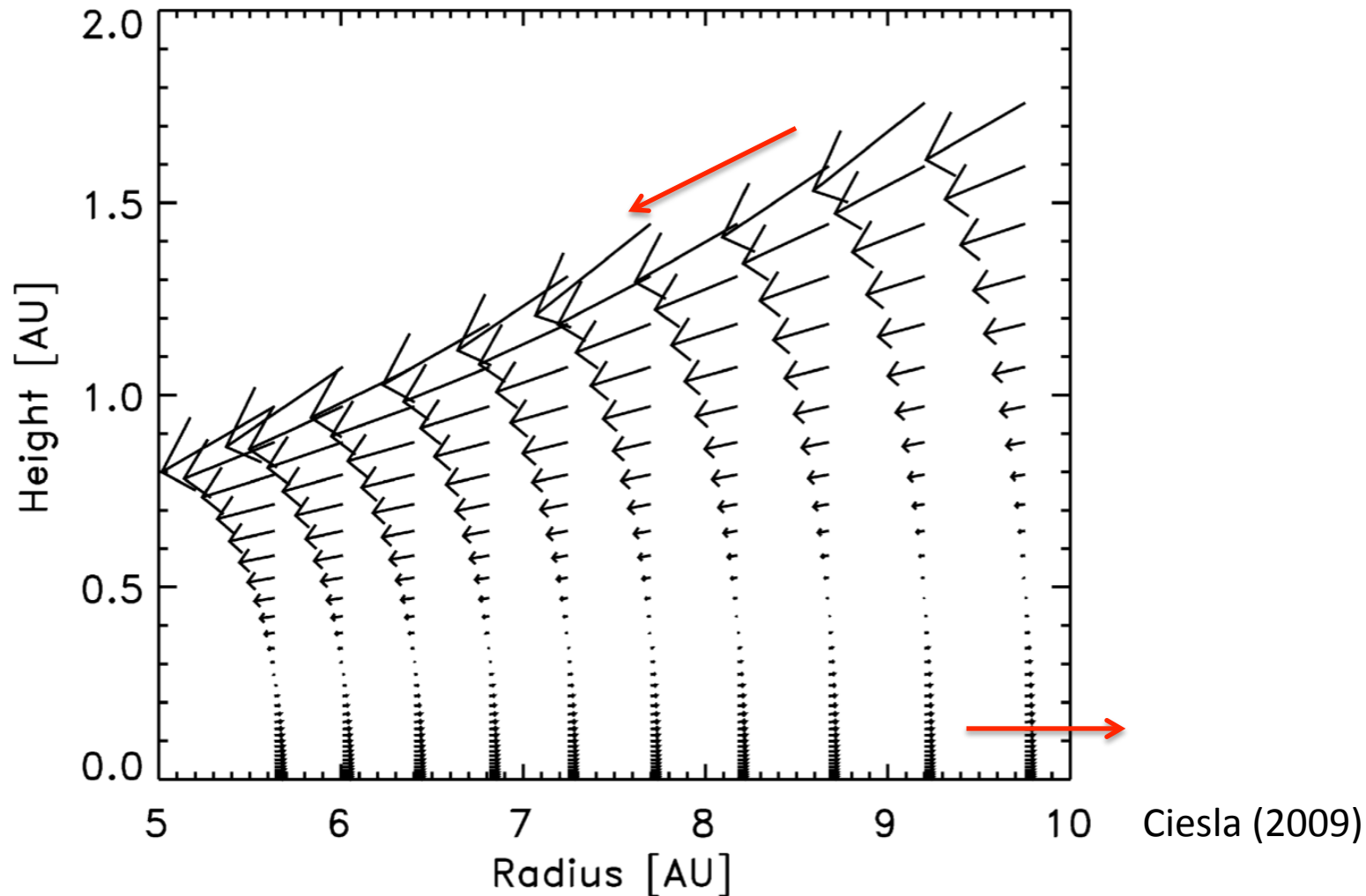
$\alpha = 1.25 \times 10^{-4}$ for $r < 1$ AU, transitioning to $\alpha = 1 \times 10^{-5}$ for $r > 10$ AU.
 $\alpha = 1 \times 10^{-2}$ in vicinity of Jupiter.

Starting composition like CI chondrites. If material ends up in region with $T > 1400$ K, 3.6% of rocky material converted into CAIs of various sizes (25% $a=800 \mu\text{m}$, 75% $a=1600 \mu\text{m}$).

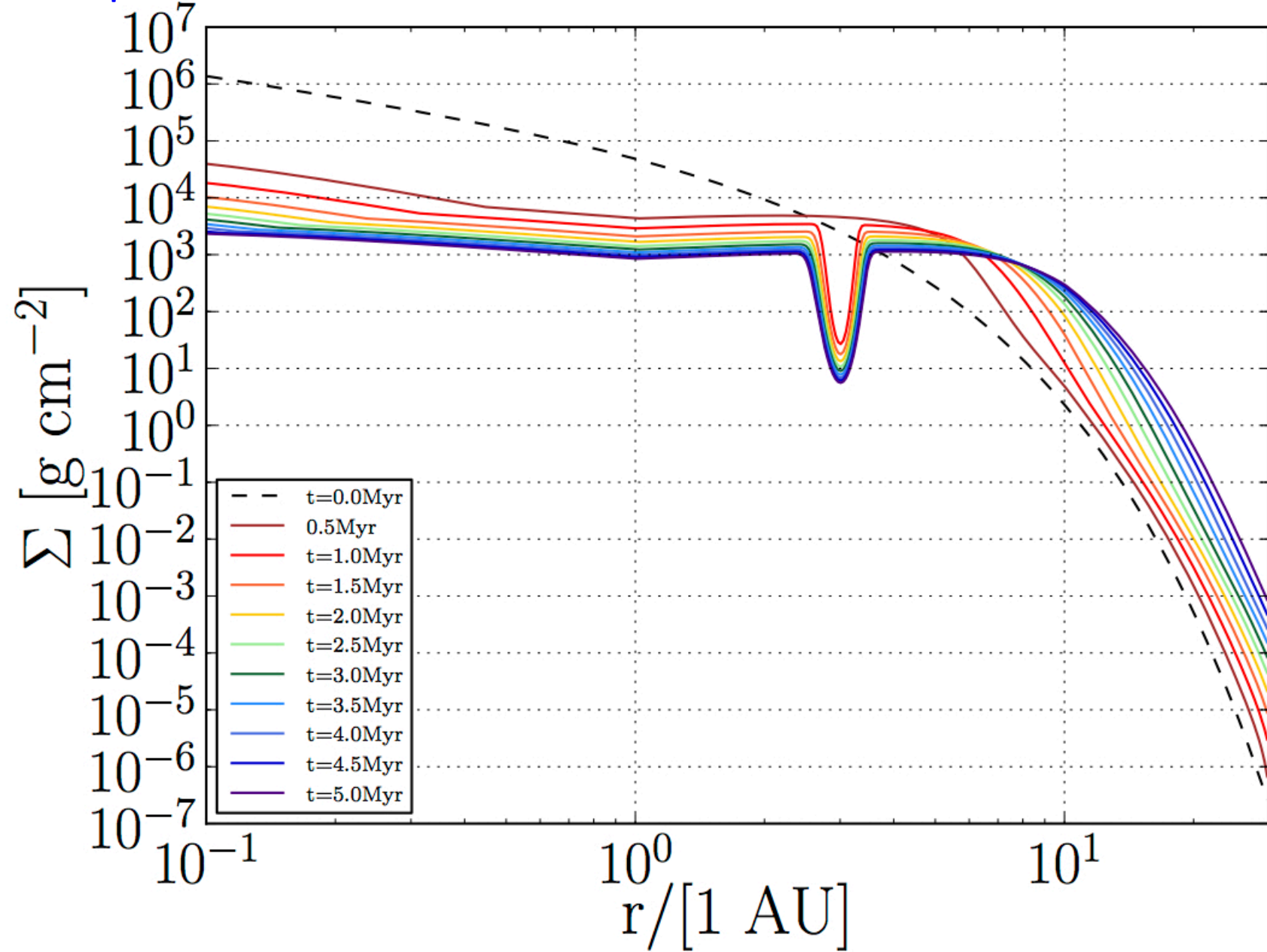
Radial transport of volatiles and particles (Desch et al. 2017).
Meridional transport included using Philippov & Rafikov (2017).

A Comprehensive Model

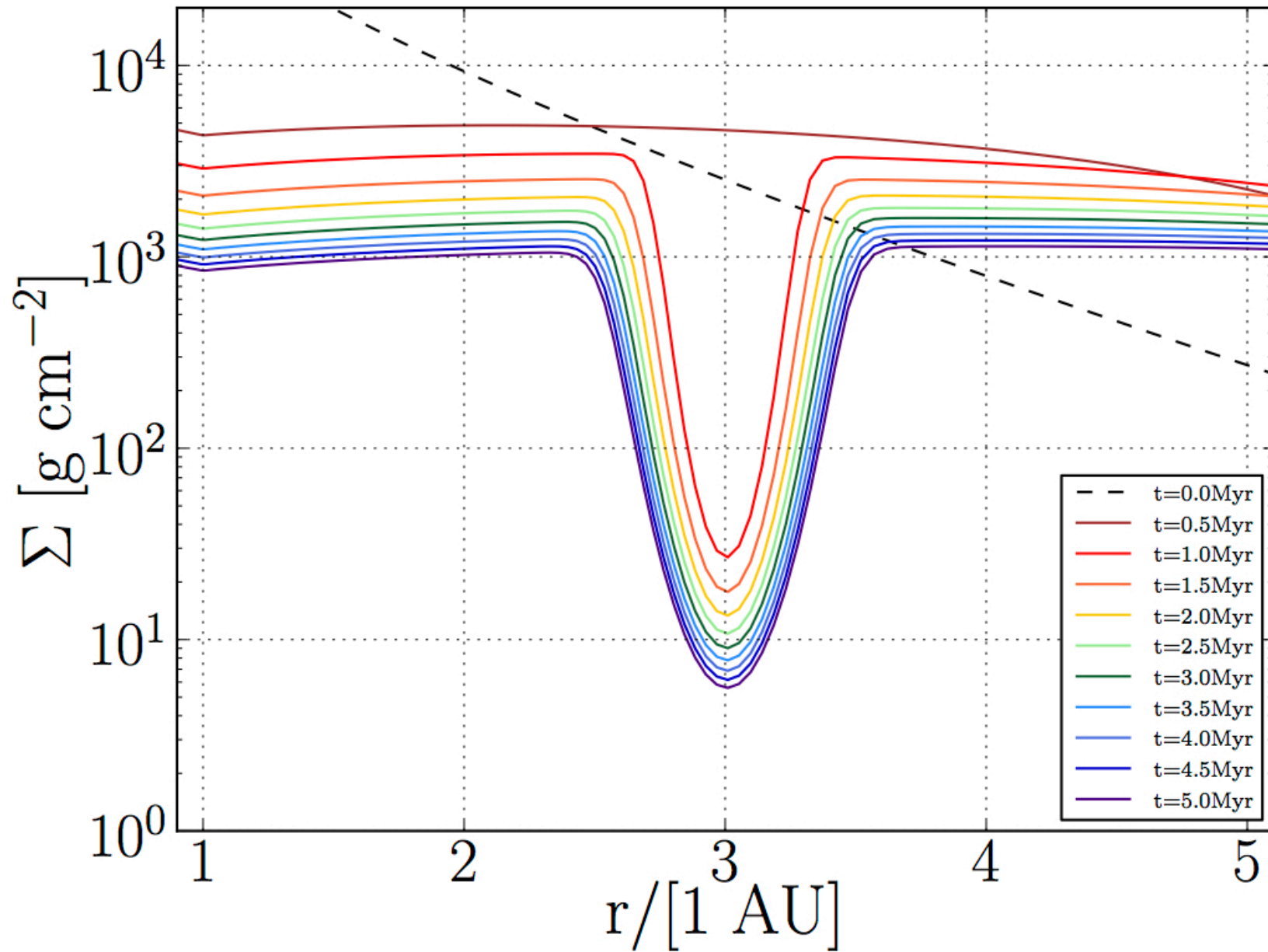
Particles participate in meridional flow if settled to midplane:
requires $St = \Omega t_{\text{stop}} > 0.2 \alpha$, particles $> 100 \mu\text{m}$



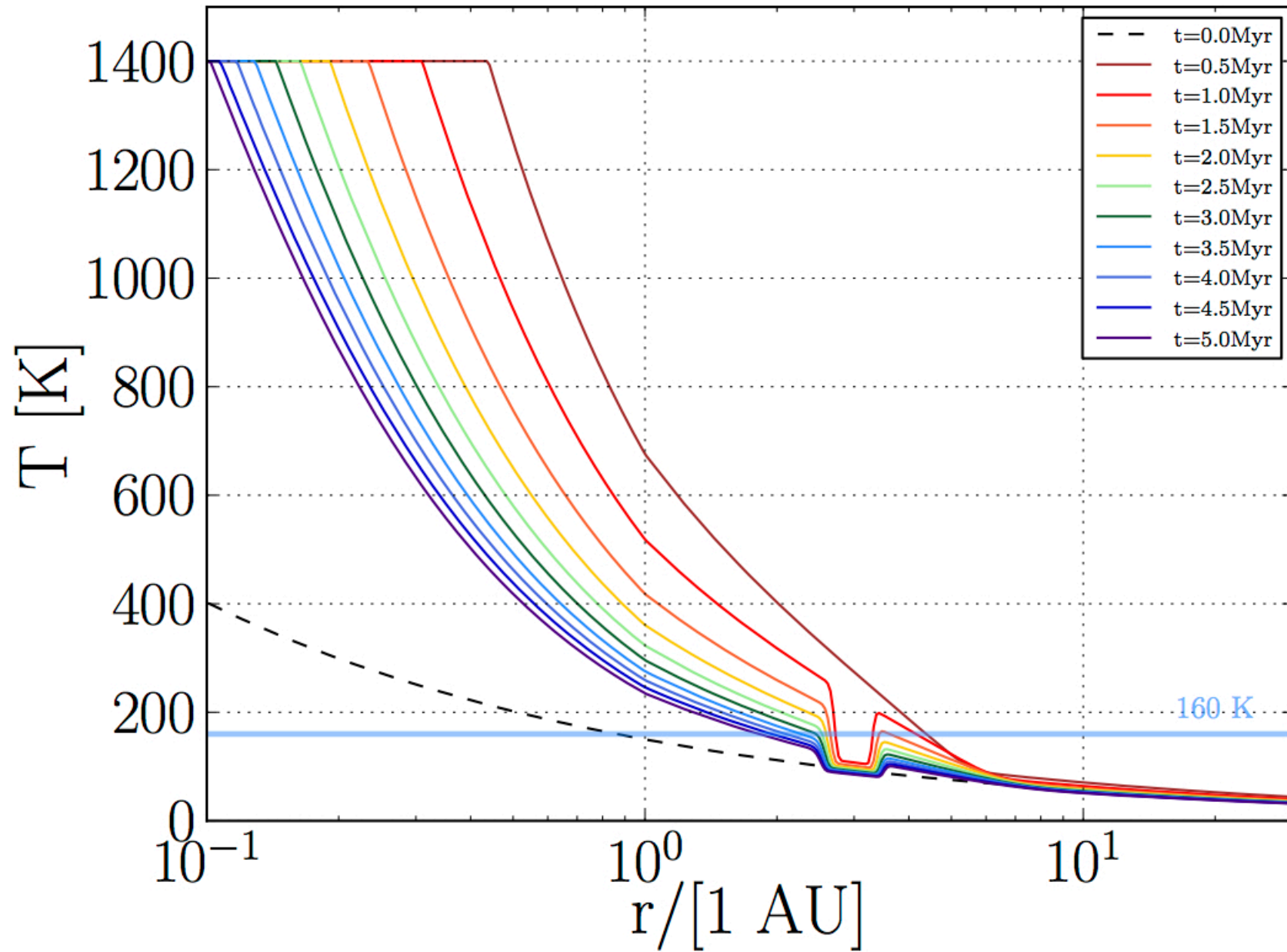
A Comprehensive Model



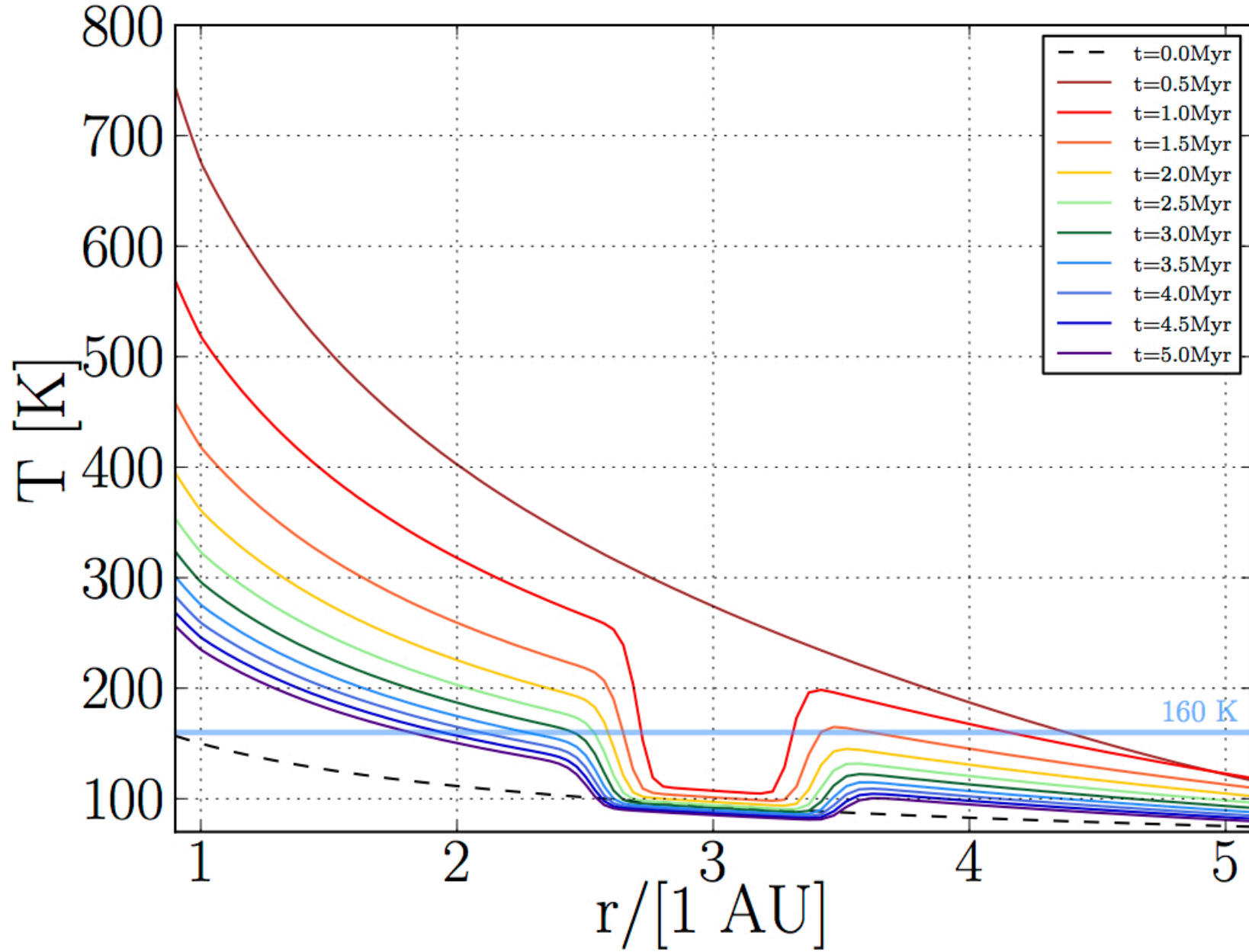
A Comprehensive Model



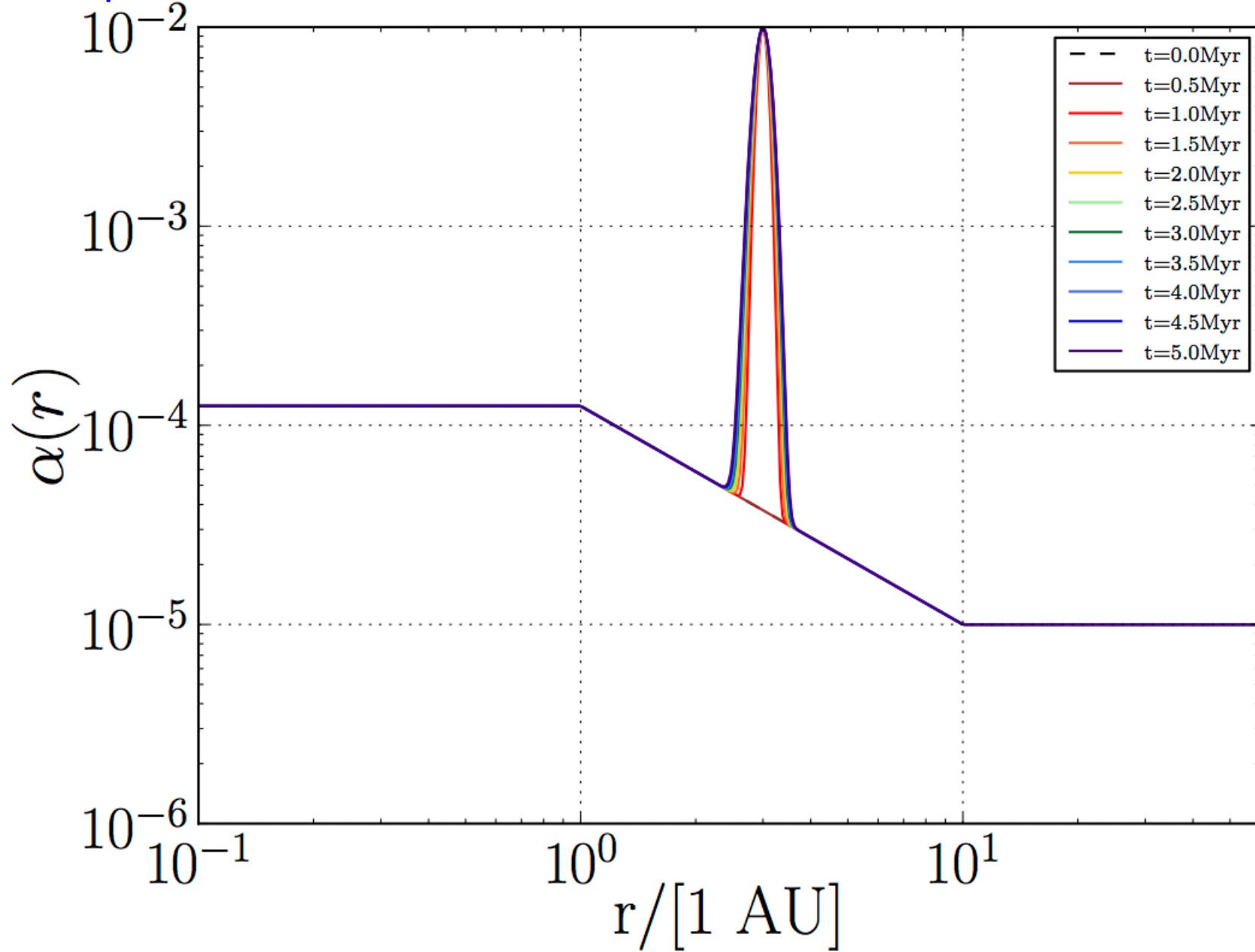
A Comprehensive Model

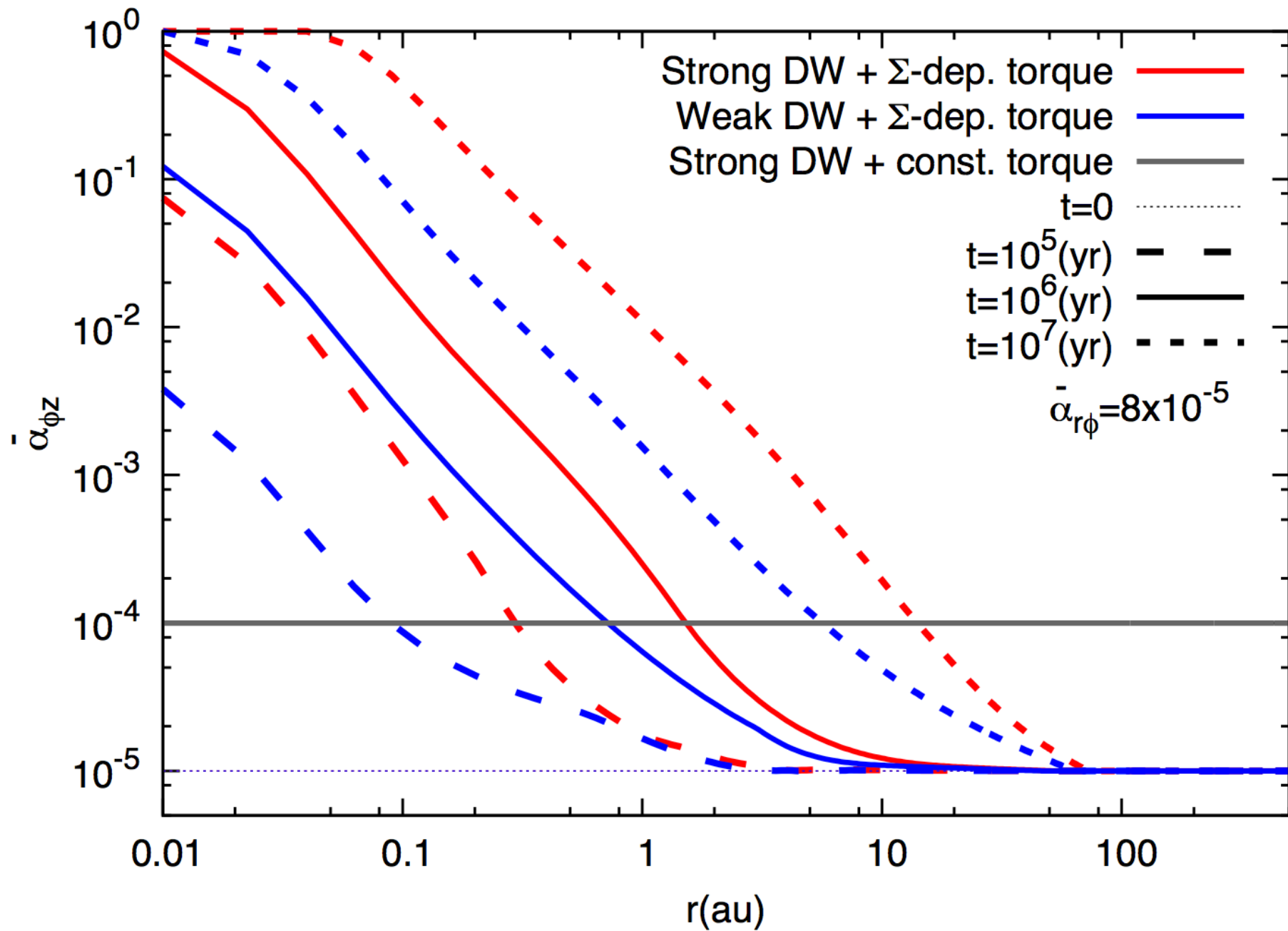


A Comprehensive Model

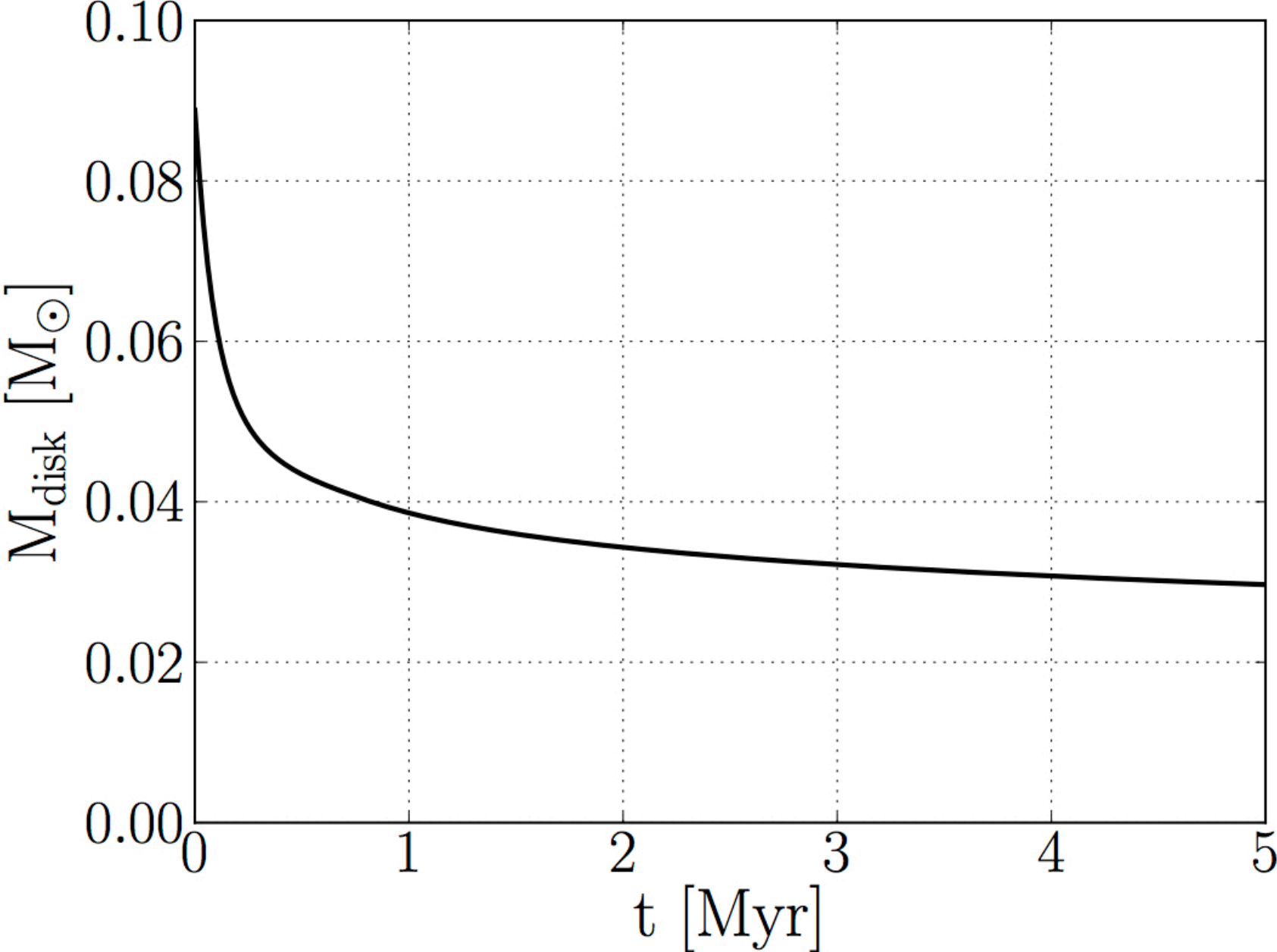


A Comprehensive Model

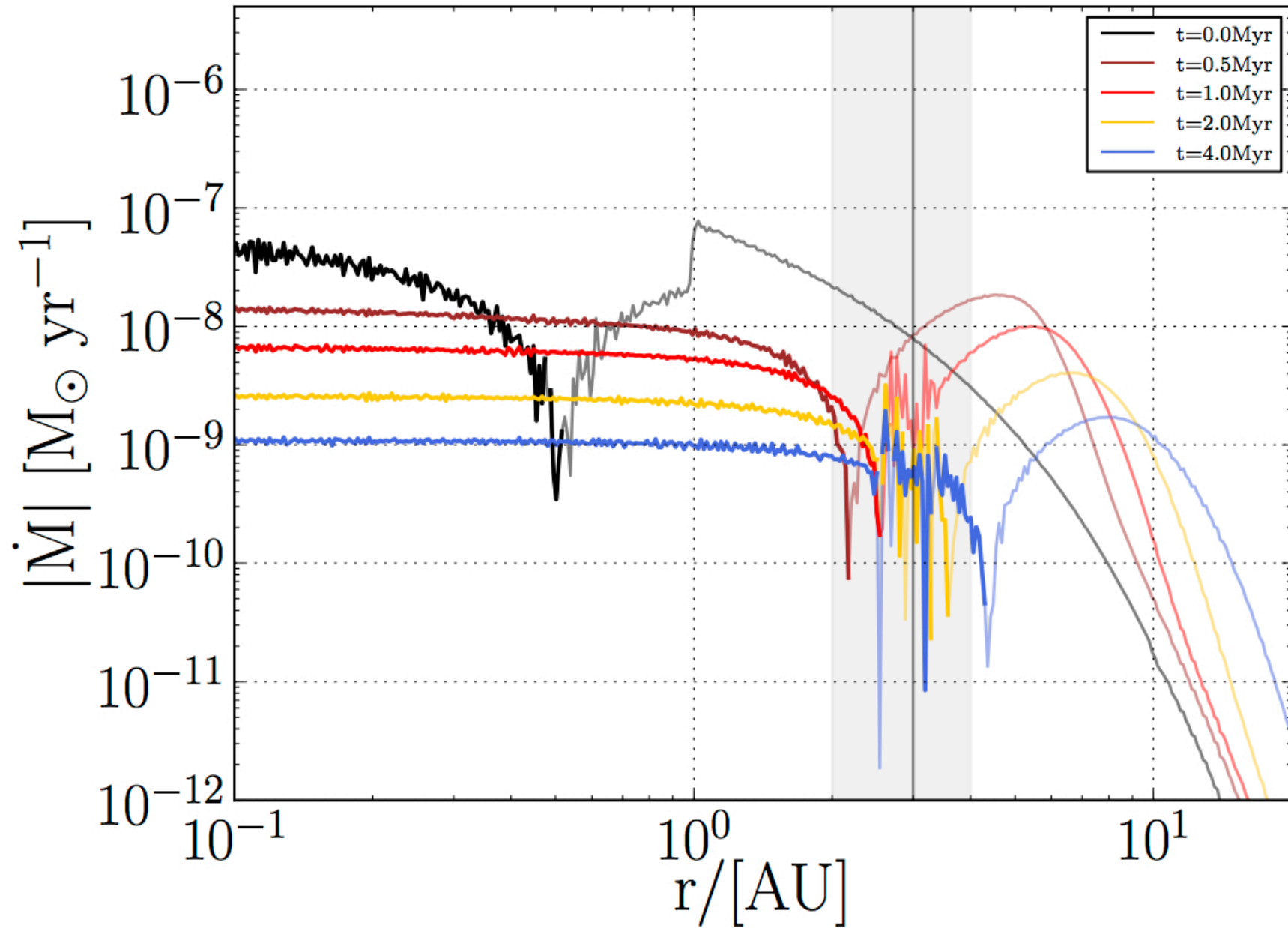




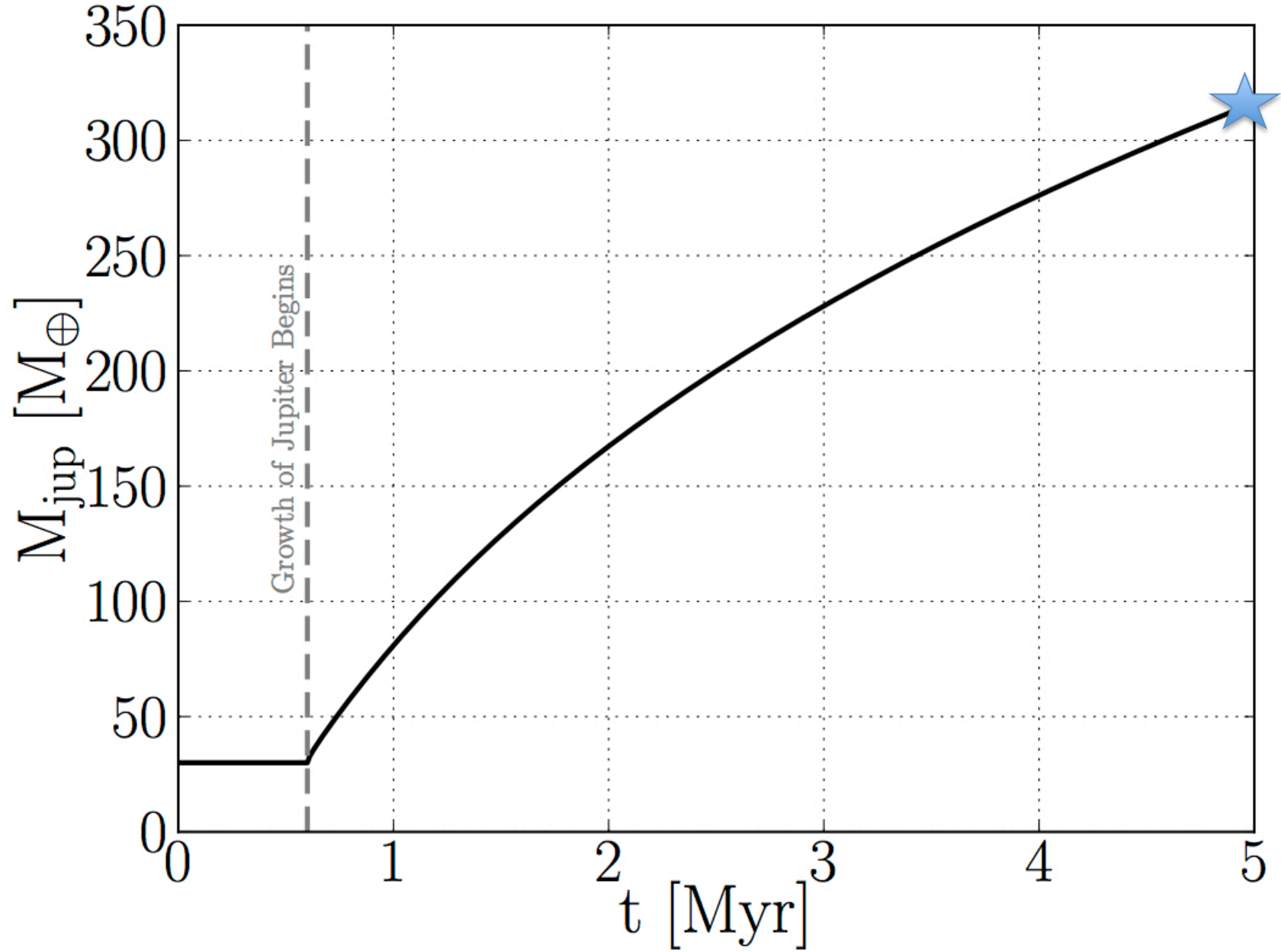
A Comprehensive Model



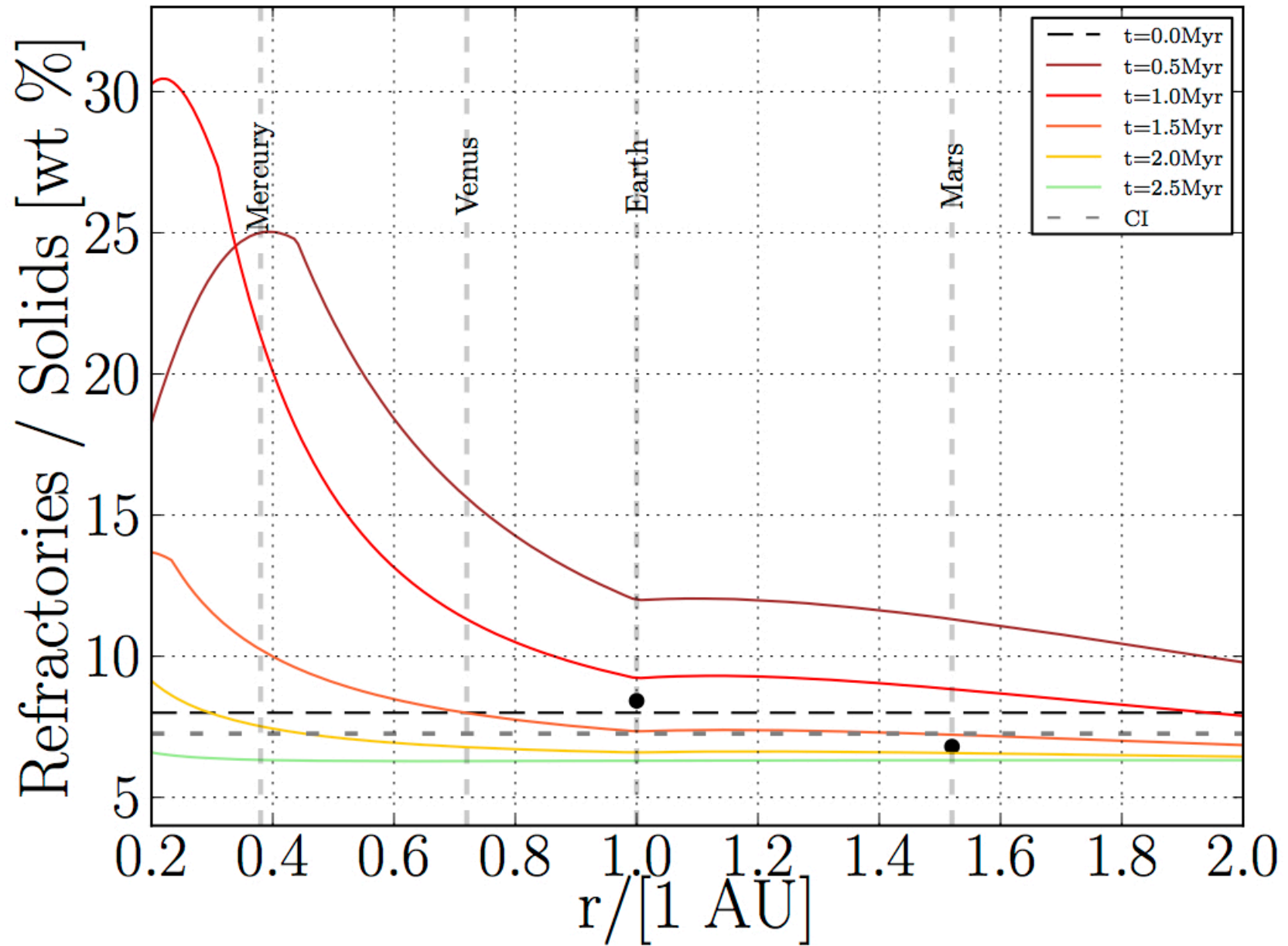
A Comprehensive Model



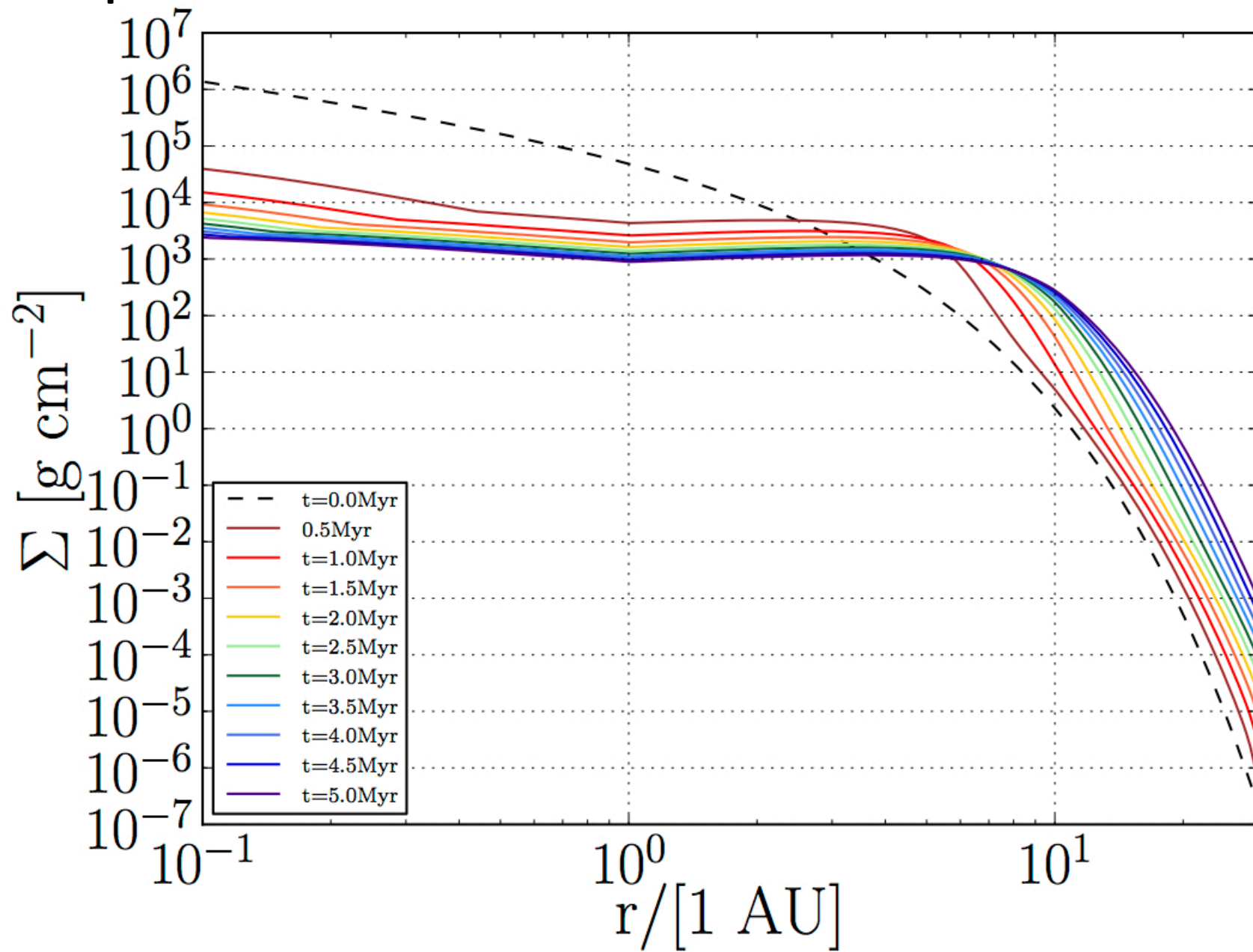
A Comprehensive Model



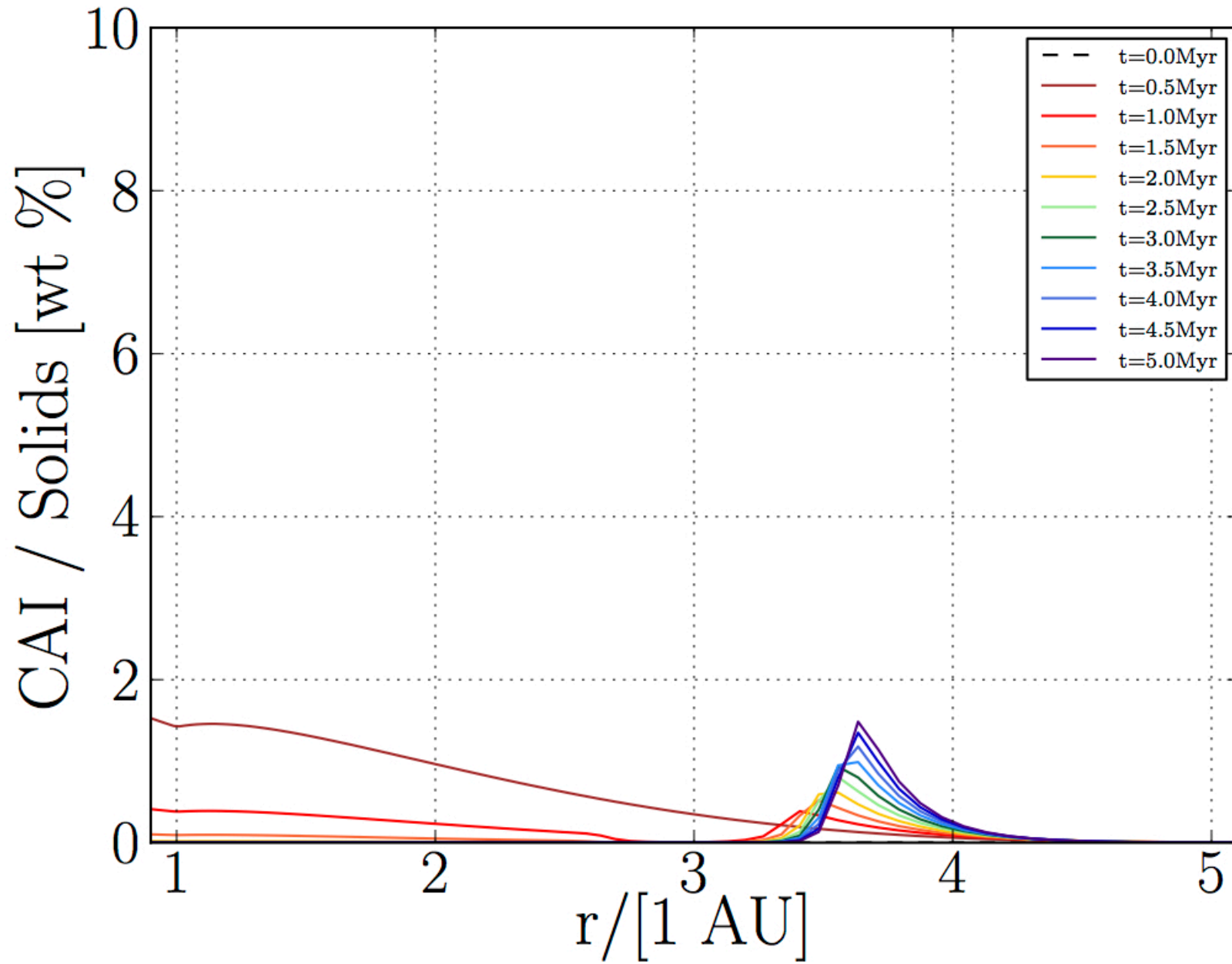
A Comprehensive Model



No Jupiter

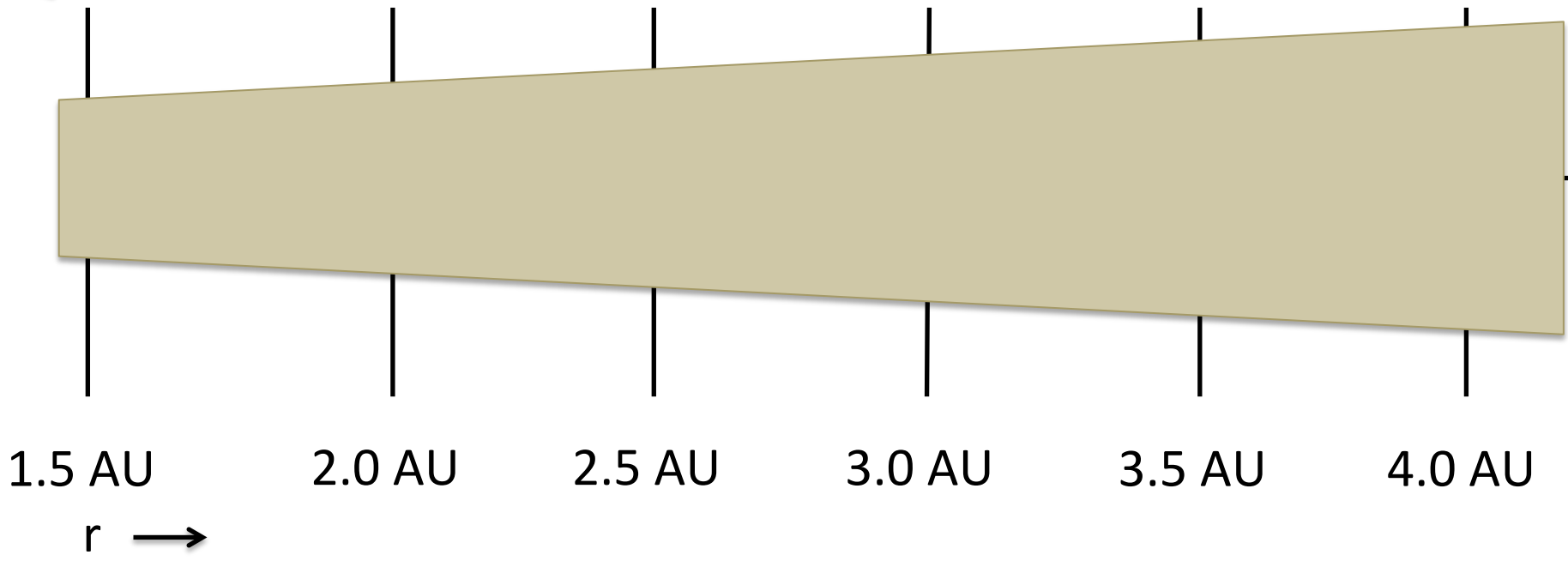


No Meridional Flow



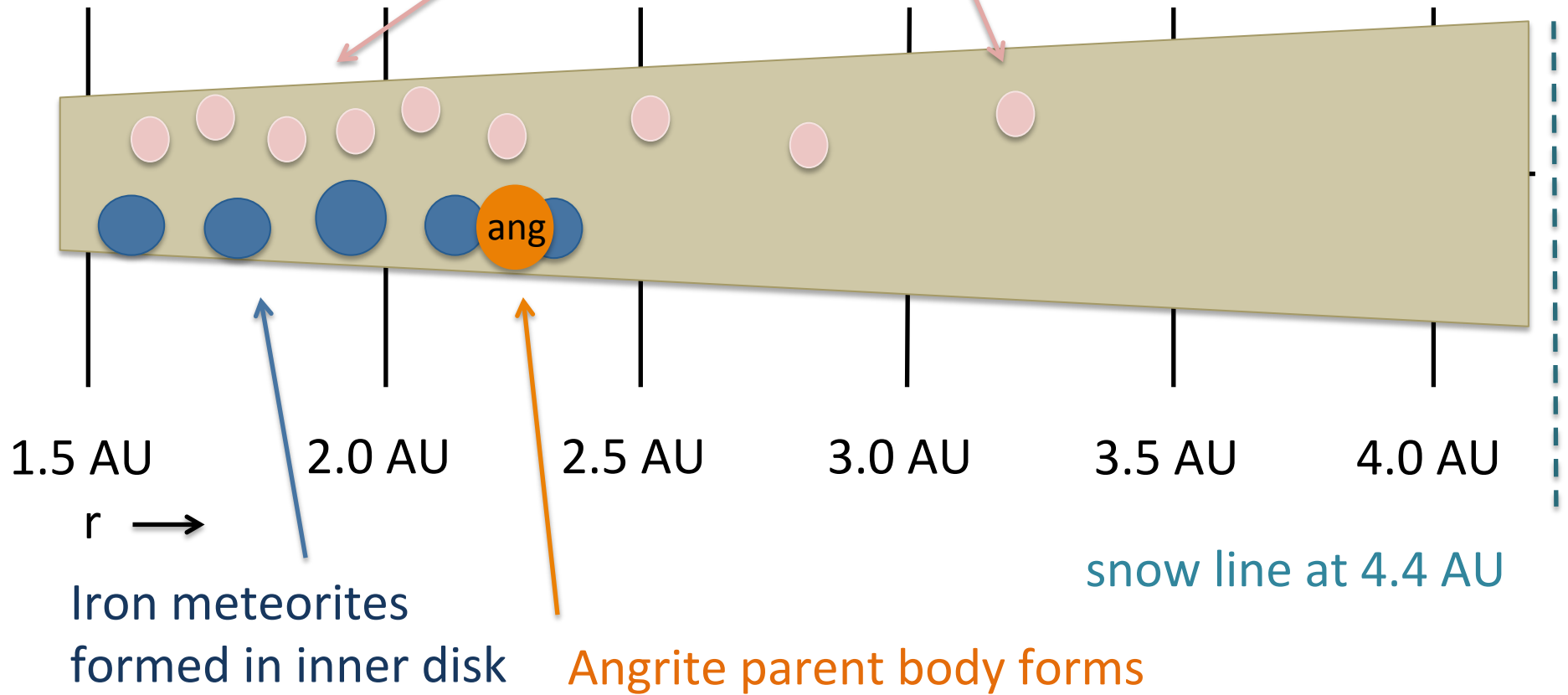
t = 0.0 Myr

CAIs created in CAI Factory < 1 AU



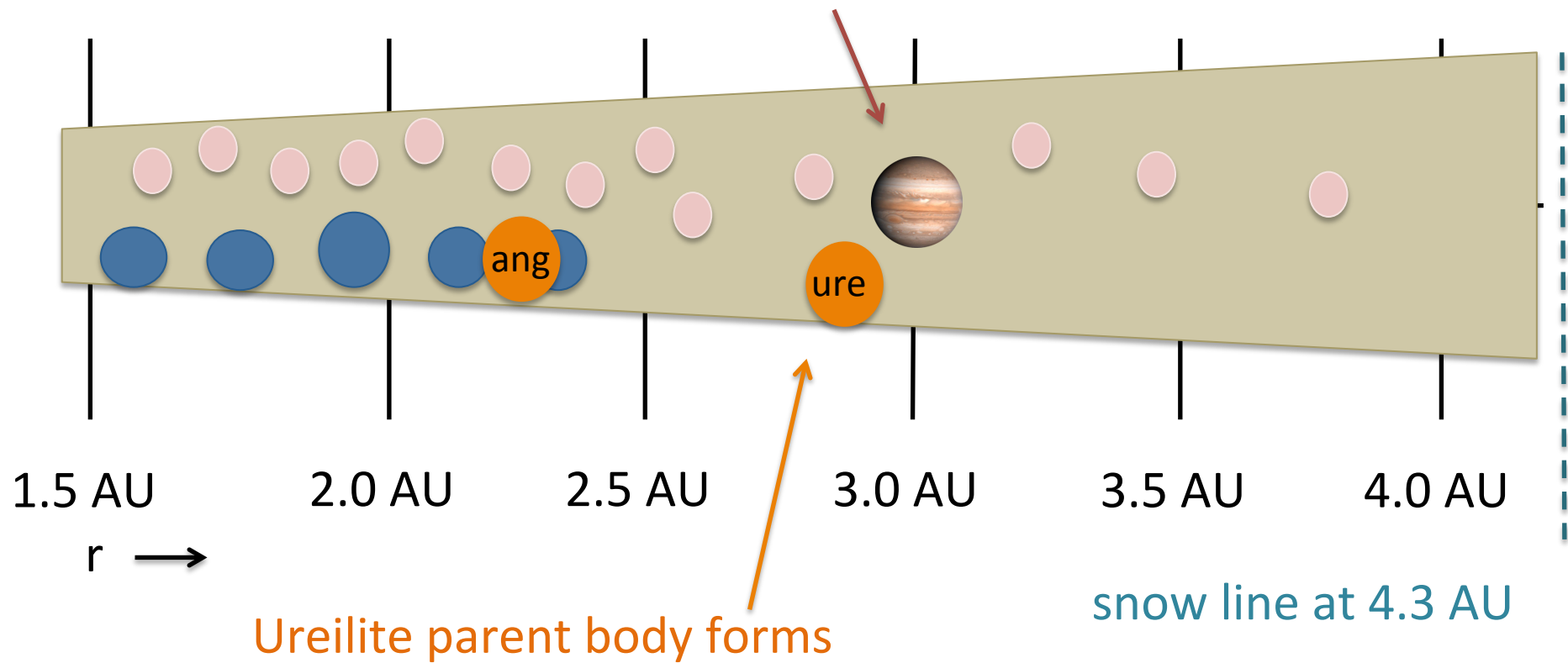
t = 0.4 Myr

CAIs transported from CAI Factory < 1 AU



t = 0.6 Myr

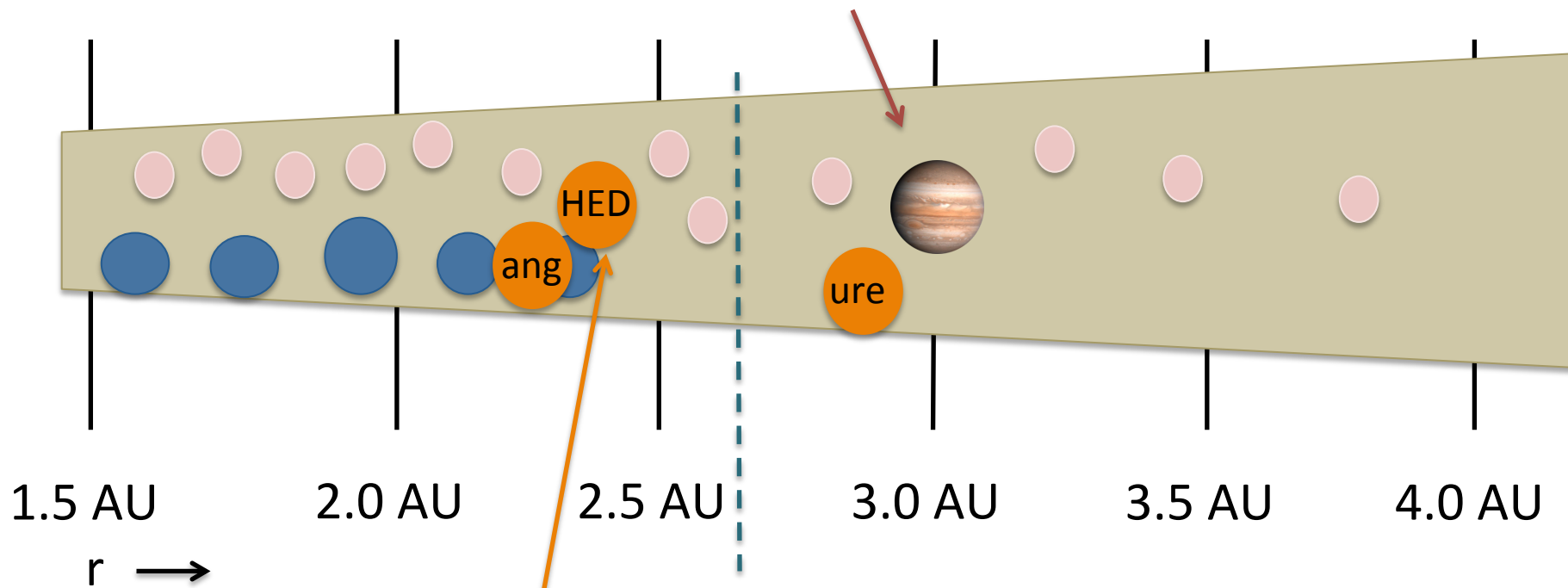
Jupiter's core forms
 $M_{\text{jup}} \sim 30 M_{\text{E}}$



snow line at 4.3 AU

t = 0.8 Myr

Jupiter's core forms
 $M_{\text{Jup}} \sim 30 M_E$



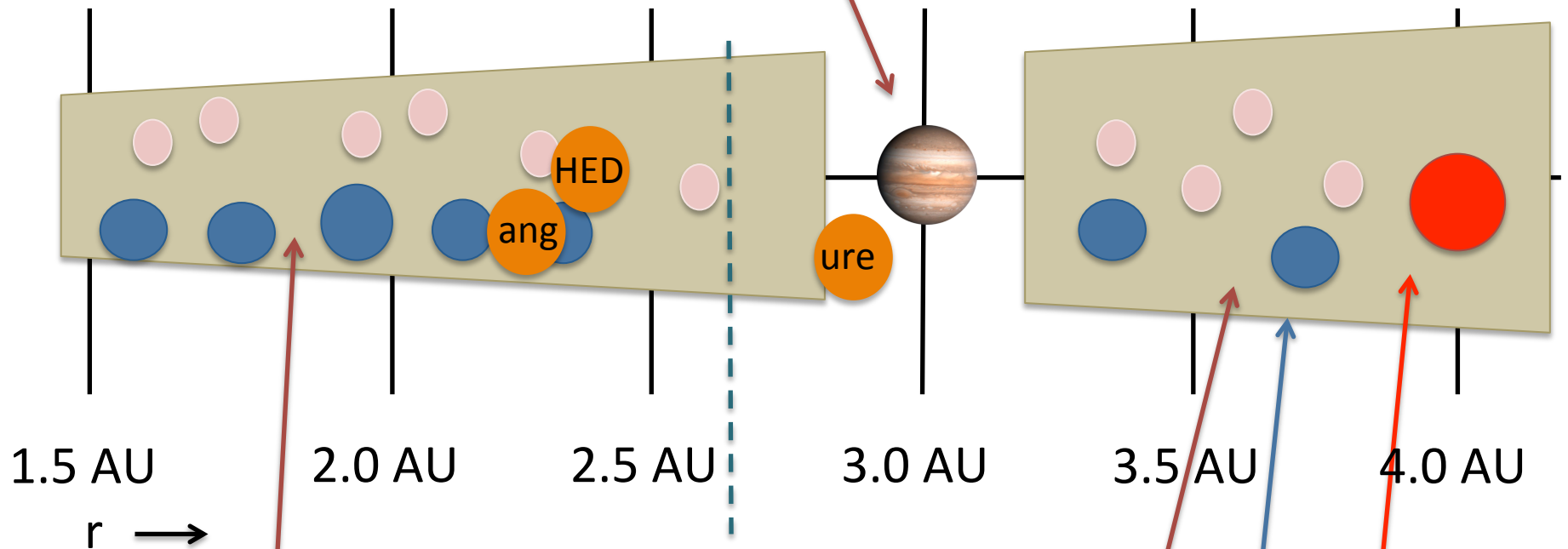
HED parent body
(4 Vesta) forms

snow line at 2.7 AU

t = 1.0 Myr

Jupiter accretes, opens gap.
 $M_{\text{jup}} \sim 80 \text{ ME}$

“PRESSURE BUMP”



CAIs drift inward, lost from inner disk

snow line at 2.7 AU

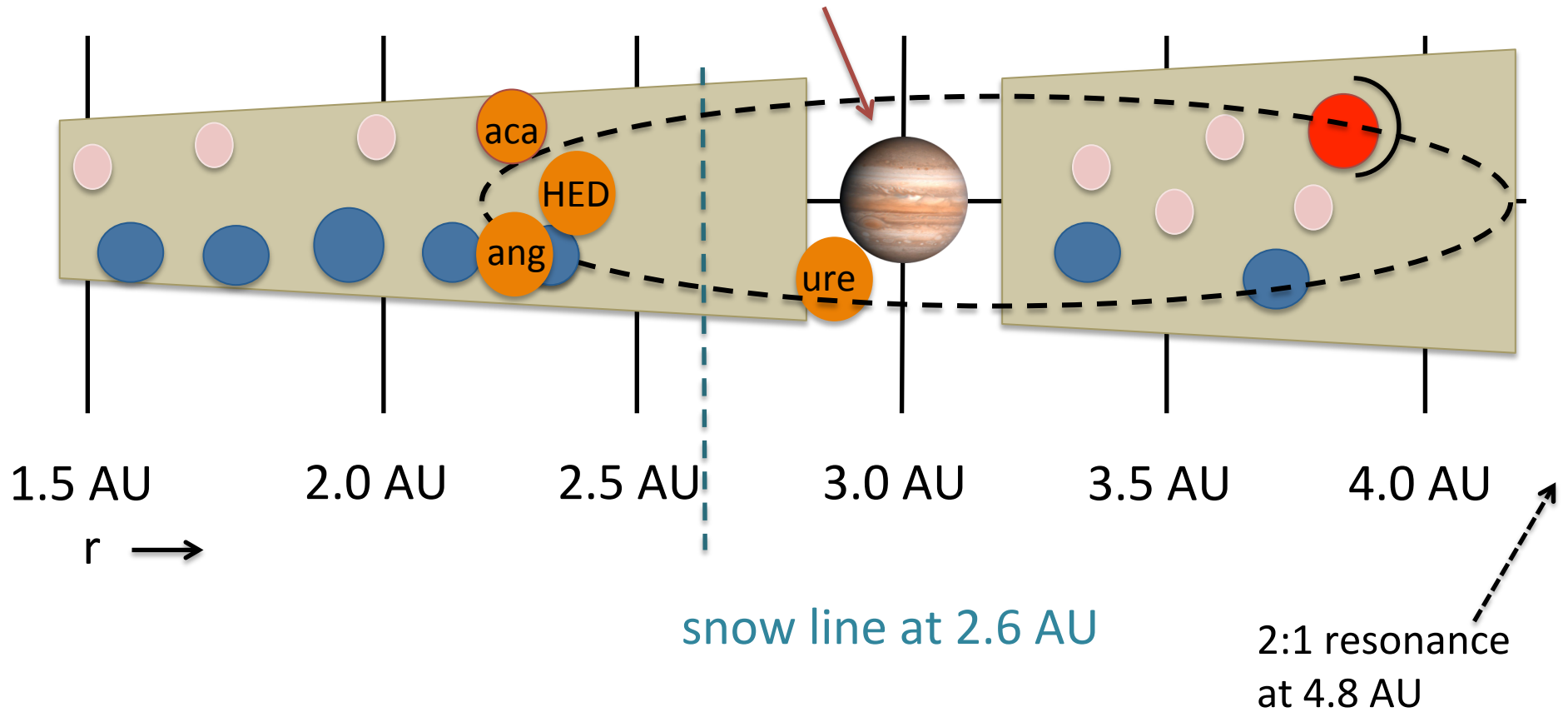
CAIs in outer disk collect in pressure bump at 3.6 AU

Iron meteorites form in outer disk

Planetary embryos form

t = 1.5 Myr

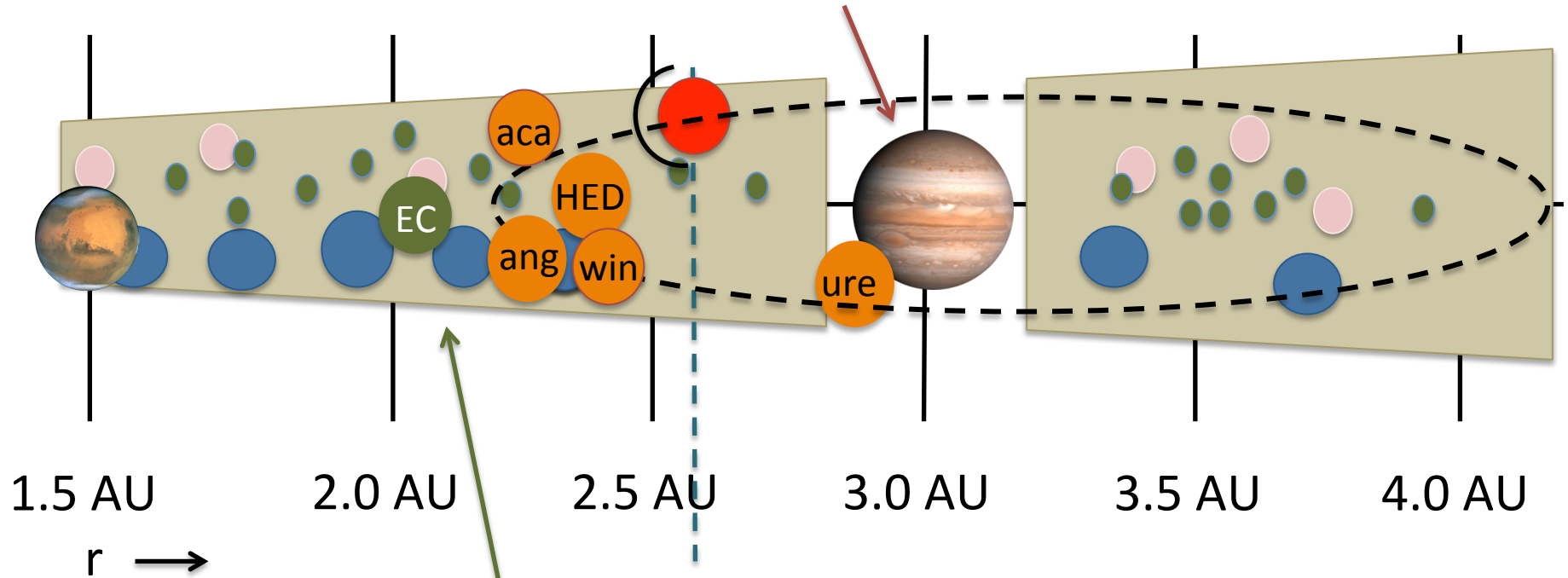
Jupiter continues to grow.
 $M_{\text{Jup}} \sim 125 \text{ ME}$



Embryos near 2:1 resonance driven to high eccentricity, driving bow shocks

t = 2.0 Myr

Jupiter continues to grow.
M_{Jup} ~ 170 ME



1.5 AU

2.0 AU

2.5 AU

3.0 AU

3.5 AU

4.0 AU

r →

Mars
embryo
formed

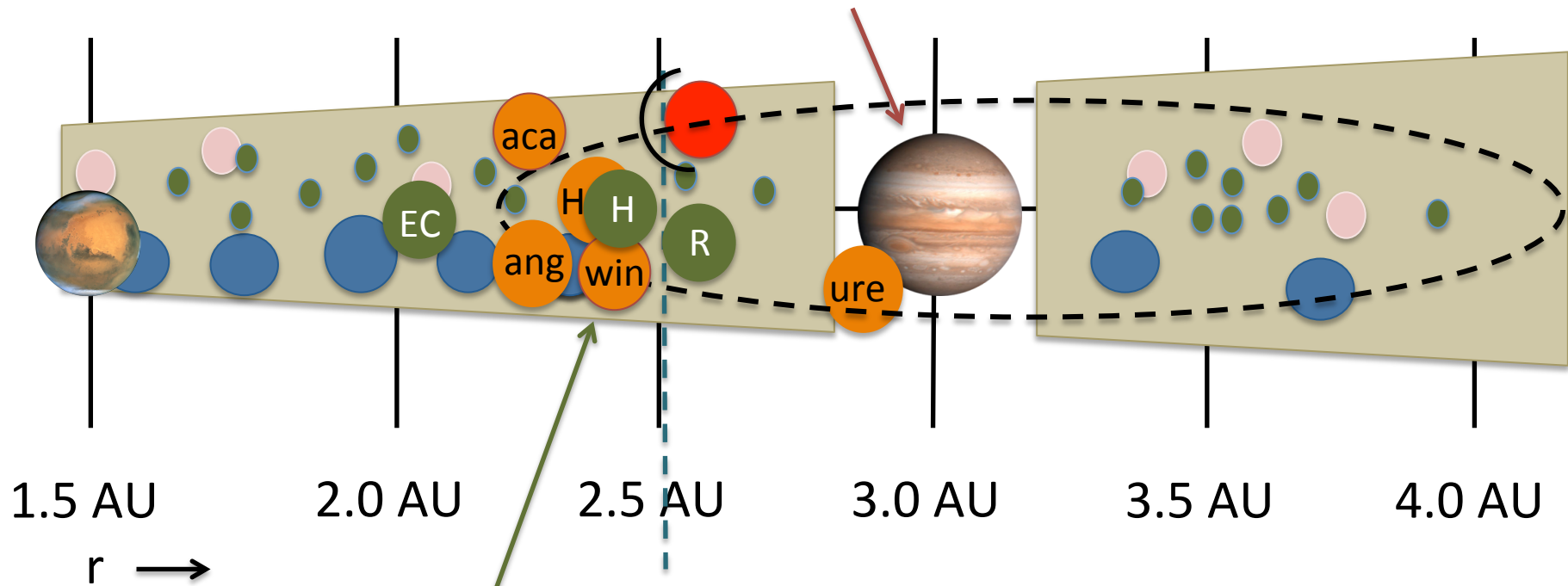
EC parent bodies form

snow line at 2.6 AU

Planetary embryos on eccentric orbits create
chondrules in both inner and outer disk

t = 2.2 Myr

Jupiter continues to grow.
M_{Jup} ~ 175 ME

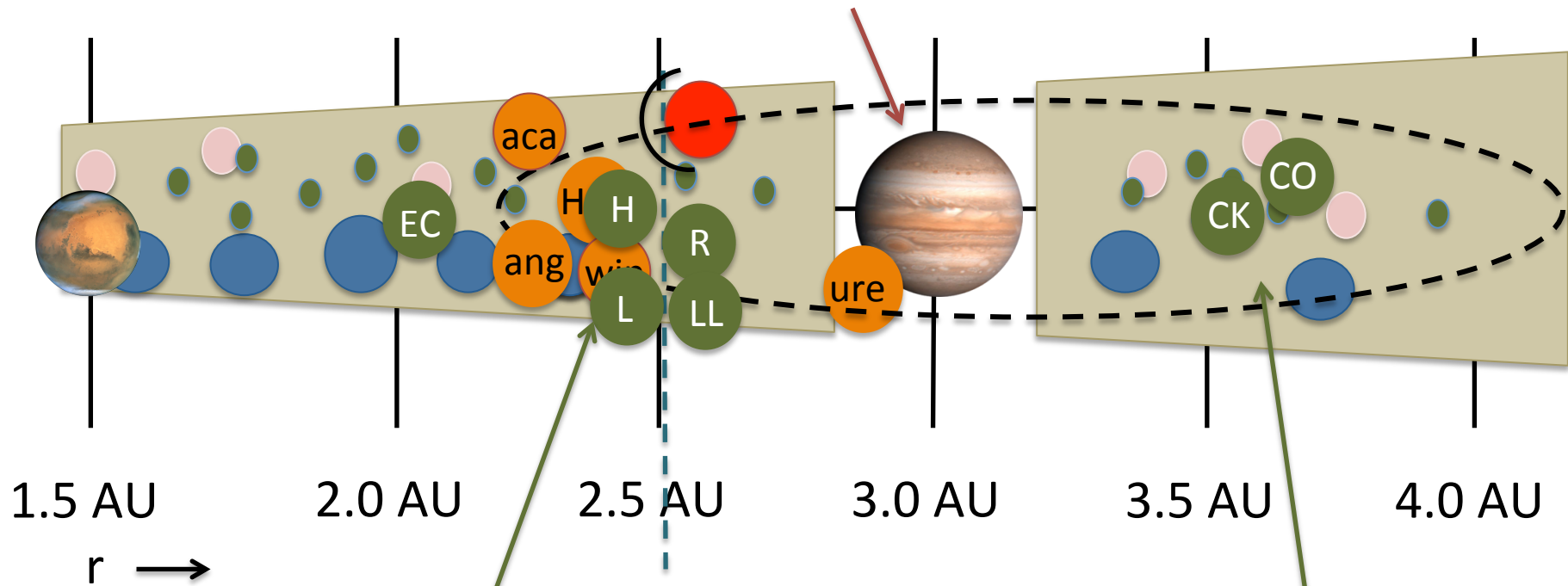


snow line at 2.5 AU

OC parent bodies form

t = 2.5 Myr

Jupiter continues to grow.
M_{Jup} ~ 200 ME



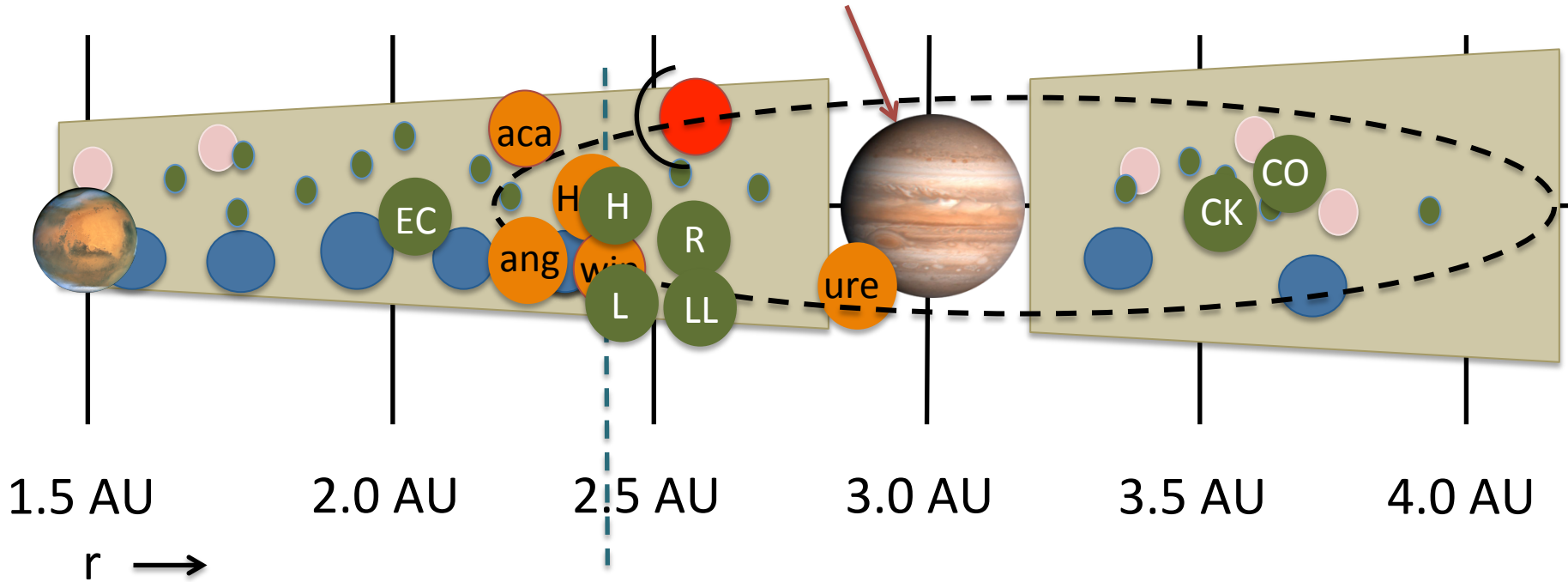
snow line at 2.5 AU

OC parent bodies formed

CK, CO parent bodies form

t = 3.0 Myr

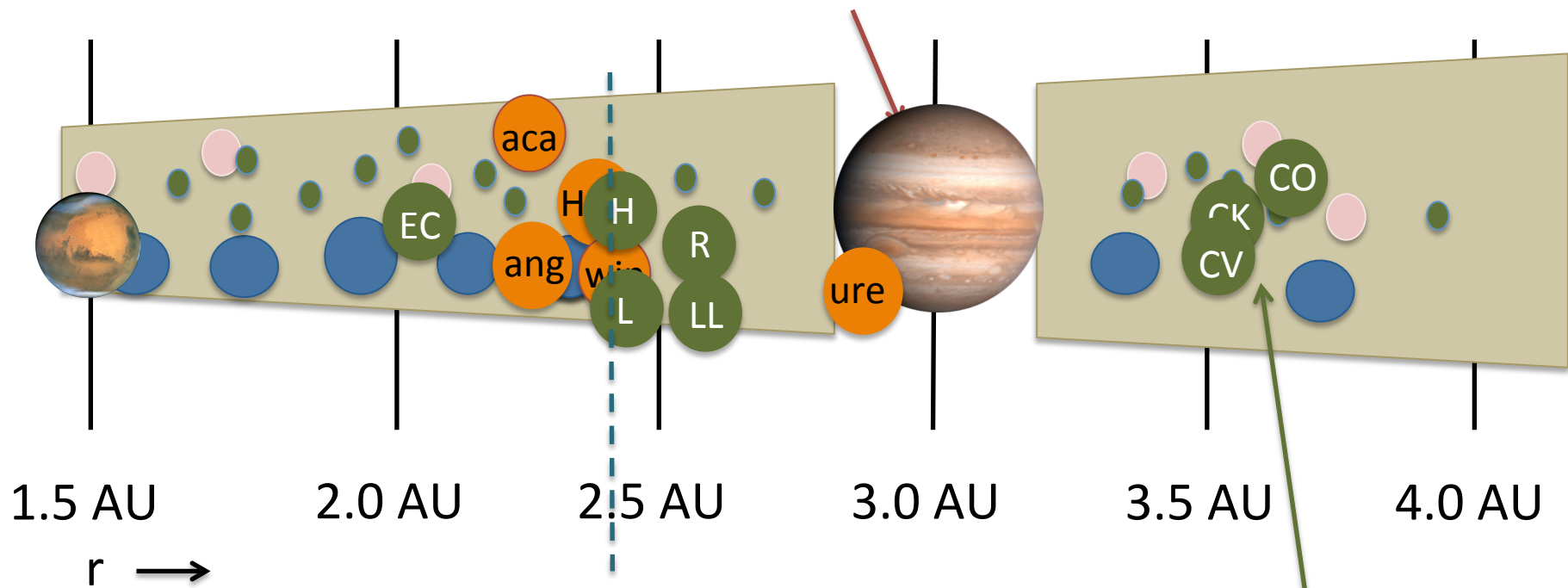
Jupiter continues to grow.
M_{Jup} ~ 230 ME



snow line at 2.4 AU

t = 3.2 Myr

Jupiter continues to grow.
M_{Jup} ~ 240 ME

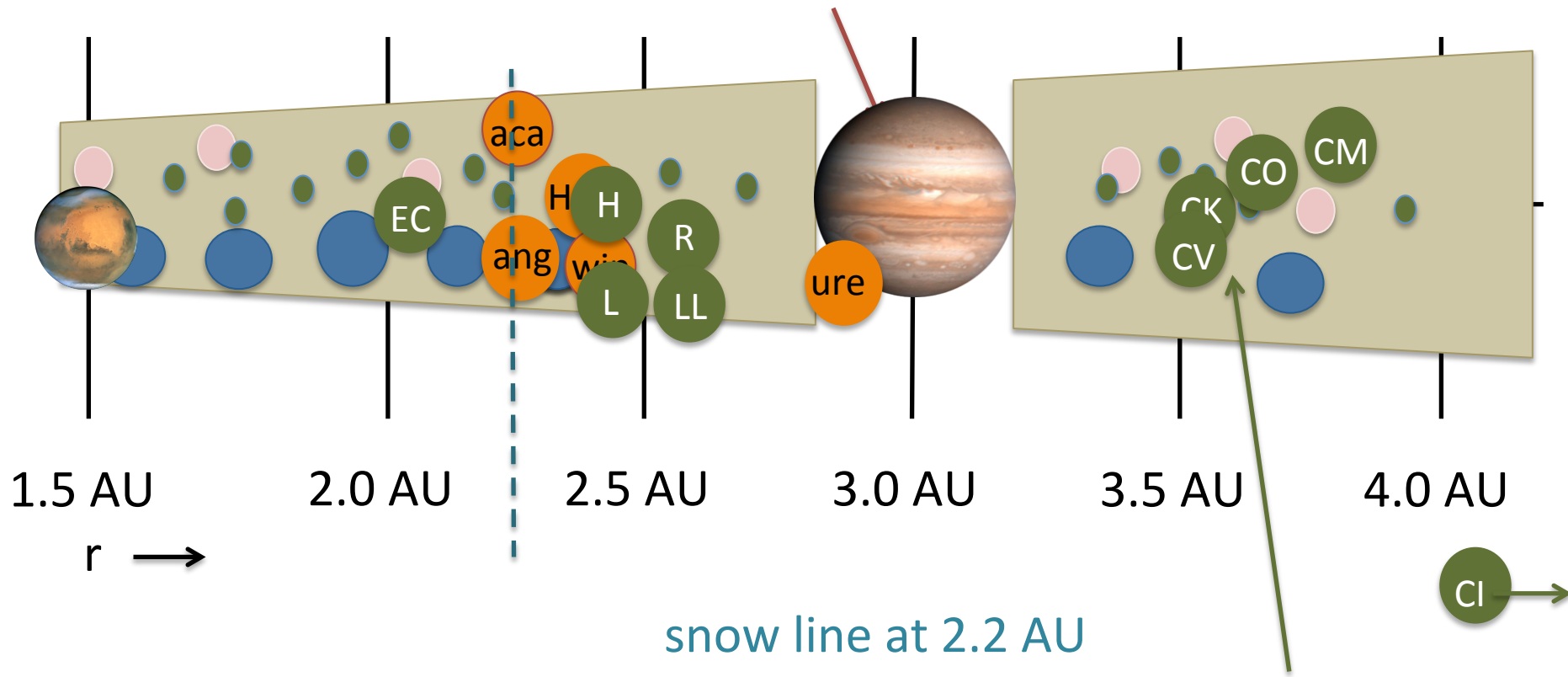


snow line at 2.4 AU

CV parent body forms

t = 3.5 Myr

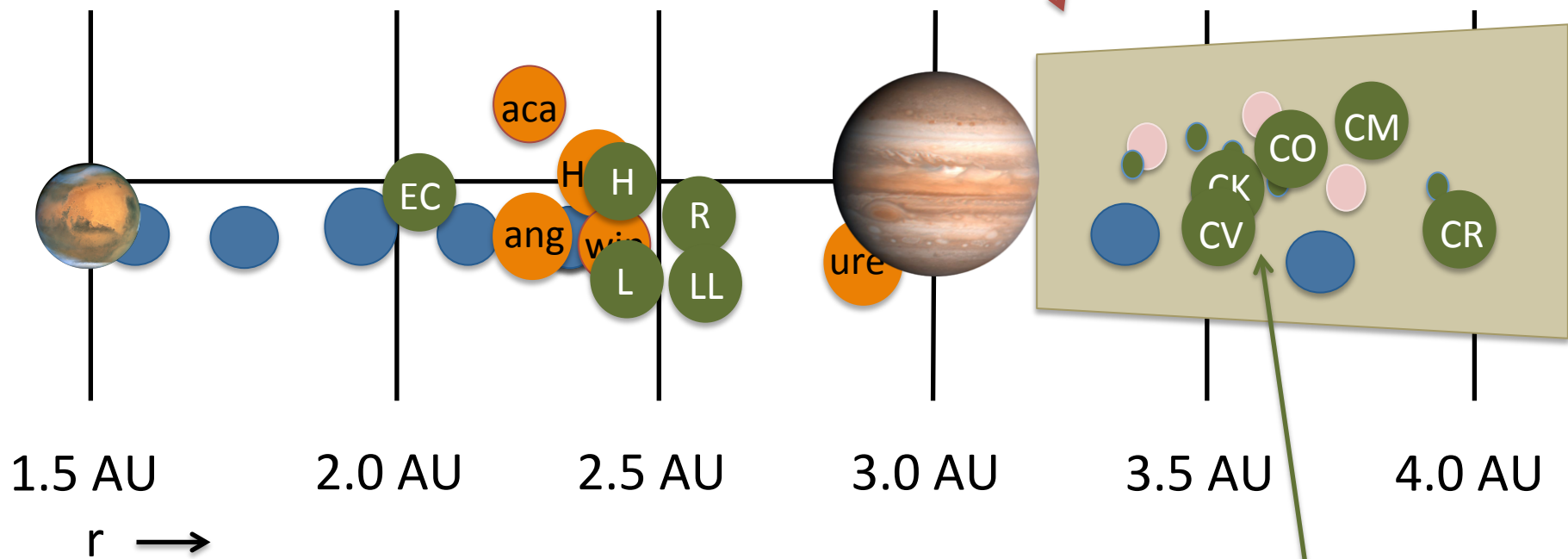
Jupiter continues to grow.
M_{Jup} ~ 260 ME



CM, CI parent bodies
form

t = 4.0 Myr

Jupiter continues to grow.
M_{Jup} ~ 280 ME

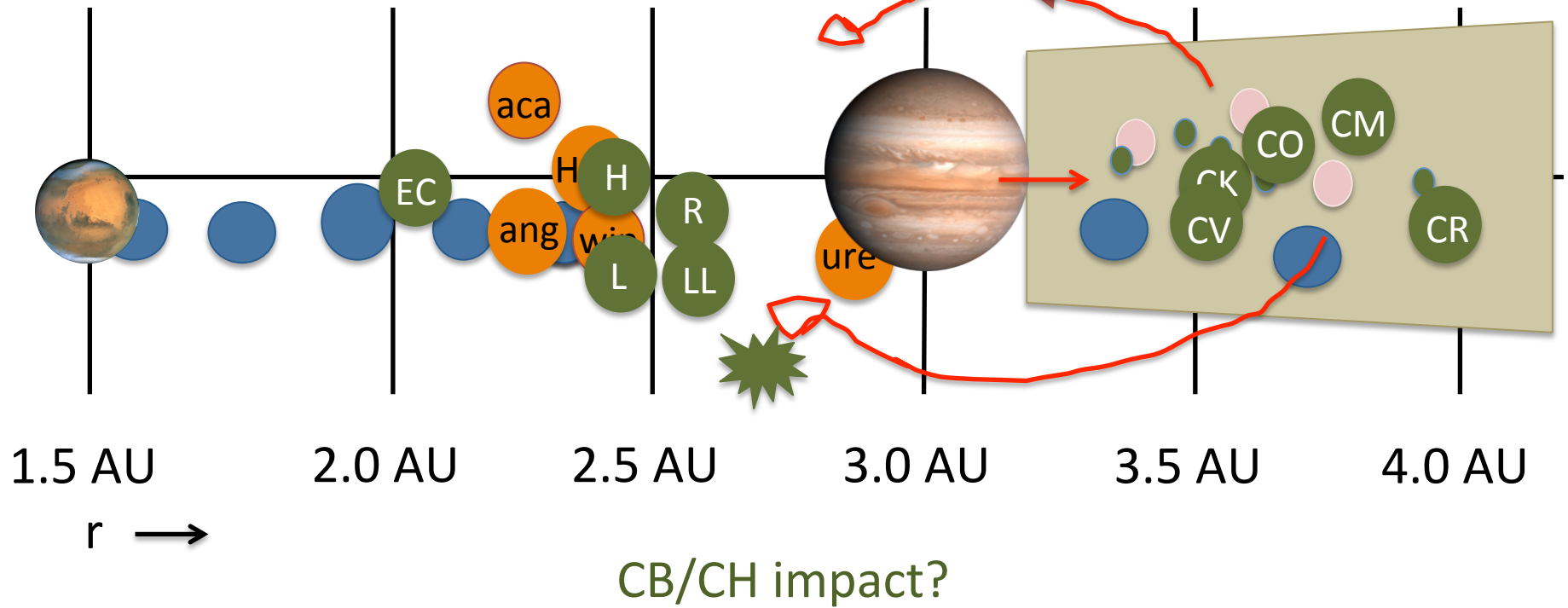


Inner disk gas dissipates

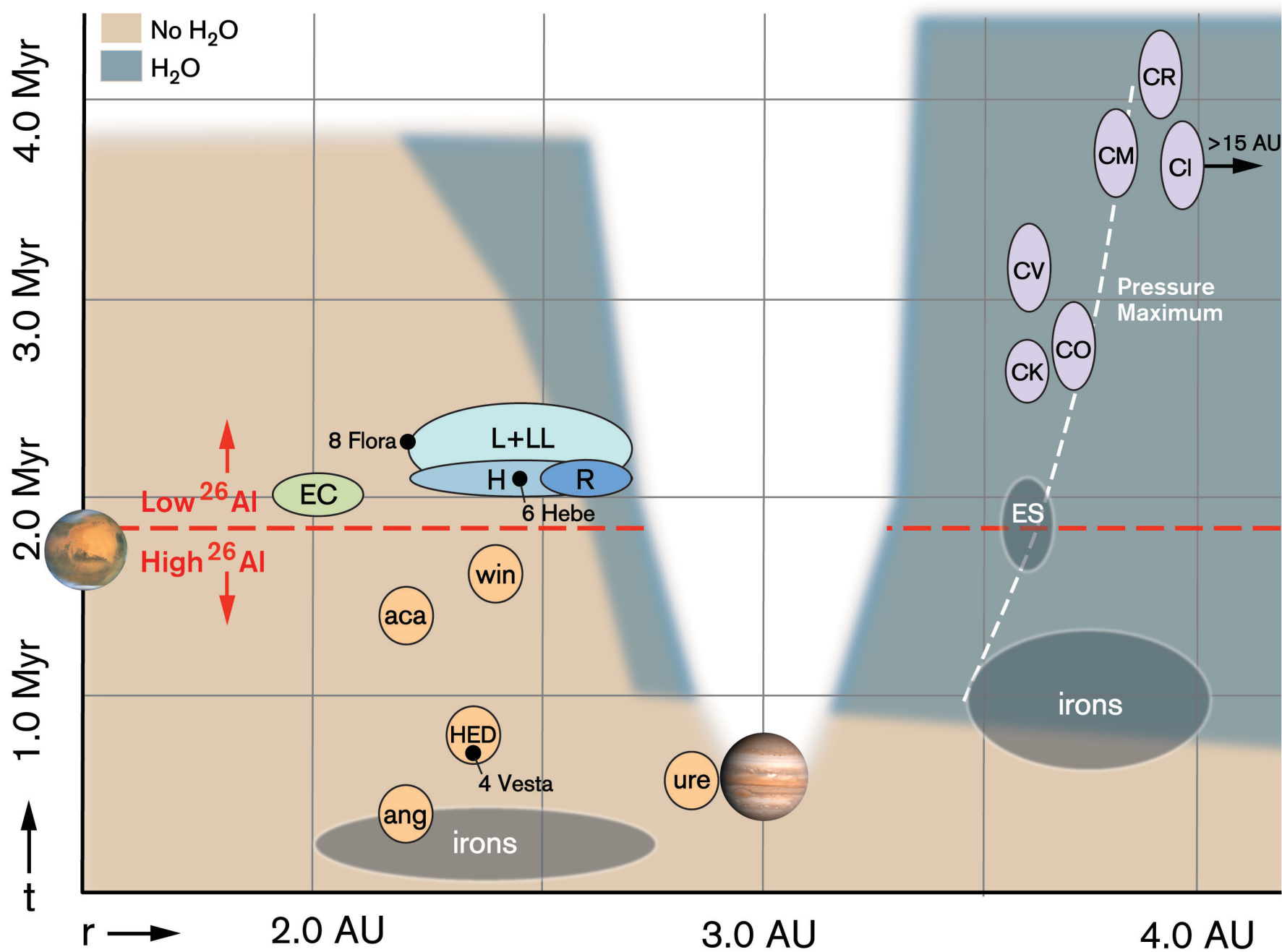
CR parent body forms

t = 4.8 Myr

Jupiter reaches $M_{\text{Jup}} \sim 318 M_E$



Jupiter migrates outward,
scatters parent bodies



A Comprehensive Model

Meteorite Type	Meteoritic Constraints				Model Predictions			
	r (AU)	formation time (Myr)	(X/Mg) / CI	CAIs (vol%)	r (AU)	formation time (Myr)	(X/Mg) / CI	CAIs (wt%)
angrites		~ 0.5	1.35		2.2	0.4	1.30	3.5
HEDs	2.36	~ 0.8	1.09		2.36	0.8	1.09	1.9
acapulcoite-lodranites		~ 1.5	0.90		2.2	1.4	0.90	0.8
ureilites	2.8	~ 0.6			2.8	0.6	1.03	1.4
winonaites		~ 1.8	0.90		2.4	1.8	0.90	0.3
ECs	1.9 - 2.1	1.7 - 1.9	0.87 - 0.88	< 0.1	< 2.1	2.0	0.90	0.27
R		2.0 - 2.2	0.97	< 0.1	2.6?	2.1	0.89	0.03
OCs	2 - 3	2.0 - 2.5	0.89 - 0.90	< 0.2	2.2 - 2.7	2.0 - 2.5	0.88 - 0.90	< 0.2
H	2.43?	2.0 - 2.2	0.90	0.01 - 0.2	2.43	2.1	0.90	0.11
L	2.20?	2.0 - 2.5	0.90	< 0.1	2.2 - 2.7	2.0 - 2.5	0.88 - 0.90	< 0.2
LL		2.0 - 2.5	0.89	< 0.1	2.2 - 2.7	2.0 - 2.5	0.88 - 0.90	< 0.2
CK	beyond	2.4 - 2.8	1.24	< 4	3.6	2.5	1.25	3.0
CV	Jupiter	~ 3.0	1.35	> 3	3.6	3.2	1.34	3.7
CO	"	~ 2.5	1.11	> 1	3.7	2.5	1.11	1.8
CM	"	3 - 4	1.13	1 - 2	3.8	3.5 - 4.0	1.13	2.0
CR	"	3.8 - 4.5	1.02	0.5 - 1	3.9	3.5 - 4.5	1.02	1.0
CI	"	3 - 4	1.00	0.0	> 15	4.0	≡ 1.00	0.0

Meteorite Type	r (AU)	Formation time (Myr)	Σ (g cm^{-2})	ρ ($10^{-10} \text{g cm}^{-3}$)	T (K)
angrites	2.2	0.4	5600	12	410
HEDs	2.36	0.8	4100	8.6	309
acapulcoite-lodranites	2.2	1.4	2700	7.0	245
ureilites	2.8	0.6	4200	7.3	276
winonaites	2.4	1.8	2200	5.5	207
ECs	< 2.1	2.0	> 2000	> 6	> 217
R	2.6?	2.1?	1400	4	< 160
H	2.43	2.1	1980	5.1	190
L	2.2 - 2.7	2.0 - 2.5	< 2000	< 6	100 - 210
LL	2.2 - 2.7	2.0 - 2.5	< 2000	< 6	100 - 210
CK	3.6	2.5	1800	3.1	> 132
CV	3.6	3.2	1500	2.7	> 119
CO	3.7	2.5	1800	3.0	> 128
CM	3.8	3.5 - 4.0	1400	2.4	110
CR	3.9	3.5 - 4.5	1300	2.2	103
CI	~ 15	4.0	6.4	0.002	≈ 44

Conclusions

CAI Storage problem solved: CAIs that diffused past Jupiter trapped in pressure bump created by Jupiter forming at 3 AU at 0.6 Myr. CAIs in 2-3 AU region lost to Sun.

Refractory abundances and CAI abundances can be used to determine where and when different meteorite formed.

Chondrites do appear to be snapshots in time of solar nebula.

H chondrite parent body formed near 2.43 AU, where 6 Hebe orbits. HEDs consistent with formation at 2.36 AU, where 4 Vesta orbits. Many ordinary and enstatite chondrite parent bodies migrated little. Carbonaceous chondrite parent bodies migrated ~ 1 AU.

Conclusions

Almost all carbonaceous chondrites formed in pressure bump beyond Jupiter, a good place to grow planetesimals!
Supports idea of common CK/CV/Eagle Station pallasite parent body (Elkins-Tanton et al. 2011).

CI Chondrites had to form beyond chondrule formation (> 7 AU), but where densities were not too low (< 10 AU), and late (3-4 Myr).
Pressure bump beyond Saturn??

Mixing inevitable. We predict CIs depleted in refractory lithophiles (Ca, Al, Ti, Sc, Y, Zr, Hf, REEs) and probably refractory siderophiles, relative to Sun, by about **9%** (0.04 dex).

Hints of this in data from Lodders et al. (2009)!

Conclusions

Planet formation was extremely rapid (Jupiter in < 1 Myr, embryos in ~ 1 Myr). Radial drift was extremely important.

$\alpha < 10^{-5}$ in disk beyond 10 AU suggests VSI or non-magnetic accretion. Higher $\alpha \sim 10^{-4}$ inside 10 AU suggests magnetic disk winds

Lots left to do: incorporate formation of disk (like Yang & Ciesla 2012), calculate water distribution, better assess chondrule formation, etc...

But careful consideration of constraints from CAIs and chondrules can provide unparalleled insights into disk processes and planet formation!

What melted chondrules?

As reviewed by Desch et al. (2012):

- Peak temperatures > 1800 K for minutes
- Cooling rates $\sim 10 - 10^3$ K/hr for hours.
- In the presence of gas, and cogenetic with matrix grains
- Densities of chondrules $\sim 10 \text{ m}^{-3}$ (suppression of evaporation, compound chondrule frequency)
- Repeatable process (relict chondrules)

Large-scale nebular shocks (Desch & Connolly 2002; Ciesla et al. 2002)

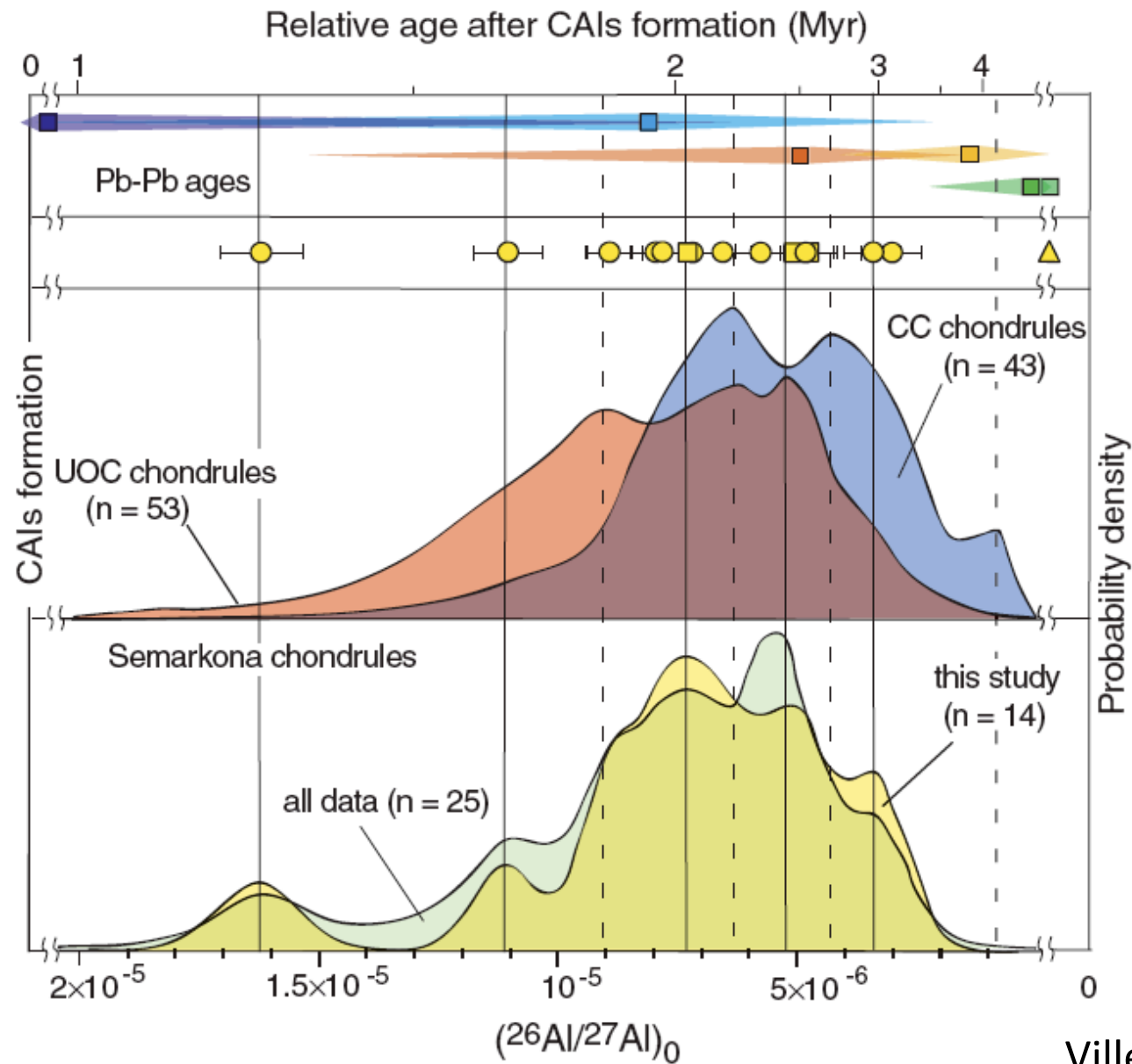
Timing and sources?

Bow shocks around planetary embryos (Morris et al. 2012, Boley et al. 2013, Mann et al. 2016). Can cooling rates match?

Impact melt droplets (Johnson et al. 2015; Lichtenberg et al. 2017).

Geochemistry? Cooling rates? Efficiency?

Most chondrules formed ~ 1.5 – 3.0 Myr after CAIs



Villeneuve et al. (2009)

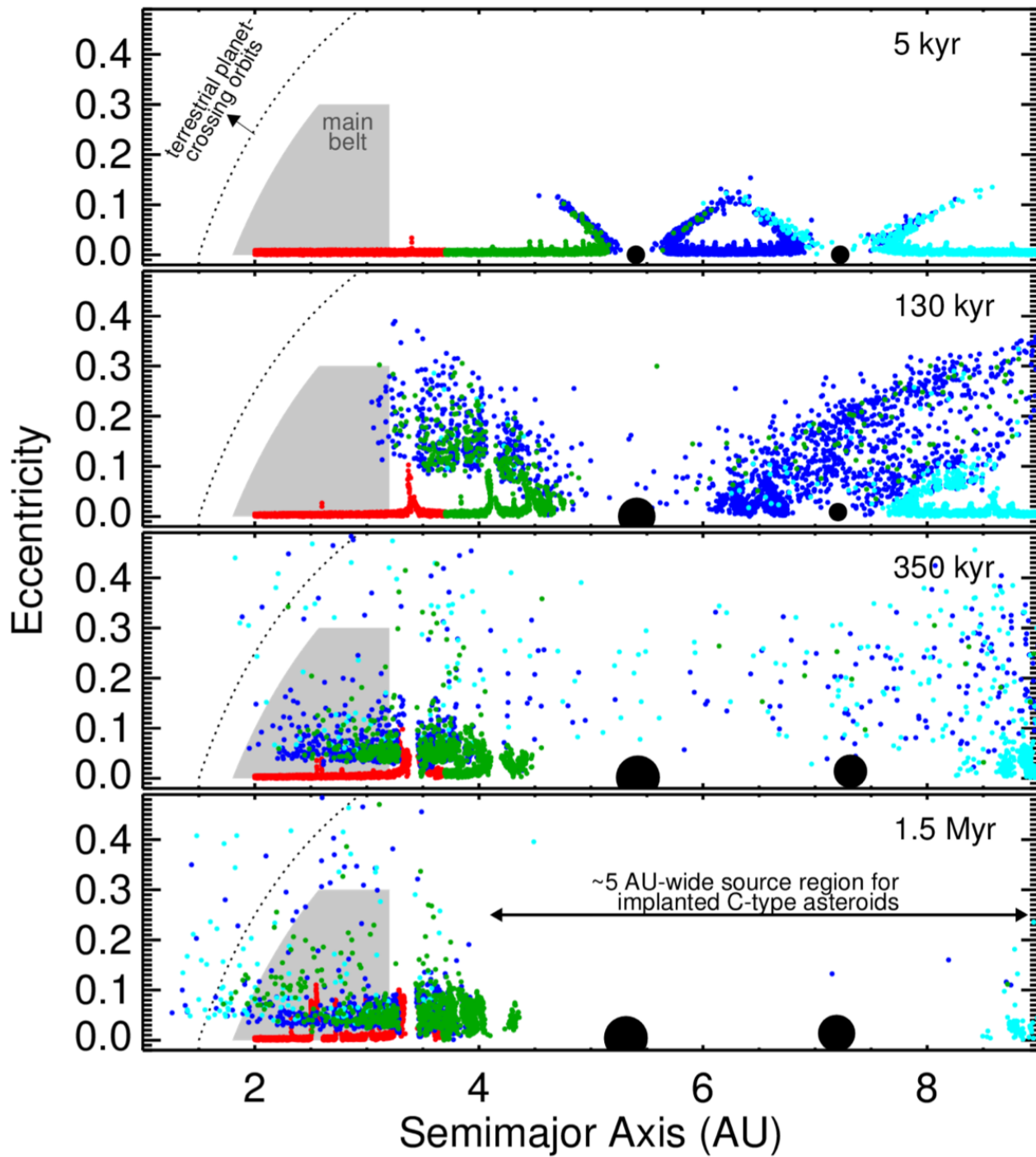
A Comprehensive Model

3 AU



2:1 resonance at 4.76 AU
 $e = 0.4, 2.9 - 6.7$ AU

Not only can Jupiter trap CAIs, it
can trigger chondrule formation



Raymond & Izidoro
(2017)

What melted chondrules?

Requires $e \sim 0.4$ and $\rho \sim 10^{-9} \text{ g cm}^{-3}$ (Boley et al. 2013).

